

**REPORT
74**



MINERAL OCCURRENCES AND EXPLORATION POTENTIAL OF THE EAST KIMBERLEY

by L. Y. Hassan



GEOLOGICAL SURVEY OF WESTERN AUSTRALIA
DEPARTMENT OF MINERALS AND ENERGY



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Cover photograph: Argyle AK1 mine and processing plant (photograph courtesy of Argyle Diamonds)

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Mineral occurrences and exploration potential of the east Kimberley

by

L. Y. Hassan

Abstract

Gold was discovered in the east Kimberley in 1885 and small quantities of gold have been produced intermittently since that time. Modern exploration of the area commenced in the 1950s and reached a peak during the late 1970s and early 1980s following the discovery of the world-class diamond deposit at Argyle. This exploration resulted in the discovery of platinum group element (PGE), nickel, rare earth element (REE), niobium, tantalum, fluorite, and base metal deposits.

The oldest tectonic units in the east Kimberley are the Palaeoproterozoic Lamboo Complex and Granites–Tanami Complex. The Lamboo Complex formed as a result of convergence and collision between the Kimberley and North Australian Cratons. Three north–northeasterly trending terranes have been identified in the Lamboo Complex: Eastern zone, Central zone, and Western zone. The Granites–Tanami Complex probably formed within the North Australian Craton. The complexes are overlain by a number of sedimentary basins, which range in age from Palaeoproterozoic to Devonian. The Halls Creek Orogen, which includes the Lamboo Complex and the deformed parts of various sedimentary basins, records periodic tectonism until the Carboniferous.

The three tectonostratigraphic zones of the Palaeoproterozoic Lamboo Complex show quite different styles and types of mineralization. Most of the gold mineralization is in the Eastern zone, where it is concentrated in the mafic volcanic and sedimentary rocks of the Biscay Formation, or in the immediately overlying sedimentary rocks of the Olympio Formation and its alkaline volcanic members (Butchers Gully Member and Maude Headley Member). In the Eastern zone, there is also significant REE–niobium–tantalum–tin–beryllium mineralization hosted by alkaline volcanic rocks of the Butchers Gully Member; base metal mineralization in distal volcanoclastic sedimentary rocks within the Biscay Formation; and stratabound tungsten mineralization, possibly volcanic exhalative in origin, hosted by the Ding Dong Downs Volcanics, Saunders Creek Formation, Biscay Formation, and Butchers Gully Member. In the Central zone, volcanogenic massive sulfide (VMS) deposits are hosted by the Koongie Park Formation, and possible porphyry copper mineralization is hosted by the Angelo Microgranite and the McHale Granodiorite. Layered mafic intrusions host PGE, chromite, and nickel–copper mineralization in the Central zone, and to a lesser extent in the Western zone.

Significant gold mineralization has recently been discovered in the Palaeoproterozoic Granites–Tanami Complex in the southeast of the area. Gold mineralization in the Larranganni area, including the Kookaburra deposit, has strong stratigraphic control, with the occurrences confined to two sequences of folded mafic volcanic and sedimentary rocks that have been informally named the Bald Hill sequence and the Quartz Ridge sequence. The relationship of these sequences to the stratigraphic succession in the Northern Territory is uncertain. Some gold mineralization is also associated with uranium at the unconformity between the Granites–Tanami Complex and the overlying Birrindudu Basin. The controls on mineralization in the newly discovered Coyote deposit are not yet clearly understood.

The c.1200 Ma diamondiferous Argyle AK1 lamproite pipe is related to a splay of the Halls Creek Fault, which is a major tectonic suture that has shown discontinuous movement since the Palaeoproterozoic. The Neoproterozoic Cummins Range Carbonatite (containing REE and tantalum) is located near the southern end of the Halls Creek Fault.

There is also a strong stratigraphic control on mineralization in some of the younger basins. Mississippi Valley-type mineralization is associated with Devonian reef complexes on the Lennard Shelf in the Canning Basin where zinc–lead–silver ore is currently mined at Kapok. Possible Mississippi Valley-type mineralization is also hosted by the Mesoproterozoic Bungle Bungle Formation in the Osmond Basin. There is minor native copper mineralization within amygdaloidal basalts of the Cambrian Antrim Plateau Volcanics. Copper mineralization of possible volcanic exhalative origin is widespread in interflow sedimentary rocks within the Carson Volcanics in the Palaeoproterozoic Kimberley Basin. Subeconomic copper mineralization of probable syngenetic origin is also widespread in the Teronis Member of the Elgee Siltstone and in shale at the base of the Middle Pentecost Sandstone in the Kimberley Basin.

Epigenetic base metal and fluorite mineralization in the Palaeoproterozoic Speewah Basin is associated with faults, possibly related to caldera collapse, and may be as young as Tertiary.

KEYWORDS: mineralization, exploration, Halls Creek Orogen, Lamboo Complex, Canning Basin, Granites–Tanami Complex, diamonds, gold, base metals, platinum group elements, nickel, chromite, tungsten, tin, rare earth elements, tantalum, niobium, beryllium, fluorite, uranium.

Introduction

Until the discovery of diamonds at Argyle, and the commencement of mining of quartz-stockwork vein gold mineralization at Palm Springs, the east Kimberley area had the reputation for hosting scattered, small, non-economic mineral occurrences. The aim of this study is to show the distribution of reported mineralization and exploration activity in a GIS database, to relate the mineralization to the geological setting, and to highlight the prospectivity of the east Kimberley area for a range of different commodities and mineralization styles.

Details of mineral exploration, mineral occurrences, and other geoscientific information for the study have been compiled from the following sources:

- the large number of open-file statutory mineral exploration reports held in the Western Australian mineral exploration (WAMEX) database at the Department of Minerals and Energy (DME);
- the database of Western Australia's mines and mineral deposits information (MINEDEX) held at DME;
- information published in books, journals, company reports to shareholders, and on the internet;
- regional geological surveys, airborne geophysical datasets, and remote-sensing datasets.

The mineral prospectivity study of the east Kimberley has three main parts: this Report, Plates 1 and 1a of this Report, and a digital dataset on CD-ROM. This Report reviews the regional geology of the area, the history of mining and exploration, examples of mineral occurrences for each style of mineralization, the mineralization controls, and the potential for further discoveries. Where mineral occurrences are mentioned in the Report for the first time, they are also identified by the WAMIN 'deposit number' shown thus: Palm Springs (2223). It should be noted that the WAMIN deposit numbers given in this Report differ from WAMIN deposit numbers in Sanders (1999) because the database was moved from MS Access to an Oracle platform between the writing of the two reports.

Plates 1 and 1a show the mineral occurrences, indicating commodity and mineralization style, on a geological map showing solid geology with a simplified regolith overprint at 1:500 000 scale. The key to the mineral occurrences on Plates 1 and 1a is provided in Appendix 1.

The accompanying CD-ROM includes all the data used to compile the map and Report, and also includes files of geophysical, remote-sensing, and topographic data. The CD-ROM contains the files necessary for viewing the data in the ArcView GIS environment, and a self-loading version of the ArcExplorer software package modified to suit this particular dataset. Metadata statements on the geological, geophysical, and topographic datasets are also provided.

Appendix 2 defines the terms used in the Geological Survey of Western Australia (GSWA) mineral occurrence database (WAMIN) and exploration activity database

(EXACT), and Appendix 3 gives a description of the digital datasets included on the CD-ROM.

Location, land access, and physiography

This Report covers the east Kimberley area of Western Australia, which includes the LISSADELL*, DIXON RANGE, GORDON DOWNS, MOUNT RAMSAY, MOUNT BANNERMAN, and BILLILUNA 1:250 000 map sheets (Fig. 1).

The sealed Great Northern Highway runs through the area in an east-northeasterly direction. The other major access road in the area is the unsealed Duncan Highway, which joins the Great Northern Highway at Halls Creek, the major town in the area (population about 1300 in 1995). Station tracks also provide access to many parts of the area. More remote areas, especially in the southern part of the area, are best accessed by helicopter.

Pastoral stations, conservation reserves, and reserves for the use and benefit of aboriginal inhabitants occupy a large percentage of the area of the east Kimberley. Standard guidelines have been developed to cover mineral exploration and mining in the different types of conservation reserves, proposed conservation reserves, and other environmentally sensitive areas. These guidelines are summarized in the DME Information Series pamphlet no. 11 (Department of Minerals and Energy, 1995).

As well as the Mining Act 1978, all new tenements and exploration licence renewals are also subject to the legislation and procedures of the Commonwealth Native Title Act 1993, except where it is determined that the applications cover land where native title has been extinguished.

Most of the project area is a wide belt of rough country and low rounded hills between the Kimberley Plateau to the west and the Victoria River Plateau to the east. The southern part of the area is an almost flat plain forming part of the Sturt Plateau and Canning Plain. More detailed descriptions of the physiography are given in Dow and Gemuts (1969), Beard (1979), and Warren (1994a).

Previous work

The first geological mapping in the East Kimberley area was carried out by Hardman (1885) who also reported on the potential for gold and other minerals. Woodward (1891) and Smith (1898) gave detailed descriptions of the gold workings. Talbot (1910) contributed to the knowledge of the geology of the area in his traverse of the country between the Tanami in the Northern Territory and Halls Creek. The first detailed maps of some of the gold mining centres and base metal and tin prospects were made by the Aerial, Geological and Geophysical Survey of Northern Australia (AGGSNA) in the late 1930s (Jones, 1938; Finucane, 1938a,b, 1939a,b; Finucane and Sullivan,

* Capitalized names refer to standard map sheets

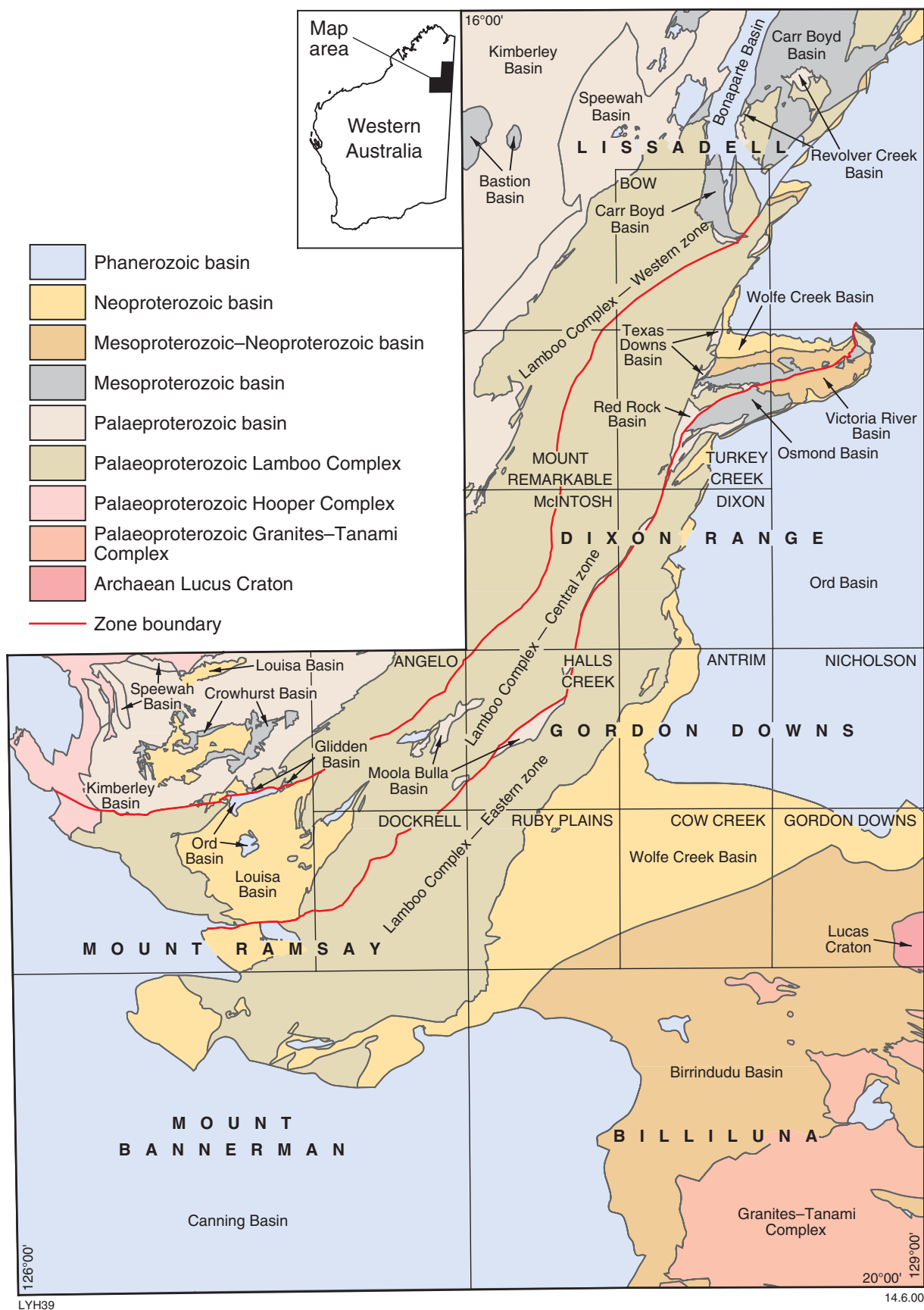


Figure 1. Tectonic sketch map of the east Kimberley showing the tectonostratigraphic subdivision of the Lamboo Complex, the Granites–Tanami Complex, the Proterozoic and Phanerozoic basins, and published 1:100 000- and 1:250 000-scale map sheets (e.g. ANTRIM). Scale — 1 cm : 20 km (approximately)

1939). Harms (1959) described the geology of the area and also gave descriptions of the gold, iron ore, chromium, uranium, base metals, tin, tungsten, and other mineral occurrences.

Geological mapping of the east Kimberley, at a scale of 1:250 000, was carried out jointly by the GSWA and the Bureau of Mineral Resources (BMR; now the Australian Geological Survey Organisation, AGSO) between 1962 and 1977. Maps include DIXON RANGE (Dow and Gemuts, 1967), LISSADELL (Plumb, 1968), GORDON DOWNS (Gemuts and Smith, 1968), MOUNT RAMSAY (Roberts et al., 1968), BILLILUNA (Blake et al., 1977), and MOUNT BANNERMAN (Yeates and Muhling, 1977). Dow et al. (1964) describes the geology of the Ord River area. A Bulletin on the geology of the east Kimberley and its mineral occurrences, with an accompanying 1:500 000 colour map, was compiled by Dow and Gemuts (1969), and Gemuts (1971) describes the metamorphic and igneous rocks of the Lamboo Complex.

During 1990–1995, part of the area was remapped at scales of 1:100 000 and 1:250 000 by GSWA and AGSO as part of the National Geoscience Mapping Accord (NGMA) Kimberley–Arunta Project. This mapping program included geochronological, whole-rock geochemical, petrological, and remote-sensing studies. The map sheets covered by this program are shown on Figure 1. Explanatory notes and maps have been published for the following 1:100 000 sheets: MOUNT REMARKABLE (Sheppard et al., 1997a), DOCKRELL (Tyler et al., 1998a), RUBY PLAINS (Blake et al., 1997), ANGELO (Griffin et al., 1998), and HALLS CREEK (Blake et al., 1999b). Maps only have been published for the following 1:100 000 sheets: ANTRIM (Blake and Warren, 1996), BOW (Sheppard et al., 1997b), TURKEY CREEK (Tyler et al., 1997a), MCINTOSH (Tyler et al., 1997b), and DIXON (Tyler et al., 1998b). Preliminary notes on the geology of DIXON, southeast MCINTOSH, and northernmost HALLS CREEK are in Warren (1997). Explanatory notes and a map have also been published for LISSADELL (1:250 000; Thorne et al., 1999), and maps for MOUNT RAMSAY (Tyler et al., 1998c) and DIXON RANGE (Tyler et al., 1998d). The Devonian Reef complexes of the Canning Basin on the western side of MOUNT RAMSAY have been mapped at a scale of 1:100 000 (Hocking and Playford, 1998a,b) and at 1:50 000 (Playford and Hocking, 1998).

Aeromagnetic and gravity data at a scale of 1:250 000 have been published by AGSO and are also available in digital format.

The Department of Resources Development and Department of Minerals and Energy (1997) reported on the prospectivity of the area as part of a joint study. Sanders (1999) discusses the mineralization of the Halls Creek Orogen and describes many of the mineral occurrences in detail; the author of this Report was involved in compiling Plate 1 (Sanders and Hassan, 1999) of Sanders (1999).

Regional geology

The east Kimberley area covers most of the Halls Creek Orogen, as well as the eastern end of the King Leopold

Orogen, part of the Granites–Tanami Complex, and parts of the Speewah, Kimberley, Ord, Bonaparte, Canning, and Birrindudu Basins (Fig. 1).

The solid geology map on Plate 1, and the digital solid geology map accompanying this Report, are based on the NGMA 1:100 000-scale geological maps where these are available, and on 1:250 000 scale geological maps. In the case of BILLILUNA and MOUNT BANNERMAN, where there is no recent mapping, there has been some interpretation based on open-file and published company data, and aeromagnetic and Landsat TM images. The regolith layer is also based on the 1:100 000-scale and 1:250 000-scale geological maps.

Halls Creek Orogen

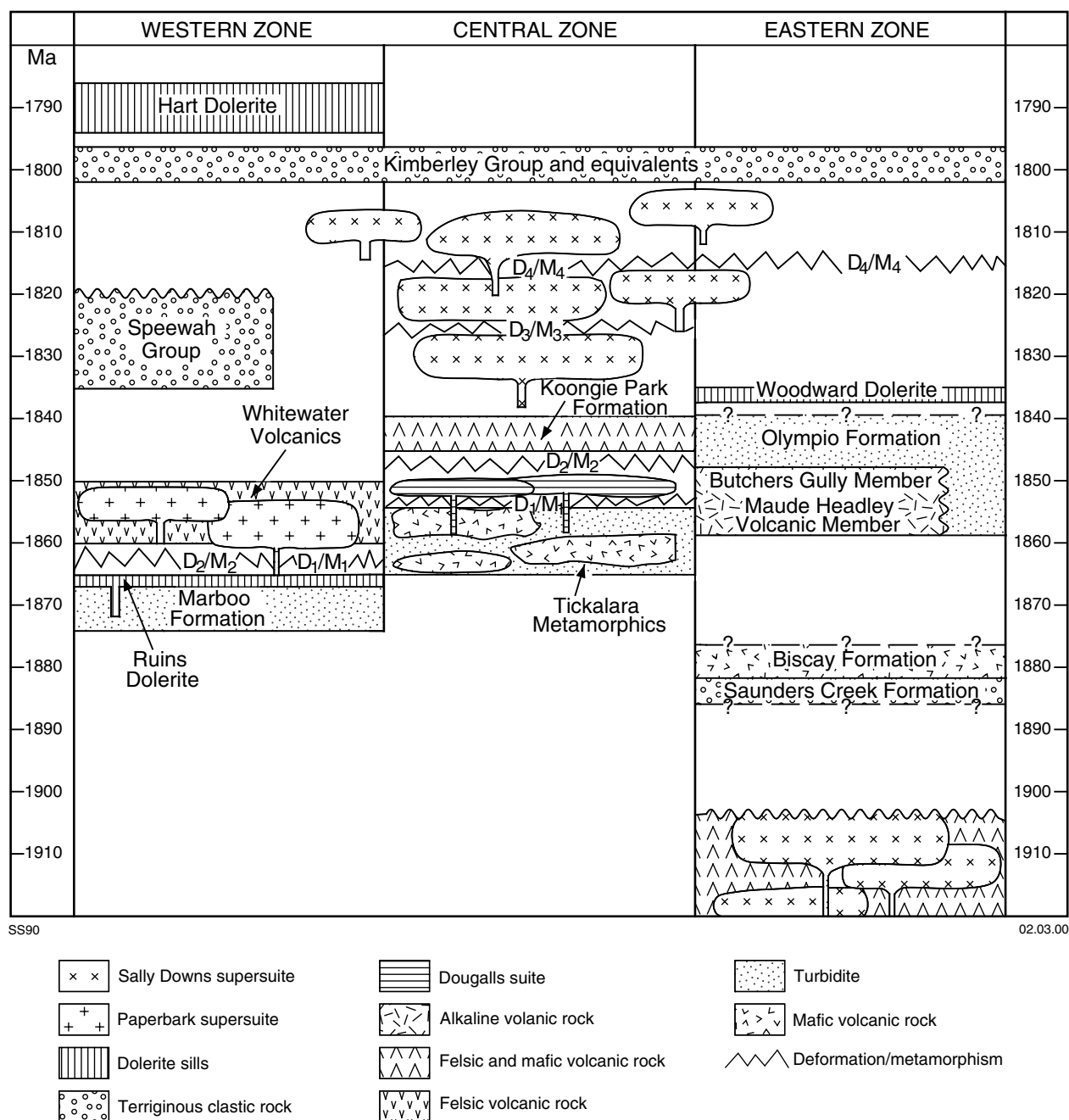
Development of the Halls Creek Orogen between the Kimberley Craton (which underlies the Kimberley and Speewah Basins) to the northwest and the North Australian Craton to the east commenced during the Palaeoproterozoic (Sheppard et al., 1999). Intermittent deformation took place until the end of the Palaeozoic (Tyler et al., 1995). The plutonic rocks, and deformed and metamorphosed Palaeoproterozoic sedimentary and volcanic rocks, are referred to as the Lamboo Complex (Gemuts, 1971; Griffin and Grey, 1990a). The folded margins of the Speewah and Kimberley Basins, and the deformed parts of sedimentary basins resulting from later tectonism during the Proterozoic and Phanerozoic are also included in the Halls Creek Orogen. A summary of the geological history of the Halls Creek Orogen is given in Table 1.

Palaeoproterozoic Lamboo Complex

Tyler et al. (1995) recognized three zones in the Lamboo Complex, each with a distinct stratigraphy, and metamorphic and deformation history (Fig. 2; Table 1). The Eastern zone is separated from the Central zone by the Angelo – Halls Creek – Osmond Fault System. The Central and Western zones are separated by a combination of the Ramsay Range Fault and the Springvale Fault; in the northern part of the Lamboo Complex, this boundary is largely obscured by granitoid. The three zones have been interpreted as tectonostratigraphic terranes that were brought into their current positions by a combination of subduction and large-scale strike-slip faulting by c. 1820 Ma, when they were all intruded by plutons of granite and gabbro (Tyler et al., 1995; Sheppard et al., 1999). The different geophysical nature of the crust on either side of the Lamboo Complex (Shaw et al., in press), and the presence of terranes, is indicative of a Palaeoproterozoic plate margin (Tyler et al., 1999). This contrasts with earlier models that described the Lamboo Complex as intracratonic on the basis of correlation of rock units across the Halls Creek Fault (Hancock and Rutland, 1984; Page and Hancock, 1988).

Eastern zone

In the Eastern zone, low- to medium-grade metasedimentary and meta-igneous rocks of the c. 1880–1840 Ma



Halls Creek Group unconformably overlies c. 1910 Ma mafic and felsic volcanic rocks of the Ding Dong Downs Volcanics, and c. 1910 Ma granitoids of the Sophie Downs suite (Griffin and Tyler, 1992; Blake et al., 1999b). The Ding Dong Downs Volcanics (*Pv*) include metamorphosed amygdaloidal basalt, metamorphosed rhyolite, and minor intercalated metasedimentary rocks. The Sophie Downs suite includes the Sophie Downs Granophyre Member (*Pvg*), the Esaw Monzogranite, and the Junda Microgranite (*Pgv*; Tyler et al., 1998a).

The Halls Creek Group as redefined by Griffin and Tyler (1992) consists of, in ascending order, the Saunders Creek Formation, Biscay Formation, and Olympio Formation.

The Saunders Creek Formation (*PHs*) consists predominantly of sandstones and conglomerates (Dow and Gemuts, 1969; Tyler et al., 1995). The sediments were deposited in a braided fluvial system with a lower energy, transitional marine environment towards the top (Hancock, 1991). Felsic, lithic tuffs have been recognized in the Castle Creek Anticline area (Buckovic et al., 1982) on HALLS CREEK and ANTRIM.

The Biscay Formation (*PHr*) conformably overlies the Saunders Creek Formation and consists predominantly of metamorphosed massive basaltic lava flows, local pillow lavas, laminated mafic volcanoclastic deposits, and fragmental deposits. The mafic volcanic rocks are inter-

Table 1. Summary of the geological history of the Halls Creek Orogen

Age (Ma)	Kimberley Basin	Lamboo Complex		
		Western zone	Central zone	Eastern zone
2500–1910 c. 1920–1910	unexposed Kimberley Craton			<ul style="list-style-type: none"> • > 2500 Ma North Australian Craton • Eruption of the Ding Dong Downs Volcanics • Intrusion of granitoids of the Sophie Downs suite • Uplift and erosion. Deposition of Saunders Creek Formation and Biscay Formation
c. 1880				
c. 1870		<ul style="list-style-type: none"> • Deposition of Marboo Formation 		
1865–1850		----- HOOPER OROGENY -----		
c. 1865 1863 1865–1850		<ul style="list-style-type: none"> • D₁ recumbent folding – M₁ • Deposition of Whitewater Volcanics • Intrusion of Greenvale and Castlereagh Hill Porphyries • Intrusion of layered mafic–ultramafic intrusions (including Springvale and Wilagee intrusions) • Intrusion of granitoid and gabbro of the Paperbark supersuite • Contact metamorphism of Marboo Formation 	<ul style="list-style-type: none"> • Deposition of protoliths of the Tickalara Metamorphics • Intrusion of Rose Bore Granite • Intrusion of massive gabbro, and layered mafic–ultramafic intrusions including Panton intrusions • Intrusion of protolith of felsic gneiss • Early thrusting (D₁) and medium- to high-grade metamorphism (M₁) 	<ul style="list-style-type: none"> • Deposition of lower part of the Olympio Formation • Eruption of Maude Headley Member
c. 1850		<ul style="list-style-type: none"> • D₂ upright folding – M₂ 	<ul style="list-style-type: none"> • Deformation and metamorphism of Tickalara • Metamorphics under high-T, low- to moderate-P metamorphic conditions (D₂–M₂) • Intrusion of gabbro–diorite and tonalite–trondhjemite sheets of the Dougalls suite 	<ul style="list-style-type: none"> • Eruption of Butchers Gully Member
<1847				<ul style="list-style-type: none"> • Deposition of upper Olympio Formation • Intrusion of Woodward Dolerite
c. 1845			<ul style="list-style-type: none"> • Deposition of Koongie Park Formation • Deposition of Winnama Formation • Intrusion of Angelo Microgranite • Intrusion of Sally Malay layered mafic–ultramafic intrusions 	
c. 1840				
1835–1805		----- HALLS CREEK OROGENY -----		
1835–1830 c. 1830	<ul style="list-style-type: none"> • Deposition of the Speewah Group in the Speewah Basin 	<ul style="list-style-type: none"> • Uplift and erosion 	<ul style="list-style-type: none"> • Intrusion of Mabel Downs Tonalite and McHale Granite • Intrusion of mafic–ultramafic intrusions including McIntosh intrusion, and Lamboo Ultramafics. Formation of D₃ folds and prominent foliation in Mabel Downs Tonalite synchronous with epidote–amphibolite facies metamorphism (M₃) 	

Table 1. (continued)

Age (Ma)	Kimberley Basin	Lamboo Complex		
		Western zone	Central zone	Eastern zone
c. 1827			<ul style="list-style-type: none"> Intrusion of Loadstone Monzogranite Deposition of Moola Bulla Formation in Moola Bulla Basin 	<ul style="list-style-type: none"> Layer-parallel shearing (D₃) related to south-westerly directed extension
1820–1810			<ul style="list-style-type: none"> Tight, upright, northeasterly plunging D₄ folds. Activation of shear zones under greenschist to lower amphibolite conditions (M₄) 	<ul style="list-style-type: none"> Upright to moderately inclined, open to isoclinal folding (D₄) accompanied by greenschist to lower amphibolite facies metamorphism (M₄)
c. 1821			<ul style="list-style-type: none"> Intrusion of Sally Downs Tonalite 	
1835–1805		<ul style="list-style-type: none"> Intrusion of gabbroic rocks, including Emull Gabbro Metamorphism of Marboo Formation and Neville Granodiorite to form Amherst Metamorphics 		
1825–1805		<ul style="list-style-type: none"> Intrusion of other Sally Downs supersuite 	<ul style="list-style-type: none"> granitoids and gabbros together with Spring Creek layered mafic–ultramafic intrusion 	
c. 1805–1785			<ul style="list-style-type: none"> Intrusion of granitoids of the San Sou suite 	
1800–1790	<ul style="list-style-type: none"> Deposition of Kimberley Group in the Kimberley Basin 		<ul style="list-style-type: none"> Deposition of the Revolver Creek, Texas Downs and Red Rock Formations 	
c. 1790	<ul style="list-style-type: none"> Intrusion of the Hart Dolerite sills into the Speewah and Kimberley Groups 			<ul style="list-style-type: none"> Intrusion of Sn-bearing pegmatites
c. 1740				
c. 1500	<ul style="list-style-type: none"> Deposition of Crowhurst and Bastion Groups 		<ul style="list-style-type: none"> Deposition of Mount Parker Formation and Bungle Bungle Dolomite in Osmond Basin Deposition of Wade Creek Formation, Ahern Formation, and Helicopter Siltstone in Victoria River Basin 	
c. 1200	<ul style="list-style-type: none"> Deposition of Glidden and Carr Boyd Groups Intrusion of Argyle lamproite pipe 			
c. 1000	----- YAMPI OROGENY -----			
	<ul style="list-style-type: none"> Large-scale sinistral strike-slip faulting and associated folding (D₅). Greenschist facies metamorphism (M₅) Intrusion of dolerite and porphyry dykes 			
c. 610 – c. 800	<ul style="list-style-type: none"> Deposition of the Louisa Downs Group in the Louisa Basin 	<ul style="list-style-type: none"> Deposition of Ruby Plains Group in Wolfe Creek Basin Deposition of Duerdin and Albert Edward Groups in Wolfe Creek Basin 		
c. 560	----- KING LEOPOLD OROGENY -----			
c. 560	<ul style="list-style-type: none"> Northwesterly directed folding and thrusting of Kimberley Group rocks. Reactivation of major faults (D₆) Deposition of the Lally Conglomerate, and eruption of the continental flood basalts of the Antrim Plateau Volcanics in the Ord and Bonaparte Basins 			
400–300	----- ALICE SPRINGS OROGENY -----			
400–300	<ul style="list-style-type: none"> Reactivation of major faults and associated north-northeasterly trending folding (D₇) 			
370–300	<ul style="list-style-type: none"> Deposition of siliciclastics and carbonates in the Ord and Bonaparte Basins 			
380–200	<ul style="list-style-type: none"> Deposition of carbonates and siliciclastics in the Canning Basin 			
70–50		<ul style="list-style-type: none"> Formation of plateau surface 		
20 to present		<ul style="list-style-type: none"> Uplift and dissection of plateau surface 		

SOURCE: Sheppard et al. (1997), Griffin et al. (1998), Tyler (2000)

calated with minor felsic volcanic rocks, pelitic and psammitic rocks, and carbonate rocks (Sheppard et al., 1999). The felsic volcanic rocks have been dated by sensitive high-resolution ion microprobe (SHRIMP) at $1880 \text{ Ma} \pm 3 \text{ Ma}$ (Blake et al., 1999b).

The Olympio Formation (*PHo*) disconformably overlies the Biscay Formation, and consists predominantly of a monotonous succession of turbiditic quartz wacke, greywacke, arkosic sandstone, and quartz sandstone, with interbedded siltstone and mudstone (Tyler et al., 1995). Near the base of the Olympio Formation there are two alkaline volcanic members: the Maude Headley Member (*PHov*) and the Butchers Gully Member (*PHob*). The Maude Headley Member consists of metamorphosed andesitic to trachytic volcanic and volcanoclastic rocks with a carbonate-rich matrix. It has been dated by SHRIMP at $1857 \pm 5 \text{ Ma}$ (Page and Hancock, 1988). The Butchers Gully Member consists of trachyandesite lavas and intrusions, together with associated volcanoclastic deposits (Tyler et al., 1995). Zircons from a pillow lava in the Butchers Gully Member have been dated by SHRIMP at $1848 \pm 3 \text{ Ma}$ (Blake et al., 1999b). Detrital zircons from the upper part of the Olympio Formation have been dated by SHRIMP at $1847 \pm 6 \text{ Ma}$, which gives a maximum age for deposition (Blake et al., 1999b). The Biscay Formation and lower part of the Olympio Formation were intruded by the Woodward Dolerite prior to deformation and regional metamorphism (Tyler et al., 1998a).

About 20 Ma separates the Biscay Formation from the lower part of the Olympio Formation, but there is no evidence that the Biscay Formation was deformed before deposition of the Olympio Formation (Tyler et al., 1995). Sheppard et al. (1999) suggests that the geological history of the Halls Creek Group, and chemistry of the basalt in the Biscay Formation, is consistent with a passive margin setting. This setting is also consistent with intraplate 'hot-spot' volcanism of the Butchers Gully and Maude Headley Members (Taylor et al., 1995).

Central zone

There are no exposures of basement rocks in the Central zone; the oldest rocks recognized are the Tickalara Metamorphics (Sheppard et al., 1999). The Koongie Park Formation post-dates the Tickalara Metamorphics in the southern part of the area. There is uncertainty about the stratigraphic relationship of a fault-bounded wedge of Milba Formation on HALLS CREEK to the other units (Blake et al., 1999b) so it is placed at the base of the map reference on Plate 1 for convenience.

The Tickalara Metamorphics (*PmT*) consists of mafic metavolcanic rocks including massive, amygdaloidal and pillowed basalt, volcanoclastic rocks, and volcanic breccia with interbedded mudstones, siltstones, sandstones, carbonate, and calc-silicate units (Tyler et al., 1995). Conformable amphibolites (*PmTo*) and possible metamorphosed mafic-ultramafic layered intrusions (*PmTon*) are also present (Sheppard et al., 1999). Detrital zircons in turbiditic sedimentary rocks give a maximum age of c. 1865 Ma for the sequence (Tyler and Page,

1996). The sequence was intruded by sheets of tonalite with subordinate granite and trondhjemite (*PmTg*) and metamorphosed at c. 1850 during the Hooper Orogeny (Tyler and Page, 1996). The metamorphic grade varies from high grade in the northern part of the zone, to medium grade to the southwest (Tyler et al., 1995). Sheppard et al. (1999) interpreted the basalts of the Tickalara Metamorphics to have formed in an oceanic island arc – backarc basin above a southeast-dipping subduction zone, or in an ensialic basin along the margin of the Kimberley Craton above a northwest-dipping subduction zone. The Hooper Orogeny is thought to reflect accretion of the island arc to the Kimberley Craton, or closure of the marginal basin (Sheppard et al., 1999).

The Koongie Park Formation (*Pe*) consists of metamorphosed mafic volcanic rocks, felsic volcanic and volcanoclastic rocks, turbiditic sandstone, siltstone, and mudstone, and interbedded chert, banded iron-formation (BIF), and carbonate rock (Orth, 1997; Griffin et al., 1998). Sewell (1999) informally divided the Koongie Park Formation in the Koongie Park area into three members: Coolibar tuff member; Camp shale member; and Weldon Creek lava member. The Coolibar tuff member consists of layered rhyolitic and dacitic lapilli and crystal tuffs. Overlying the Coolibar tuff member is the Camp shale member, which consists of siltstones and interbedded andesitic to dacitic volcanoclastic rocks and lavas, and includes the base-metal bearing 'Mimosa interval'. Spherulitic dacite and rhyolite lava flows of the Weldon Creek lava member overlie the Camp shale member. Orth (1997) proposed the term Onedin Member to include the Camp shale member and 'Mimosa interval' of Sewell (1999). SHRIMP U–Pb zircon dating of three rhyolite samples from the Koongie Park Formation on HALLS CREEK gave ages of 1845–1840 Ma (Page et al., 1994). The Koongie Park Formation is interpreted to have been deposited in a basin near the southeastern edge of the Kimberley Craton, associated with extension behind the convergent–strike-slip plate margin (Sheppard and Hassan, 2000).

The Milba Formation (*Pz*) is confined to fault-bounded wedges along the eastern edge of the Central zone and its relationship to other rock units is not known. It consists of greywacke, siltstone, sandstone, marble, impure calcareous rocks, chert, and minor mafic lavas and sills (Blake et al., 1999b).

Western zone

The Western zone is a continuation of the Hooper Complex in the King Leopold Orogen (Tyler et al., 1995; Sheppard and Hassan, 2000; Griffin et al., in press). The oldest rocks exposed are those of the Marboo Formation. The Marboo Formation (*Pm*) consists of metamorphosed turbiditic sandstone, siltstone, and mudstone (Griffin and Tyler, 1992; Tyler et al., 1999). The maximum age of deposition is c. 1870 Ma based on SHRIMP U–Pb dating of detrital zircon (Tyler et al., 1995, 1999); the minimum age is given by a high-level intrusive stock in the Marboo Formation dated at $1858 \pm 5 \text{ Ma}$ (Tyler et al., 1999).

The Whitewater Volcanics (*Pw*) unconformably overlies the Marboo Formation, and consists of dacitic to

rhyolitic ignimbrites, minor lava flows, lapilli tuff, and volcanogenic sedimentary rocks (Thorne et al., 1999). Two samples of Whitewater Volcanics have been dated at c. 1855 Ma using SHRIMP U–Pb geochronology (Griffin et al., in press). The volcanic rocks have been intruded by high-level porphyry intrusions (*PgPh*), and granitoid (*PgPc*) and gabbro (*PgPo*) of the Paperbark supersuite dated at 1865–1850 Ma. The volcanic rocks, porphyries, and granitoids are all cogenetic (Griffin et al., in press).

The Amherst Metamorphics consists of high metamorphic grade paragneisses (*PmAm*) and orthogneisses (*PmAgn*) that were intruded by c. 1820 Ma granitoids (Griffin et al., 1998). The Amherst Metamorphics is interpreted to have formed as a result of high-grade metamorphism of the Marboo Formation and c. 1860 Ma Neville Granodiorite (Paperbark supersuite) between 1860 and 1805 Ma (Griffin et al., 1998).

Younger granitoids and associated gabbros of the Lamboo Complex

The distribution of the younger granitoids in the Halls Creek Orogen is shown on Figure 3. The Sally Downs supersuite ranges in age from c. 1835 to 1805 Ma and consists of sixteen separate granitoid intrusions (*PgSg*) and coeval gabbroic intrusions (*PgSo*; Sheppard et al., 1995). The older intrusions are restricted to the Central zone, but intrusions in the supersuite that are younger than c. 1820 Ma (e.g. the 1808 Ma Mount Christine Granitoid) also intrude the Eastern and Western zones.

The San Sou suite (*PgSe*) is restricted to the southern part of the Lamboo Complex and includes the c. 1805 Ma Eastman Granite and the c. 1790 Ma San Sou Monzogranite. The Eastman Granite intrudes the southern part of the Central zone. The San Sou Monzogranite outcrops in the southern part of the Eastern zone on DOCKRELL. The San Sou Monzogranite is the youngest known granite in the Lamboo Complex, but is of similar age to The Granites Granite in the Granites–Tanami Inlier (Page and Sun, 1994).

Layered mafic–ultramafic intrusions in the Lamboo Complex

Layered mafic–ultramafic intrusions are confined to the Central and Western zones of the Lamboo Complex. Hoatson (1997, in press) divided the layered mafic–ultramafic intrusions in the Halls Creek Orogen into seven major groups on the basis of their field relationships, geochronology, style and intensity of deformation, and types of mineralization. The groups (numbered I to VII) were emplaced during at least three separate periods: 1855 Ma, 1845 Ma, and 1830 Ma (Hoatson, 1997).

Group I (e.g. Panton) intrusions (*Paa*) are coeval with Group II and Group III intrusions at c. 1855 Ma. The Group I intrusions are strongly differentiated and have the highest ratio of ultramafic to mafic cumulates. They have been deformed to produce steeply dipping folded bodies (Hoatson, 1997) and they are restricted to the Central zone. Group II (e.g. Springvale) intrusions (*Pas*) are

restricted to the Western zone. They are moderately dipping, sheet-like mafic bodies (Hoatson, 1997). The Group III Toby Gabbro in the Western zone is the sole representative of the Group III intrusions. It is a high level intrusion consisting of gabbroic and doleritic rocks of non-cumulus origin (Hoatson, 1997), and is considered to be part of the Paperbark supersuite (*PgPo*; Sheppard et al., 1997c; Tyler et al., 1997b).

The Group V (Sally Malay) intrusions (*Pay*) are differentiated mafic–ultramafic intrusions that have been dated at c. 1840 Ma (Hoatson, 1997), and are restricted to the Central zone. These intrusions are coeval with mafic and felsic volcanic rocks of the Koongie Park Formation (Hoatson and Blake, in press).

The Group IV (Wild Dog Creek) intrusions of Hoatson (1997) are sheet-like gabbroic bodies coeval with the Sally Downs supersuite (*PgSo*) in the Central zone. Group VI (McIntosh; *Pam*) intrusions also occur in the Central zone and include the large 1830 Ma funnel-shaped, gabbroic McIntosh intrusion and smaller sheet-like gabbroic bodies, such as the Armanda intrusion (Hoatson, 1997). Group VII (Black Hills Yard; *Pa*) intrusions intrude unlayered gabbroic rocks of the Paperbark supersuite in the Western zone (Hoatson, 1997).

The layered mafic–ultramafic Lamboo and Eastman Bore intrusions (*Pa*), which intrude the c. 1845 Ma Koongie Park Formation in the southern part of the Central zone, are not included in the classification of Hoatson (1997). Layered mafic–ultramafic intrusions within the Tickalara Metamorphics are also excluded.

King Leopold Orogen

The King Leopold Orogen includes the Palaeoproterozoic metasedimentary and igneous rocks of the Hooper Complex and the deformed margins of the Speewah and Kimberley Basins (Griffin and Grey, 1990a). The Hooper Complex outcrops on the western side of MOUNT RAMSAY. It includes the Marboo Formation, the Whitewater Volcanics, and granitoids of the Paperbark supersuite. The Hooper Complex is a continuation of the Western zone of the Lamboo Complex (Tyler et al., 1995, 1999; Griffin et al., in press).

Hooper Orogeny

The Hooper Orogeny was first recognized in the Hooper Complex of the King Leopold Orogen (Tyler and Griffin, 1990; Griffin et al., 1993). During the Hooper Orogeny, rocks of the Marboo Formation in the Western zone of the Lamboo Complex, and the Tickalara Metamorphics in the Central zone, were affected by two phases of deformation. The Marboo Formation was deformed and metamorphosed (D_1 – M_1) between c. 1870 and c. 1865 Ma before being unconformably overlain by the Whitewater Volcanics (Tyler et al., 1995, 1999). A second period of deformation and metamorphism (D_2 – M_2) occurred either prior to, or contemporaneously with, intrusion of the Paperbark supersuite between 1865 and 1850 Ma (Griffin

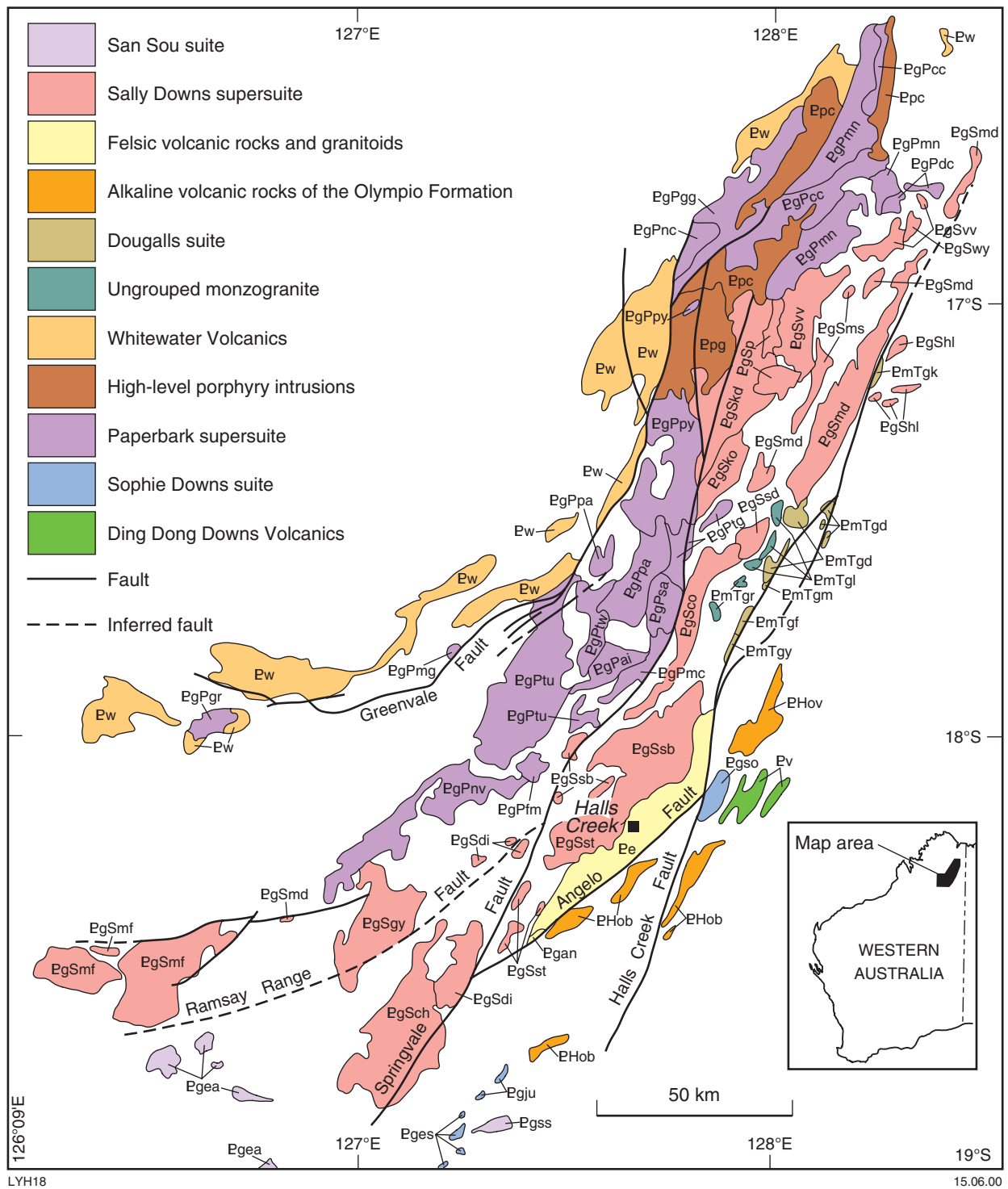


Figure 3. Granitoids and felsic volcanic rocks of the Halls Creek Orogen (after Sanders, 1999)

Key to granitoid and felsic volcanic rocks of the southern Halls Creek Orogen

San Sou suite (1805–1790 Ma)

Egea	Eastman Granite
Egss	San Sou Monzogranite

Eastern zone

Sally Downs supersuite (c. 1835–1805 Ma)

EgSco	Corrara Granite	EgSmf	Mount Fairbairn Monzogranite
EgSch	Mount Christine Granitoid	EgSms	Magotty Springs Monzogranite
EgSdi	Dillinger Monzogranite	EgSp	unnamed monzogranite and granodiorite
EgSgy	Grimpy Monzogranite	EgSsb	Shepherds Bore Granite
EgShl	McHale Granodiorite	EgSsd	Sally Downs Tonalite
EgSkd	Kevin Dams Monzogranite	EgSst	Loadstone Monzogranite
EgSko	Koondooloo Monzogranite	EgSvv	Violet Valley Tonalite
EgSmd	Mabel Downs Tonalite	EgSwy	Wesley Yard Monzogranite

Western, Central and Eastern zones

Felsic volcanic rocks and granitoids (c. 1845 Ma)

Pe	Koongie Park Formation
Egan	Angelo Microgranite

Central zone

Alkaline volcanic rocks of the Olympio Formation (1848–1856 Ma)

PHob	Butchers Gully Member
PHov	Maude Headley Member

Eastern zone

Tickalara Metamorphics granitoids

Dougalls suite (c. 1850 Ma)

EmTgd	Dougalls Tonalite	EmTgm	Monkey Yard Tonalite
EmTgf	Dead Finish Tonalite	EmTgy	Black Duck Tonalite
EmTgk	Corkwood Tonalite		

Central zone

Ungrouped monzogranite (c. 1865 Ma)

EmTgl	Fletcher Creek Monzogranite
EmTgr	Rose Bore Granite

Whitewater Volcanics (1865–1850 Ma)

Ew	Whitewater Volcanics
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High-level porphyry intrusions (1865–1850 Ma)

Epc	Castlereagh Hill Porphyry
Epg	Greenvale Porphyry

Paperbark supersuite (1865–1850 Ma)

EgPai	Airfield Granodiorite	EgPnc	Neil Creek Monzogranite
EgPcc	Crooked Creek Granite	EgPnv	Neville Granodiorite
EgPdc	Dinner Creek Tonalite	EgPpa	Paperbark Granite
EgPfm	Gnewing Granodiorite	EgPpy	Pandanus Yard Monzogranite
EgPgg	Gordons Gorge Granite	EgPsa	Sandy Dam Monzogranite
EgPgr	Little Gold River Microgranodiorite	EgPtg	Togo Monzogranite
EgPmc	Mussel Creek Granite	EgPtu	Tumagee Granite
EgPmg	Mad Gap Monzogranite	EgPtW	Top Water Tonalite
EgPmn	Mount Nyulasy Granite		

Western zone

Sophie Downs suite (c. 1910 Ma)

Eges	Esaw Monzogranite
Egju	Junda Microgranite
Egso	Sophie Downs Granophyre

Eastern zone

Ding Dong Downs Volcanics (c. 1910 Ma)

Ev	Ding Dong Downs Volcanics
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et al., 1993; Tyler et al., 1995, 1999). The Hooper Orogeny is interpreted to correspond to the accretion of the island arc basalts (Tickalara Metamorphics) of the Central zone to the Kimberley Craton (Myers et al., 1996; Sheppard et al., 1997c, 1999).

Halls Creek Orogeny

The earliest period of deformation in the Eastern zone involved layer-parallel shearing (D_3) related to extension during the Halls Creek Orogeny. This was followed by a second period of deformation (D_4) that resulted in the formation of open to isoclinal folds with east-northeasterly to northeasterly trending fold axes (Tyler et al., 1998a). The Halls Creek Orogeny also affected rocks of the Tickalara Metamorphics in the Central zone and the syn-orogenic granitoids and mafic rocks of the Sally Downs supersuite (Tyler and Page, 1996; Thorne et al., 1999). The Halls Creek Orogeny is interpreted to be the result of collision of the Eastern zone with the combined Central and Western zones at c.1800 Ma (Tyler and Page, 1996).

Palaeoproterozoic Granites–Tanami Complex and Archaean Lucas Craton

There has been no recent mapping by the GSWA in the area of the Granites–Tanami Complex on BILLILUNA and the geology on Plates 1 and 1a and the digital map is largely based on Blake et al. (1977). However, PNC Exploration (Australia) (Pearcey et al., 1988) and Blake (1996a) mapped late Archaean gneiss, granite, quartzite, conglomerate, and calc-silicate rock in the vicinity of Browns Range Dome. Uranium–lead SHRIMP ages on detrital zircons from the metasedimentary rocks of the Browns Range Dome area indicate a maximum depositional age of 2507 ± 22 Ma (Page et al., 1995). These rocks were intruded by leucogranite dated at 2510 Ma and have been interpreted as part of the Archaean Lucas Craton (AmB; Myers et al., 1996; Tyler, 2000).

Greywacke, sandstone, siltstone, and mudstone, with minor quartzite, banded chert, basalt, dolerite, gabbro, and felsic porphyry (EmG), was mapped as Killi Killi 'Beds' by Blake et al. (1977). However, recent mapping by the Northern Territory Geological Survey (Hendrickx et al., 2000) suggests that there are three main lithological units present in the Northern Territory: the MacFarlane Peak Group, overlain by the Dead Bullock Formation, which in turn is overlain by the Killi Killi Formation. The MacFarlane Peak Group consists of mafic volcanic and volcanoclastic rocks with minor clastic sedimentary rocks and calc-silicate rocks. The Dead Bullock Formation consists of dark-grey massive to laminated siltstone, carbonaceous shale with minor BIF, and calc-silicate rocks. The Killi Killi Formation consists predominantly of turbiditic sandstone and siltstone.

Glengarry Resources (1999, written comm.*) suggested that the sedimentary and mafic rocks of the informally named 'Bald Hill Sequence', which hosts the

Kookaburra and other deposits, are correlated with the Mount Charles Formation in the Northern Territory. However, Hendrickx et al. (2000) regard the Mount Charles Formation as equivalent to the Pargee Sandstone (see below) and restrict it to the Tanami Mine basalt sequence just north of the Tanami mines in the Northern Territory. The 'Bald Hill Sequence' is possibly correlated with the Dead Bullock Formation, which hosts the Dead Bullock soak and Granites gold deposits (Hendrickx et al., 2000).

Because of the uncertainty in correlating the newly mapped Northern Territory sequence with the rocks previously mapped as Killi Killi 'Beds' on the Western Australian side of the border, they are shown as undifferentiated (EmG) on Plate 1 and the digital map.

The Pargee Sandstone (EmGpg) unconformably overlies the rocks previously mapped as Killi Killi 'Beds'. It consists of sandstone and minor conglomerate (Blake et al., 1977).

The Slatey Creek Granite (PgGsl) intrudes the rocks previously mapped as Killi Killi 'Beds', 75 km east-southeast of Billiluna Homestead. Elsewhere, the granite is exposed only in cliffs where it is overlain by sandstone of the Lewis Range Sandstone (Redcliffe Pound Group; Blake et al., 1977).

Palaeoproterozoic basins

Sedimentary and volcanic rocks of the Speewah and Kimberley Basins unconformably overlie the Western zone of the Lamboo Complex and the Hooper Complex. The Kimberley Basin covers an area of about 160 000 km² (Plumb et al., 1981) and underlies most of the Kimberley Plateau. The Moola Bulla Formation is thought to be a time equivalent of the Speewah Group, and the Red Rock, Revolver Creek, and Texas Downs Basins are possible time equivalents of the Kimberley Basin (Thorne et al., 1999).

Speewah Basin

The Speewah Group (PS) is considered to have been deposited in a separate basin from the Kimberley Group, on the basis of a significant unconformity between the Speewah Group and the overlying Kimberley Group and different palaeogeographies (Griffin et al., 1993; Thorne et al., 1999). The Speewah Group includes the O'Donnell Formation, Tunganary Formation, Valentine Siltstone, Lansdowne Arkose, Luman Siltstone, and Bedford Sandstone. These formations form part of a transgressive–regressive cycle with fluvial sands passing into or alternating with shallow-marine facies and then back into fluvial sands (Plumb et al., 1981). Palaeocurrent measurements suggest that sediment was transported from the northeast and east (Gellatly et al., 1970). Zircons from a felsic volcanic unit in the Valentine Siltstone gave a SHRIMP U–Pb date of 1834 ± 3 Ma (AGSO OZCHRON

* Internet document on the world-wide web. Accessed 9 December 1999.

database). Thus, the group was deposited during the early stages of the Halls Creek Orogeny, and is older than most granites in the Sally Downs supersuite.

Moola Bulla Basin

Sedimentary rocks of the Moola Bulla Formation (*El*) in the Moola Bulla Basin are disconformable on the Koongie Park Formation (Griffin et al., 1998) in the Central zone of the Lamboo Complex. The Moola Bulla Formation may be a time equivalent of the Speewah Group (Dow and Gemuts, 1969; Griffin and Tyler, 1992; Blake et al., 1999b). The Moola Bulla Formation includes sandstones, conglomerates, mudstones, and shales, which are interpreted to have formed in a fluvial and lacustrine to marine environment with a provenance to the southwest (Blake et al., 1999b).

Kimberley Basin

The Kimberley Group unconformably and disconformably overlies the Speewah Group (Griffin et al., 1993; Thorne et al., 1999). The Kimberley Group is interpreted to have been deposited within a broad, semi-enclosed, shallow-marine basin (Plumb et al., 1981). Palaeocurrent indicators in the Kimberley Group indicate sediment transport from the north and north-northwest (Gellatly et al., 1970).

The King Leopold Sandstone (*EKl*) at the base of the Kimberley Group consists of medium- to coarse-grained sandstone and has medium- to large-scale trough cross-stratification (Thorne et al., 1999).

The Carson Volcanics (*EKc*) conformably overlies the King Leopold Sandstone, and consists of interlayered massive to amygdaloidal basalt, quartz sandstone, feldspathic sandstone, siltstone, and chert (Thorne et al., 1999).

The Warton Sandstone (*EKw*) conformably overlies the Carson Volcanics and consists of coarse-grained quartz sandstone with minor feldspathic sandstone. The sandstone has medium- to large-scale trough cross-stratification and horizontal planar stratification (Thorne et al., 1999).

The Elgee Siltstone (*EKe*) conformably overlies the Warton Sandstone and consists predominantly of red-brown siltstone with thin interbedded quartz sandstone (Plumb, 1968). Bruinsma (1970) subdivided the Teronis Member, at the base of the Elgee Siltstone, into four units: Teronis I consisting of grey-green sandstone with minor siltstone and shale; Teronis II consisting of grey-green shale and siltstone with minor fine-grained sandstone and oolitic and algal dolomite; Teronis III consisting of red-brown siltstone and minor shale; and Teronis IV consisting of green shale, siltstone, sandstone, calcareous sandstone and siltstone, and algal dolomite. Subeconomic syngenetic copper mineralization is widespread in the Teronis IV unit. Owen (1970) suggested that the Teronis Member was deposited on a tidal flat, or open, shallow lagoon.

The Pentecost Sandstone (*EKp*) has been subdivided into three units. The lower unit consists of thinly bedded to laminated quartz sandstone; the middle unit consists of

planar-stratified or cross-stratified quartz sandstone and siltstone, with glauconitic sandstone and shale at the base; and the upper unit consists of massive, trough cross-bedded quartz sandstone and pebbly sandstone (Thorne et al., 1999).

Red Rock Basin

The Red Rock Formation (*Ek*), which is a time equivalent of the Kimberley Group, is faulted against the Olympio Formation in the Eastern zone of the Lamboo Complex, is unconformable on the Winnama Formation in the Central zone in the Osmand Range, and overlies the Moola Bulla Formation (east of Halls Creek; Tyler, 2000). The Red Rock Formation consists of sandstone, conglomerate, siltstone, mudstone, massive to amygdaloidal basalt flows, and basaltic breccia (Tyler et al., 1997a).

Texas Downs Basin

Sedimentary rocks of the Texas Downs Formation (*Ex*), in the Texas Downs Basin, overlie the Red Rock Formation and McHale Granodiorite with angular unconformity (Thorne et al., 1999). The formation consists of sandstone, siltstone, minor conglomerate, and basalt (Thorne et al., 1999). The age of the Texas Downs Formation is poorly constrained but it is believed to be broadly coeval with the lower part of the Kimberley Group (Thorne et al., 1999).

Revolver Creek Basin

In the Revolver Creek Basin, the Revolver Creek Formation (*Er*) unconformably overlies the Marboo Formation, Whitewater Volcanics, and Castlereagh Hill Porphyry, and is overlain by the Carr Boyd Group (Thorne et al., 1999). The formation consists of sandstone, conglomerate, siltstone, dolerite, and basalt, and is thought to be coeval with the lower part of the Kimberley Group (Thorne et al., 1999).

Hart Dolerite

The Hart Dolerite (*Edh*) consists of a series of massive dolerite sills and less extensive granophyre that intrude the Speewah Group and lower part of the Kimberley Group. The Hart Dolerite has been dated at c. 1790 Ma using SHRIMP U–Pb geochronology (Thorne et al., 1999). Alvin (1993) proposed the term Yilingbun granophyre for the granophyric phase, which intrudes the dolerite in the Speewah Valley.

Mesoproterozoic basins

Bastion Basin

The Bastion Group (*EB*) unconformably overlies the Kimberley Group in the Bastion Basin. The Bastion Group includes three formations: the Mendena Formation, consisting of sandstone with minor siltstone and

carbonate; the Wyndham Shale, consisting of siltstone with minor sandstone; and the Cockburn Sandstone, consisting of quartz sandstone and micaceous sandstone (Thorne et al., 1999). The age of the Bastion Group is uncertain but it is younger than the c. 1790 Ma Hart Dolerite and older than the c. 540 Ma Antrim Plateau Volcanics, which overlie the group.

Crowhurst Basin

The Crowhurst Group (*CR*) unconformably overlies the Pentecost Sandstone of the Kimberley Group in the Crowhurst Basin, and may be equivalent to the Birrindudu Basin (Tyler et al., 1998e). The group consists of four formations: the Hilfordy Formation, consisting of sandstone with interbedded siltstone and shale; the Liga shale, consisting of shale and micaceous siltstone; the Collett Siltstone, consisting of laminated siltstone with minor interbedded dolomite; and the Hibberston Dolomite, consisting of dolomite, which is commonly stromatolitic (Griffin and Grey, 1990b; Tyler et al., 1998c).

Osmond Basin

The Mount Parker Formation (*Pt*) in the Osmond Basin, consists of sandstone and conglomerate that was deposited conformably on the Texas Downs Formation (Tyler et al., 1997a). The Mount Parker Formation is conformably overlain by the Bungle Bungle Dolomite (*Bu*). The Bungle Bungle Dolomite includes dolomite, which is commonly stromatolitic, dolomitic sandstone, siltstone, mudstone, stromatolitic chert, and chert breccia (Tyler et al., 1997a). Felsic, lithic, vitric tuffs have been reported near the base of the Bungle Bungle Dolomite (de Angelis, 1984). The Mount Parker Formation and Bungle Bungle Dolomite may be equivalents of the Birrindudu Group.

Carr Boyd Basin

The Carr Boyd Group (*BC*) of the Carr Boyd Basin unconformably overlies Palaeoproterozoic rocks of the Lamboo Complex and the Revolver Creek Basin (Thorne and Tyler, 1996). Lower units of the Carr Boyd Group have been

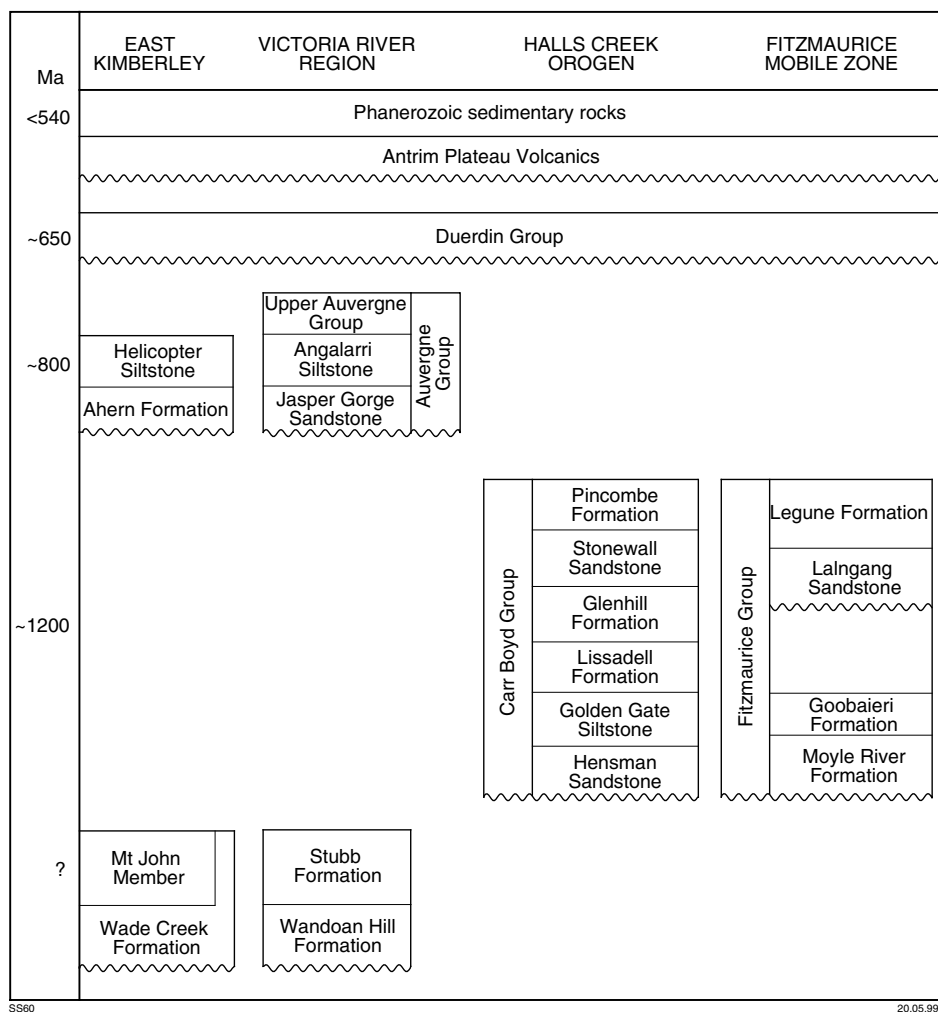


Figure 4. Regional correlation of Mesoproterozoic to Phanerozoic rocks in the east Kimberley (after Thorne et al., 1999)

correlated with the Auvergne Group in the Victoria River Basin (Sweet, 1977), but more recent work suggests that the Carr Boyd Group is older and probably correlates with the Fitzmaurice Group of the Fitzmaurice Mobile Zone in the Northern Territory (Thorne and Tyler, 1996; Thorne et al., 1999; Fig. 4). The Carr Boyd Group is interpreted to have formed in a sandy, braided delta complex and an adjacent siliciclastic marine shelf (Thorne and Tyler, 1996).

The Carr Boyd Group includes the Hensman Sandstone, Golden Gate Siltstone, Lissadell Formation, Glenhill Formation, Stonewall Sandstone, Pincombe Formation, and Bandicoot Range beds, but the upper two units do not outcrop in the study area. The Hensman Sandstone consists predominantly of quartz arenite with minor siltstone and conglomerate. The base of the Golden Gate Siltstone consists of a thin green siltstone and sandstone unit; this is overlain by a 7–10 m-thick ferruginous sandstone and sandy ironstone unit, which hosts the Pompeys Pillar and Matsu iron deposits. The remainder of the Golden Gate Siltstone consists of interbedded argillite and quartz sandstone. The Lissadell Formation consists of quartz arenite and lithic quartz sandstone interbedded with siltstone. The Glenhill Formation consists of interbedded siltstone and sandstone. The Stonewall Sandstone consists of quartz sandstone and pebbly sandstone with minor interbedded siltstone.

Glidden Basin

The Glidden Group (*BG*) of the Glidden Basin overlies the Amherst Metamorphics and Carson Volcanics of the Kimberley Group on MOUNT RAMSAY (Tyler et al., 1998c). The group includes the Harms Sandstone (quartz sandstone with interbedded shale), Matheson Formation (shale, siltstone, and minor sandstone), Forman Sandstone (quartz sandstone), and Maddox Formation (black shale, siltstone, and minor shale; Grey and Griffin, 1990).

Mesoproterozoic Yampi Orogeny

The Yampi Orogeny produced large-scale shearing in the crystalline rocks of the Hooper Complex, and north-easterly trending folding and thrusting in the Kimberley Basin (Tyler and Griffin, 1990; Griffin et al., 1993). Ductile strike-slip faulting (D_5) along north-northeasterly trending structures that initially formed during the Palaeoproterozoic in the Halls Creek Orogen is also interpreted to be related to the Yampi Orogeny (Tyler et al., 1995). Potassium–argon ages from sheared granitoid rocks from the Hooper Complex place age limits of between 1475 ± 12 and 999 ± 9 Ma on the Yampi Orogeny (Shaw et al., 1992).

Mesoproterozoic to Neoproterozoic basins

Victoria River Basin

The Mesoproterozoic Wade Creek Formation has been correlated with lower units of the Victoria River Basin in

the Northern Territory, whereas the Neoproterozoic Ahern Formation and Helicopter Siltstone correlate with the Auvergne Group of the Victoria River Basin in the Northern Territory (Dow and Gemuts, 1969; Thorne et al., 1999; Fig. 4).

The Wade Creek Formation (*Es*) lies unconformably on the Bungle Bungle Dolomite (Grey and Griffin, 1990; Tyler et al., 1997a). It consists of sandstone, conglomerate, siltstone, mudstone, and chert breccia (Tyler et al., 1997a). The Neoproterozoic Ahern Formation and Helicopter Siltstone (*Ef*) unconformably overlie the Wade Creek Formation. The Ahern Formation consists of sandstone with minor interbedded pebbly sandstone, conglomerate, and siltstone (Thorne et al., 1999). The Helicopter Siltstone, which conformably overlies the Ahern Formation, consists of laminated micaceous siltstone interlayered with quartz sandstone (Thorne et al., 1999).

Birrindudu Basin

The Birrindudu Basin contains the Birrindudu and Redcliffe Pound Groups, and rocks correlated with the Redcliffe Pound Group.

The Birrindudu Group (*BY*) rests unconformably on rocks of the Granites–Tanami Complex. It consists of the Gardiner Sandstone, Talbot Well Formation, and Coomarie Sandstone. The Gardiner Sandstone consists predominantly of sandstone with minor siltstone and shale, and glauconitic sandstone beds. Conglomerate is present at the base of the Gardiner Sandstone. The Talbot Well Formation consists of stromatolitic chert, sandstone, siltstone, and limestone. The Coomarie Sandstone is composed of sandstone (Blake et al., 1977).

The Redcliffe Pound Group (*PD*) includes the Lewis Range Sandstone, the Murraba Formation, and the Erica Sandstone. The Denison beds, the Jawilga beds, and the Boee beds are probable correlatives of the Redcliffe Group (Blake et al., 1977; Grey, 1990).

Neoproterozoic basins

Wolfe Creek Basin

Neoproterozoic glaciogene rocks of the Wolfe Creek Basin unconformably overlie rocks of the Victoria River and Carr Boyd Basins, and rocks of the Lamboo Complex. The Wolfe Creek Basin includes the Ruby Plains Group, Duerdin Group, and Albert Edward Group (Blake et al., 1997; Thorne et al., 1999).

The Ruby Plains Group (*BP*) is interpreted to represent shallow marine-shelf deposits (Blake et al., 1997). The group includes the Mount Kinahan Sandstone, the Eliot Range Dolomite, and the Iljarra Sandstone.

The Duerdin Group (*DE*) is a glacial sequence that includes the Fargoo Tillite, Frank River Sandstone, Moonlight Valley Tillite, Jarrod Sandstone Member, Johnny Cake Shale Member, and Ranford Formation (Dow and Gemuts, 1969). It is correlated with the lower

part of the Marinoan glacial sequence in central and southern Australia (Corkeron et al., 1996).

The Albert Edward Group (*EL*) includes the Mount Forster Sandstone, Elvire Formation, Boonal Dolomite, Timperley Shale, Nyuleless Sandstone, and Flat Rock Formation (Dow and Gemuts, 1969). The group is interpreted to have been deposited in shallow water and is equivalent to the upper part of the Marinoan glacial sequence in central and southern Australia (Corkeron et al., 1996).

Louisa Basin

Sedimentary rocks of the Columbo Sandstone, Kuniandri Group, and Louisa Downs Group were deposited in the Louisa Downs Basin. They are correlated with units in the Wolfe Creek Basin (Corkeron et al., 1996). The Columbo Sandstone (*Pc*) includes sandstone and chert-pebble breccia that overlie the Kimberley and Crowhurst Group (Tyler et al., 1998c). The Kuniandri Group (*LN*) includes the Landrigan Tillite, Stein Formation, Wirara Formation, and Mount Bertram Sandstone (Tyler et al., 1998c). Five formations are included within the Louisa Downs Group (*LI*): the Egan Formation, Yurabi Formation, McAlly Shale, Tean Formation, and Lubbock Formation (Tyler et al., 1998c).

Neoproterozoic King Leopold Orogeny

The King Leopold Orogeny (*D₆*) produced extensive west-northwesterly trending folding and thrusting in the King Leopold Ranges along the southwestern margin of the Kimberley and Speewah Basins (Griffin and Myers, 1988; Tyler and Griffin, 1990), together with the reactivation of shear zones in the Hooper Complex (Tyler et al., 1991; Shaw et al., 1992). Reactivated shear zones have given K–Ar ages of c. 560 Ma (Shaw et al., 1992). In the Lamboo Complex, the King Leopold Orogeny produced sinistral strike-slip faulting, folding, and uplift of Neoproterozoic glaciogene rocks (Griffin et al., 1998; Thorne et al., 1999).

Phanerozoic basins

Ord and Bonaparte Basins

The Ord Basin lies to the east of the Halls Creek Fault and the Bonaparte Basin lies to the west and north of the Halls Creek Fault. The Cambrian Antrim Plateau Volcanics (*Ca*) is the only formal unit common to both basins. In the northern part of the area, the Antrim Plateau Volcanics is about 200 m thick and unconformably overlies Palaeoproterozoic to Mesoproterozoic rocks, whereas further south, it is about 1000 m thick, and unconformable on the Neoproterozoic Duerdin and Albert Edward Groups of the Wolfe Creek Basin (Mory and Beere, 1988; Thorne et al., 1999). The Antrim Plateau Volcanics consist of massive to amygdaloidal basalt and basaltic breccia with minor interbedded sedimentary rock. The Mount Close Chert

Member consists of laminated, stromatolitic chert and chert breccia (Mory and Beere, 1988; Thorne et al., 1999).

The Cambrian Goose Hole Group (*CG*) overlies the Antrim Plateau Volcanics in the Ord Basin. It includes the Headleys Limestone, Nelson Shale, Linnekar Limestone, Panton Formation, Eaglehawk Sandstone, and Overland Sandstone (Mory and Beere, 1988; Tyler et al., 1997b; Thorne et al., 1999).

The Devonian Mahony Group (*DM*) conformably overlies the Cambrian Goose Hole Group in the Ord Basin, and includes sandstone, pebbly sandstone, and conglomerate (Mory and Beere, 1988).

The Cockatoo Group (*DC*) is interpreted as a mixed alluvial-fan–fluvial–eolian–shallow-marine succession deposited in response to active strike-slip faulting along the Glenhill – Revolver Creek – Carr Boyd Fault system (*D₆*; Alice Springs Orogeny) in the Bonaparte Basin (Thorne and Tyler, 1996; Thorne et al., 1999). In the study area, the Cockatoo Group includes the Ragged Range Conglomerate, the Cyril Sandstone, the Kellys Knob Sandstone, and the Galloping Creek Formation (Thorne et al., 1999).

Canning Basin

Lennard Shelf

Devonian limestones (*Dr*) form part of the Devonian ‘Great Barrier Reef’ on the Lennard Shelf (Griffin et al., 1993). The reef complexes include oolitic, cyanobacterial, stromatoporoid, and stromatolitic limestones, and debris flow deposits. This grouping includes the Pillara Limestone, Sadler Limestone, Gogo Formation, Virgin Hills Formation, Nullara Limestone, Winjana Limestone, Bugle Gap Limestone, and Piker Hills Formation on MOUNT RAMSAY (Tyler et al., 1998c).

Devonian boulder, cobble, and pebble conglomerates (*Dg*) are interpreted to have formed in platform and basal settings. They include the Sparke Conglomerate, Elma Conglomerate, and Barramundi Conglomerate. Locally, the conglomerates interfinger with the reef complex limestones (Tyler et al., 1998c).

The Late Devonian to Early Carboniferous Fairfield Group (*DCF*) consists of limestone, sandstone, siltstone, and shale deposited in a shallow-marine environment (Tyler et al., 1998c). The Fairfield Group is overlain by the Late Carboniferous to Early Permian Grant Group (*P*).

Remainder of Canning Basin

Shallow-marine to fluvial sandstones and conglomerates of the Permian Noonkanbah Formation and Liveringa Group (*P*) were deposited in the Fitzroy Trough on MOUNT RAMSAY (Tyler et al., 1998c). These formations are also present on MOUNT BANNERMAN and BILLILUNA where they are overlain by Triassic sedimentary rocks. Ordovician marine sedimentary rocks (*Oc*) and Devonian fluvial sedimentary rocks (*Dk*) lie beneath Carboniferous and Permian sedimentary rocks on the Billiluna Shelf (Blake et al., 1977).

Palaeozoic Alice Springs Orogeny

The 400–300 Ma Alice Springs Orogeny (D₇) is the youngest phase of major faulting and folding in the east Kimberley. During this period there was reactivation of the Halls Creek Fault and other northeasterly trending faults in the Halls Creek Orogen, and also a period of north-northeasterly sinistral strike-slip faulting in the Pillara Range in the Canning Basin (Thorne and Tyler, 1996).

Regolith

A veneer of unconsolidated or partly consolidated sediments including sand, silt, soil, calcrete, and silcrete covers much of MOUNT BANNERMAN and BILLILUNA, and is locally developed in the vicinity of rivers and creeks in the remainder of the area (Plumb and Gemuts, 1976; Warren, 1994a). There are also areas of ferricrete and ferruginous pisolitic gravel that represent remnants of a pre-Miocene land surface (Blake, 1996b).

Exploration and mining

1882–1964

Gold was the first mineral to be discovered in the east Kimberley region in September 1882 by two prospectors, Philip Saunders and Adam Johns. Johns was alerted to the possible gold potential of the Kimberley by Alexander Forrest when they met at Pine Creek at the end of Forrest's 1879 Kimberley expedition (Playford and Ruddock, 1985). Due to lack of water and Johns' illness, Saunders and Phillips were unable to follow the gold to its source. Saunders' account of their discovery was published in the *South Australian Advertiser* in 1883 (reprinted in Clement and Bridge, 1991). During the Kimberley Survey Expedition in 1884, Edward Hardman (Government geologist attached to the expedition) found auriferous quartz reefs and alluvial gold over a large area traversed by the Margaret, Mary, Elvire, Panton, and Ord rivers (Hardman, 1885). Hardman (1885) also noted copper mineralization on the Margaret River (Me No Savvy), lead and zinc mineralization on the Ord River, and alluvial tin and semiprecious stones. Although Hardman (1885) did not have sufficient time in the area to show that payable gold deposits existed, his report and coloured geological map showing auriferous areas enabled the prospecting party of Charles Hall and John Slattery to obtain the first payable amounts of gold near Old Halls Creek in 1885 (Playford and Ruddock, 1985).

The ensuing Kimberley gold rush reached its peak in mid 1886 when thousands of diggers from the southern parts of Western Australia, the Eastern colonies, and New Zealand came to work the alluvial deposits and rich narrow quartz reefs at the Ruby Creek, Mount Bradley – Brockman, Mount Dockrell, Halls Creek, and Mary River centres. This led to the proclamation in May 1886, of the Kimberley as the first goldfield in Western Australia. However, by the time Woodward visited the area as

Government geologist in 1890, most of the alluvial and underground workings had been abandoned (Woodward, 1891) due to the remoteness of the area, lack of water, illness and disease, shortage of fuel, small size of the deposits, and the discovery of richer fields in the Pilbara, Murchison, and Eastern Goldfields. The Ruby Queen mine, the largest of the historical mines, continued production until 1908, then produced sporadically between 1938 and 1964. There was also minor production from the other mining centres mentioned above between 1935 and 1960.

Recorded historical (pre-1964) production is relatively low, totalling 1056.8 kg, and consisting of 511.5 kg of alluvial and dollied gold and 545.3 kg of lode gold. However, production was probably significantly higher as there were no reporting requirements for gold produced prior to 1896 and gold produced between 1896 and 1898 during the peak of the gold rush probably went largely unreported in order to avoid the gold tax of 2s 6d per ounce.

The Tanami gold mine in the Northern Territory was discovered in 1900. Talbot found quartz with up to 5 g/t of gold to the east of Larranganni Bluff in the Granites–Tanami Complex in Western Australia in 1909. This followed his visit to the Tanami mine when he made a traverse from the Tanami mine to Halls Creek to determine if the formation hosting the Tanami deposit continued into Western Australia. The Kookaburra deposit was found by Glengarry Resources NL in 1995 not far from one of Talbot's sampling sites (Doust, 1997).

The State's largest fluorite deposit was discovered on Speewah Station in 1905 by botanist W. V. Fitzgerald (Simpson, 1951). Blatchford (1928) visited the occurrence and discovered an additional three fluorite veins; he also reported on the nearby Martins lead–silver prospect. A lead–silver deposit was also discovered near Mount Amherst during an AGGSNA survey (Jones, 1938). However, there was no historical production of base metals or fluorite because the distance to port made them uneconomical to mine.

Exploration by mining companies commenced in the 1950s. During 1954, the BMR carried out an airborne radiometric survey over part of the area and published a list of 54 well-defined anomalies (Goodeve, 1955). Torbernite had previously been discovered in the Dunham River area by United Uranium NL and Western Uranium, but the survey resulted in a sudden increase in uranium exploration in the area (Harms, 1959). Amongst these companies was Rio Tinto Finance and Exploration who searched for uranium deposits in black shale, chlorite schist, carbonate, and calc-silicate rocks adjacent to granites (i.e. similar to Rum Jungle, Northern Territory), for uranium in major faults cutting Proterozoic sedimentary and volcanic rocks, and for uranium in shear zones in granite (similar to Edith River, Northern Territory). Most radiometric anomalies coincided with lateritic ridges and were considered to be caused, in part, by topographic effects combined with moderate background radiation.

The iron ore deposits at Pompeys Pillar were discovered and explored by the Broken Hill Company during the late 1950s (Harms, 1959). Further exploration and resource evaluation were carried out by Bell Brothers, Planet Management and Research, Sentinel Mining Company, and Western Mining Corporation between 1961 and 1973.

Between 1960 and 1961, New Consolidated Goldfields Australia carried out exploration for sandstone–conglomerate-hosted gold and uranium mineralization in the basal conglomerates of the Upper Proterozoic rocks in the Gardiner Range area, and discovered uranium–xenotime mineralization in the Killi Killi area. However, a considerable amount of follow-up exploration in the area failed to find economic mineralization. New Consolidated Goldfields also found traces of gold in Slaty Creek, and in a creek draining the northern slopes of Junction Hill in the Granites–Tanami Complex.

Between 1961 and 1963, Peko Mines carried out exploration of previously identified copper occurrences at Mount Angelo, New Homestead (near Lamboo Homestead), and the Venture 3 or Golf Course prospect near Halls Creek.

Base metal gossans at Little Mount Isa and Ilmars were discovered in 1962 during regional mapping of GORDON DOWNS (Gemuts, 1963). Platinum-bearing chromite in the Panton Sill was also discovered between 1962 and 1963 during the regional mapping program on DIXON RANGE; assays of up to 3 g/t Pt were recorded (Dow and Gemuts, 1967).

Between 1963 and 1966, Pickands Mather and Company International carried out regional stream-sediment sampling over most of the east Kimberley area. Follow-up of the anomalies outlined by this program led to the discovery of many base metal (e.g. Gentle Annie), nickel–copper (e.g. Bow River), and chromite (e.g. Lamboo) prospects. This was the beginning of modern intensive mineral exploration in the Kimberley region.

1965–1975

Between the mid 1960s and the mid 1970s, the main targets of exploration were volcanogenic massive sulfide (VMS) and volcanic exhalative base-metal deposits; copper–nickel sulfides and platinum group elements (PGE) associated with layered mafic intrusions; and sandstone-hosted and unconformity-type uranium deposits. There was also some exploration for hydrothermal lode gold and porphyry-type copper and, from 1972 onwards, for Mississippi Valley-type mineralization and diamonds.

Pickands Mather and Company International discovered VMS-style mineralization in the Koongie Park Formation at Moola Bulla and at Koongie Park, southwest of Halls Creek between 1966 and 1968. Follow-up work by Kennecott Explorations (Australia) between 1972 and 1976 led to significant intersections of copper, lead, and zinc in the Koongie Park Formation at the Gordon Downs One (Golf Course), Gordon Downs Two (Mount Angelo

North), Gordon Downs Three (Koongie Park), Eastman, and Mount Ramsay Three (Dusty Bore) prospects.

Platinum group element mineralization was discovered in an ultramafic intrusion, informally referred to as the Eastman intrusion, by Kennecott Explorations (Australia) Limited during 1974.

Broken Hill Proprietary Company (BHP), Australian Anglo American, Planet Management and Research, and Planet Mining Company found widespread copper mineralization in the Carson Volcanics between 1968 and 1972. Concentrations of native copper were found in the Antrim Volcanics in 1968 by Metals Exploration NL at Cuddys prospect, but no economic copper mineralization was found.

In 1972, Endeavour Oil Company NL found gossans in Devonian limestone of the Lennard Shelf at Ross Hill that assayed up to 27% Zn and 10% Pb. Mississippi Valley-type base metal gossans were also found in carbonates of the Winnama Formation in the Osmand Valley by Australian Anglo American in 1972. Trend Exploration intersected significant lead mineralization in the Bungle Bungle Dolomite in 1973.

Australian Anglo American discovered the Sally Malay nickel–copper–cobalt deposit during a stream-sediment and gossan survey during 1973. Nickel–copper mineralization in mafic–ultramafic intrusions was also found at the Keller Creek, McKenzies Spring, Coolumbooloo, Springvale, Billymac, Corkwood, and Melon Patch prospects.

Although uranium exploration did not locate any significant uranium deposits during this period, Trend Exploration discovered highly anomalous niobium, yttrium, tin, and fluorite mineralization, associated with a radiometric anomaly, at a prospect that is now known as the Brockman deposit.

1975–present

Exploration boomed in the east Kimberley from the mid 1970s following the discovery of diamonds in the north Kimberley in 1973 and the diamondiferous lamproite pipes at Ellendale in the west Kimberley in 1977. The results of this exploration for various commodities are discussed below.

Diamond

From the mid-1970s, diamond became a major exploration target in the east Kimberley area. The Ashton Joint Venture recovered two diamonds from Smoke Creek during a reconnaissance-sampling program in August 1979. Follow-up sampling in October 1979 led to the discovery of the Argyle AK1 pipe, the largest diamond deposit in the world (Boxer and Jaques, 1990). Mining commenced in 1985. Alluvial diamonds have also been mined from drainages and palaeodrainages draining the Argyle AK1 pipe (Upper Smoke Creek, Limestone Creek, Bow River). Exploration for other lamproites and kimberlites has continued throughout the east Kimberley.

Gold

Most recent gold exploration in the Lamboo Complex has been in the vicinity of known lode gold and alluvial workings. There was minor production of alluvial gold from the Majeed area (Red Gully, Sam Hazlett Creek, Chinaman Creek, Majeed Creek, Blue Bar Creek, and White Elvire River) by 125 Nominees between 1981 and 1983 and from the Mary River and its tributaries by Roebuck Resources NL and Alcott Holdings between 1988 and 1989. The development of carbon-in-pulp technology led to exploration for stockwork deposits in the vicinity of the old high-grade lode mines that could be mined by low-cost open-pit techniques. The Palm Springs mine (incorporating the old Butchers Gully workings) and the Nicolsons Find mine (50 km to the west) are examples of old mines with resources in quartz stockwork. Precious Metals Australia commenced open-pit production at the Palm Springs mine in April 1995, and 2272 kg of gold was produced before the mine was placed on care and maintenance in April 1997 as a result of falling gold prices (Precious Metals Australia Limited, 1997).

Gold exploration has also been carried out in the Granites–Tanami Complex. Anomalous gold was first reported in the Larranganni Bluff area by Talbot (1910), but it was not until 1983 to 1988 that CRA Exploration, Dry Creek Mining NL, and Mr A. R. Burns carried out exploration for Granites-style gold mineralization in this area. A costean in an area with anomalous arsenic exposed two 40 cm-thick quartz veins assaying 1.29 and 2.18 g/t Au respectively, but the tenement was relinquished; this is now the Cuckoo prospect of Glengarry Resources NL. In 1995, Glengarry Resources NL completed a program of regional geochemical sampling using a power auger. Follow-up of a gold anomaly led to the announcement of the discovery of the Kookaburra deposit on 24 November 1995 (Doust, 1997) and other prospects including Sandpiper, Hawk, Finch, Osprey, Albatross, and Cuckoo North.

Acacia Resources (now AngloGold) is also actively exploring for gold in the Tanami region of Western Australia. Acacia Resources recently announced the discovery of extensive high-grade gold mineralization in several zones over an area of 800 × 200 m with intersections up to 8 m at 458 g/t at the Coyote prospect (Acacia Resources Limited, 1999).

Base metals

Exploration for volcanogenic massive sulfide (VMS) and volcanogenic exhalative base metal deposits, mainly in the vicinity of the earlier discoveries, has continued since 1975. The most significant deposits defined are Mount Angelo North and Sandiego in the Koongie Park area (Sewell, 1999).

Intense exploration for Mississippi Valley-type mineralization in the Lennard Shelf continued during this period, particularly by BHP Minerals and Billiton Australia. Base-metal bearing gossans were discovered in the Horse Spring Range area (Enigma gossan) and at

Findlay Hill. The Kapok deposit, a blind deposit 350 m below surface and on the western margin of the study area, was discovered by BHP Minerals in 1989 as the result of aeromagnetic interpretation (Ferguson, 1999). Mining of the Kapok deposit by Western Metals commenced in 1997.

Nickel–copper and PGE

Exploration for nickel–copper and PGE mineralization in mafic–ultramafic intrusions has continued since 1975, mainly in the vicinity of previously discovered occurrences.

Uranium

Between 1980 and 1984, Mineral Reserves and Energy Reserves Canada carried out extensive radiometric and geobotanical exploration for unconformity-type uranium at the contact between the Tanami Complex and overlying Gardiner Sandstone in the Birrindudu area (Gardner Range, Laranganni Bluff, and Killi Killi Hills). This resulted in the discovery of the Don uranium–gold occurrence. BHP Minerals and PNC Exploration Australia carried out further exploration in this region, but no significant uranium deposits were found.

Rare earth elements and niobium–tantalum

The Cummins Range carbonatite was discovered by CRA Exploration during the follow-up of an aeromagnetic anomaly. The airborne survey was flown in 1978 over the inferred junction of the Halls Creek Orogen and the King Leopold Orogen, south of Cummins Range, primarily in the search for kimberlites. Drilling carried out between 1982 and 1984 gave an indicated resource of 3 to 4 Mt at a grade of 2 to 4% total rare earth oxides (REO), but this was considered too low to be mined economically (Andrew, 1990).

West Coast Holdings followed up an area of anomalous niobium, yttrium, tin, and fluorite discovered by Trend Exploration in tuffs near Halls Creek and by 1988 had defined a resource at the Brockman deposit.

Tungsten

Union Oil Development Corporation explored for tungsten in the eastern zone of the Lamboo Complex using ultraviolet (UV) lamping. About 90 scheelite occurrences were found over a strike length of 20 km in the Castle Creek (Berthas Butt) area. The best intersection from follow-up drilling was 2.5 m at 0.96% WO₃ in BB1 at Glory Rock Hole (Blight, 1983).

Mineralization

The 1333 mineral occurrences in the east Kimberley are shown on Figure 5 and on Plates 1 and 1a. The mineral occurrences on Plates 1 and 1a are grouped by mineral

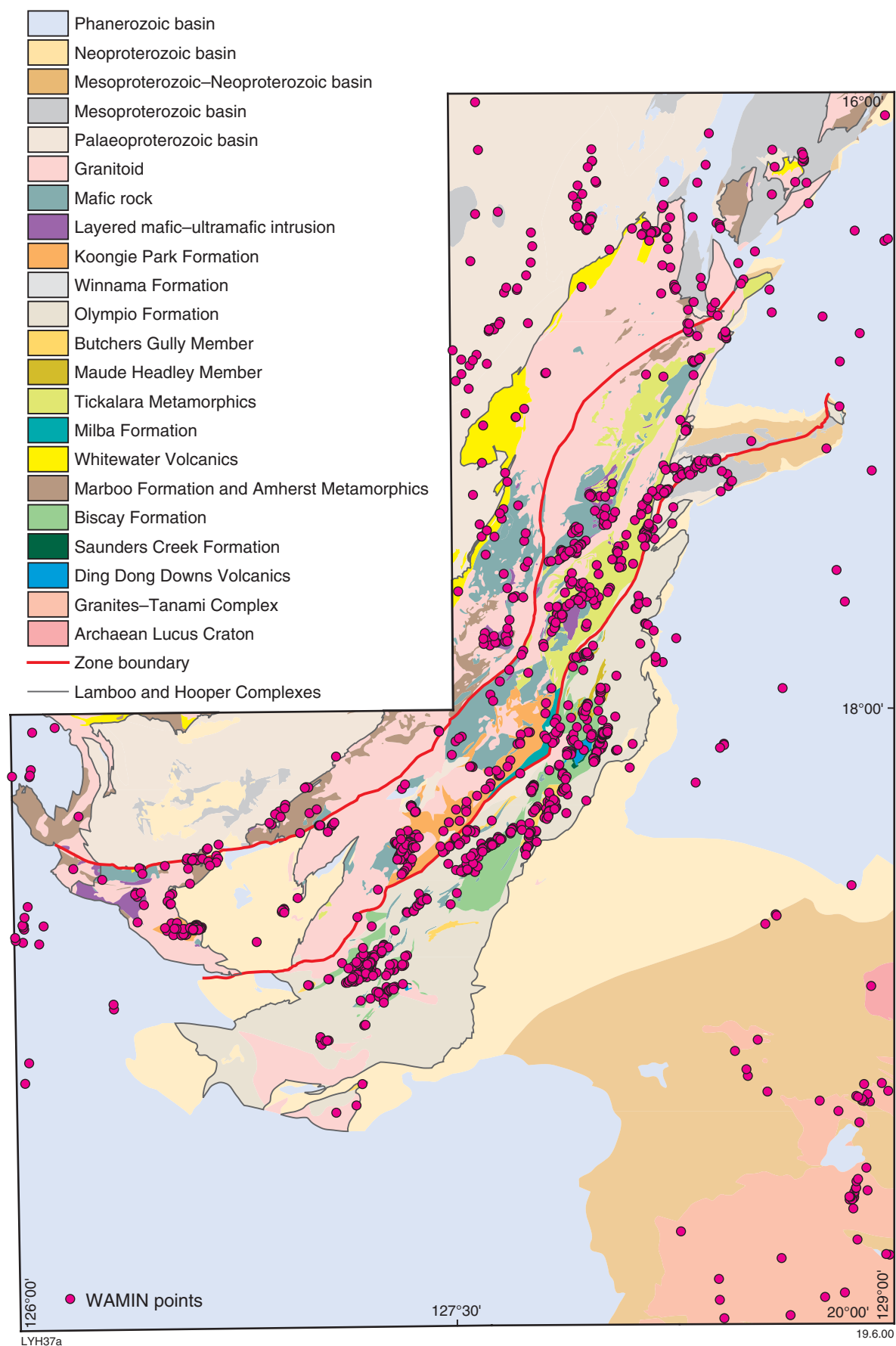


Figure 5. Tectonic sketch showing distribution of mineral occurrences in the east Kimberley. Scale — 1 cm : 20 km (approximately)

commodity and by mineralization style. Symbol colours on the Plates are used to distinguish commodity groups and symbol shapes are used to distinguish mineralization styles. For the convenience of description in the section below, mineralization style is used as a main heading followed by each of the commodity groups under various subheadings. Mineral occurrences referred to below are also identified by the WAMIN 'deposit number' shown thus: Palm Springs (2223).

Mineralization in kimberlite and lamproite intrusions

Only three mineral occurrences in kimberlite and lamproite intrusions have been identified in the east Kimberley area, but one of these, the Argyle AK1 pipe diamond deposit (1106; Fig. 6), is by far the most economically important mineral deposit discovered to date in the area.

Precious mineral — diamond

Argyle AK1

The Argyle AK1 pipe is a diatreme consisting of olivine lamproite pyroclastic and reworked pyroclastic rocks, intruded by rare olivine lamproite dykes. The pyroclastic rocks have been divided into 'sandy tuffs' that contain grains of country rock, and 'non-sandy tuffs' that do not contain any quartz grains. A detailed petrographic

description of the pipe is given in Jaques et al. (1986). The pipe intrudes Palaeoproterozoic to Mesoproterozoic sedimentary rocks of the Carr Boyd Group and Revolver Creek Formation, and has been dated at 1178 ± 47 Ma using Rb–Sr dating of phlogopite and the whole-rock isochron (Pidgeon et al., 1989). At the surface, the pipe is about 2 km long and varies from 500 m wide at the northern end to 150 m at the southern end. The shape of the pipe has been affected by post-intrusion faulting and regional tilting of about 30° to the north. Diamonds are most abundant in the sandy tuff, where they are present in juvenile lamproite clasts and as liberated grains in the matrix of the tuffs; diamonds are also present in the non-sandy tuff and the lamproite dykes (Boxer and Jaques, 1990). The pre-mining surface grade was about 5 cts/t, which is much higher than that of other diamond deposits around the world, but only about 5% of the diamonds are gem quality with the remainder being cheap-gem (40%) and industrial diamonds (55%; Jaques et al., 1986). Since mining commenced in 1985, recorded production to December 1999 has been 366.8 Mcts. The resources and reserves for Argyle AK1, as at 31 December 1999, are proved reserves of 62.9 Mt at 2.6 cts/t, probable reserves of 0.9 Mt at 3.3 cts/t, and measured, indicated, and inferred resources of 145.2 Mt at 2.9 cts/t (Ashton Mining Limited, 2000).

Lissadell Road Dykes

The Lissadell Road Dykes (1110) are an en echelon series of altered lamproite dykes and stringers that range from



LYH32

23.05.00

Figure 6. Argyle AK1 mine — terrace of east ridge wall (photograph courtesy of Argyle Diamonds)

a few millimetres to one metre in width and have been emplaced in tension gashes resulting from east-northeast shear movement (Jaques et al., 1986). Moonstone Diamond Corporation recovered 500 macrodiamonds during bulk testing of the dykes in 1992, with spot grades up to 24 cts/t (Bester, 1999).

Maude Creek Dyke

CRA Exploration recovered a small diamond (0.05 cts) from a 4.6 t sample from the Maude Creek Dyke (1115; Fielding, 1983) and five microdiamonds from two samples totalling 80 kg from this dyke (Smith, 1978; Fielding, 1983). The dyke intrudes the Hart Dolerite and is 1.5 m wide. An aerial photo-lineament, together with the distribution of kimberlitic indicator minerals, suggests that the dyke extends for about 2 km (Smith, 1978). Petrographically, the dyke resembles a kimberlite in that it contains abundant picroilmenite and pyrope macrocrysts, dunitic xenoliths, and traces of diamond; however, the spinel composition is not typical of kimberlite (Jaques et al., 1986).

Mineralization in carbonatite and alkaline igneous intrusions

Only four mineral occurrences have been classified as carbonatite and alkaline igneous intrusions. By far the largest of these is the Cummins Range (REE) carbonatite deposit (1119). The Cummins Range intrusion is a composite, subvertical stock consisting of an outer layer of pyroxenite and inner zones of carbonated mica-rich pyroxenite with numerous steeply dipping carbonatite veins that have been weathered to silicified ironstone breccia at the surface. The stock measures 1.8 by 1.7 km in plan (Fitton, 1981; Andrew, 1990). The intrusion has been dated at 800 Ma (Pidgeon et al., 1989). Both sovite (calcite dominant) and beforosite (dolomite dominant) carbonatite phases are present. The carbonatite is anomalous in REE, P_2O_5 , and tantalum. Zircon, sphene, baddeleyite, monazite, aeschynite, pyrochlore, columbite, and allanite are the main accessory minerals in the carbonatite (Andrew, 1990). There is an indicated resource of 6.3 Mt to 50 m depth at 0.5–1.0% total REO using a 0.5% cutoff (4.6 Mt at 1.0–2.0% if 1% cutoff is used; 0.6 Mt at 2.0–2.5% if 2.0% cutoff is used; Weir, 1989). The European grade of 50 to 100 ppm Eu_2O_3 is comparable with that of the Mountain Pass deposit in California (Weir, 1989). However, the overall REO grade was considered too low to be mined economically (Andrew, 1990). Cummins Range (REE 2) (5961) is a possible satellite body of the main Cummins Range (REE) deposit as it contains anomalous P_2O_5 and REE (Fitton, 1981; Dundas Gold Corporation NL, 1984).

Pods of carbonatite breccia with anomalous REE have been found in a small altered syenite intrusion that intrudes the Tickalara Metamorphics at Copperhead (1122; Rugless and Pirajno, 1994, 1996). A possible carbonatite dyke is also associated with fluorite mineralization at West Ridge (2399; Alvin, 1993; Rogers, 1998).

Orthomagmatic mafic and ultramafic mineralization

Mineralization in layered mafic intrusions

A total of 188 occurrences have been classified as 'orthomagmatic mafic and ultramafic — layered mafic'. Of these, 176 are in the Central zone of the Lamboo Complex; the remaining 12 are in the Western zone. The occurrences have been divided according to the main commodity into steel-industry metal (94 occurrences), precious metal (52 occurrences), base metal (35 occurrences), industrial mineral (six occurrences), and construction material (one occurrence).

Layered mafic–ultramafic intrusions in the Halls Creek Orogen occur in the Central and Western zones of the Lamboo Complex (Fig. 7). According to Hoatson (1997, in press) the intrusions were emplaced during at least three separate periods, and they may be divided into seven major groups on the basis of their field relationships, geochronology, style and intensity of deformation, and types of mineralization (see **Regional Geology**, p. 4, and Fig. 8). Significant mineralization occurs in two of Hoatson's seven groups: the 1855 Ma Group I (e.g. Panton intrusion) and the 1845 Ma Group V (e.g. Sally Malay intrusion). There is also some mineralization in the 1855 Ma Group II (e.g. Springvale intrusion) and the Group IV Wild Dog Creek Gabbro. There is potential for mineralization in the 1830 Ma Group VI (e.g. McIntosh intrusion).

Significant mineralization also occurs in the Lamboo intrusion and the Eastman Bore intrusions, which intrude the c. 1845 Koongie Park Formation, but these are located outside the area studied by Hoatson (1997, in press). In addition, there are mineralized layered mafic intrusions within the Tickalara Metamorphics that have been metamorphosed to granulite facies and that pre-date the seven groups of Hoatson (1997).

Steel-industry metal — nickel (copper, cobalt)

Most of the nickel occurrences are in Group V intrusions in the Central zone of the Lamboo Complex and the most significant is the Sally Malay deposit (653). The deposit is hosted by norite at the base of the Sally Malay intrusion where mineralization occurs as two keel-shaped lenses of massive pyrrhotite–pentlandite–chalcopyrite, 3–40 m thick. The massive sulfides are overlain by minor stringer-style chalcopyrite–pyrrhotite–pentlandite mineralization (Shedden and Barnes, 1996). On the basis of their textures, compositions, and basal distribution, Thornett (1981) suggested that the sulfides in the Sally Malay deposit formed by the gravity segregation of an immiscible sulfide liquid, with the sulfides possibly being derived from graphite–pyrrhotite gneisses within the Tickalara Metamorphics. The Sally Malay deposit has an indicated and inferred resource of 3.85 Mt averaging 1.79% Ni, 0.73% Cu, and 0.10% Co (Shedden and Barnes, 1996). Group V (Sally Malay) intrusions also host nickel–copper sulfide mineralization at the Keller Creek (e.g. 642),

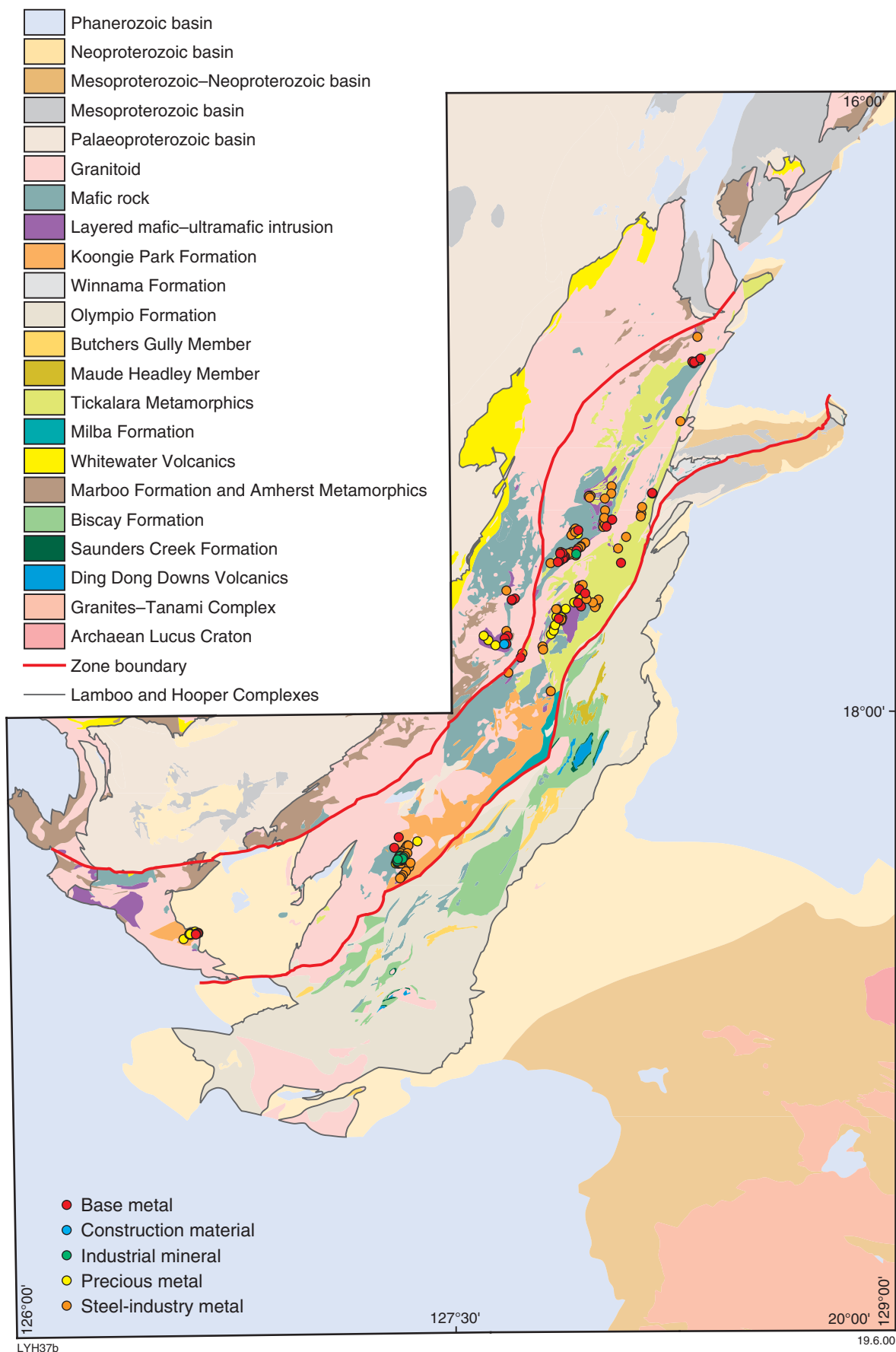


Figure 7. Tectonic sketch showing distribution of mineralization associated with layered mafic–ultramafic intrusions. Scale — 1 cm : 20 km (approximately)

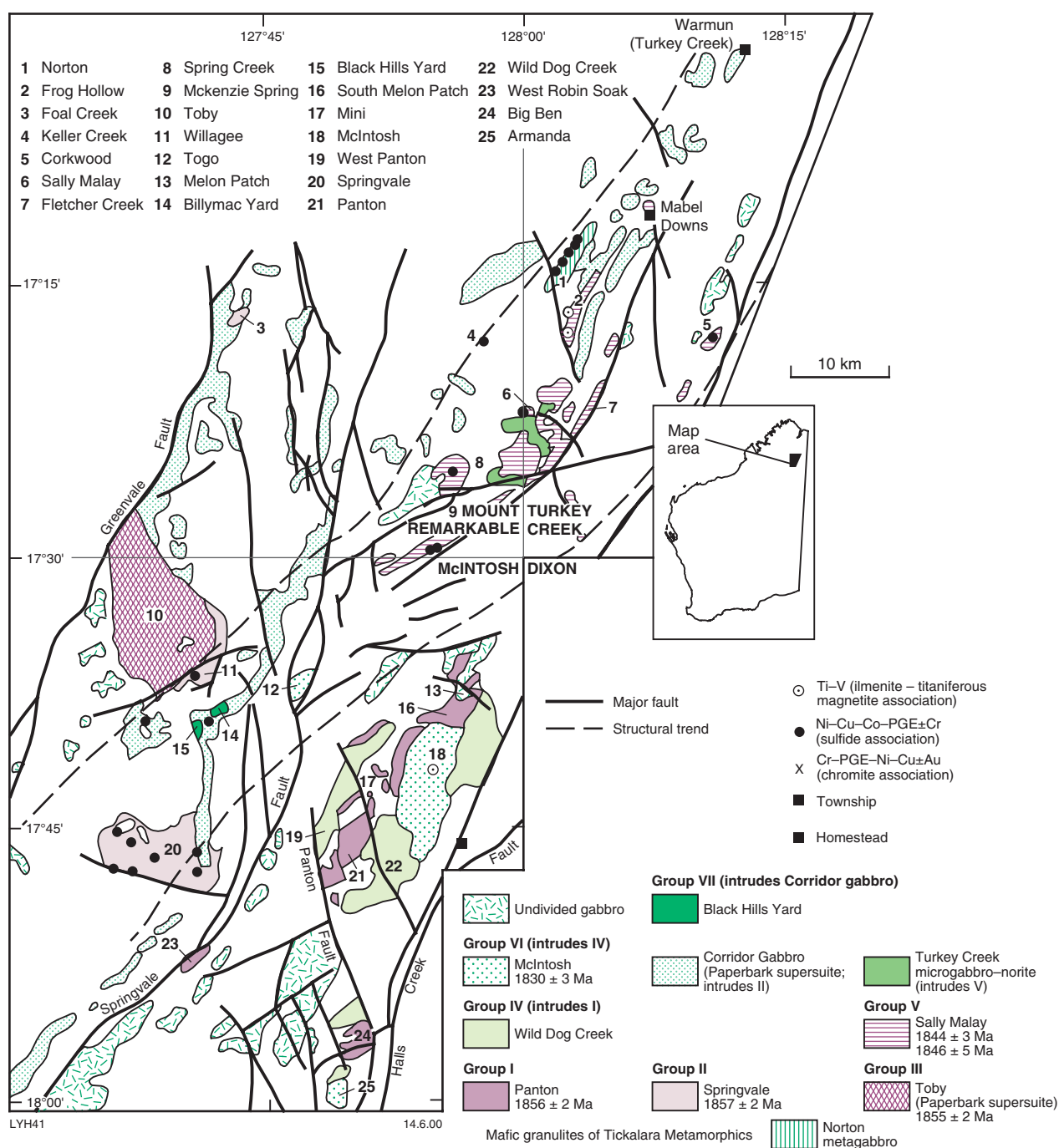


Figure 8. Distribution and classification of layered mafic intrusions in part of the Lamboo Complex (after Hoatson, 1997)

McKenzies Spring (e.g. 661), Corkwood (e.g. 641), Bulldust Flat (762), and Melon Patch prospects (e.g. 663). Nickel–copper–cobalt-bearing gossans have also been found in the Group V Spring Creek intrusion (e.g. Spring Complex 2, 7; and Coolumbooloo Yard, 4703). At Spring Complex 3 (10), a nickel–copper–cobalt gossan is interpreted to be related to mobilized sulfides in xenoliths of ultramafic–mafic rocks of the Spring Complex within granitoid (Lemcke, 1981).

A layered mafic–ultramafic intrusion is clearly recognizable in Tickalara Metamorphics in the Central zone of the Lamboo Complex in the Bow River area

despite granulite facies metamorphism. This intrusion contains low-grade, primary disseminated sulfides in norite, and also bands of massive pyrrhotite and minor chalcopyrite, which could be primary or metamorphic in origin at Bow River (Ni–Cu) (667). The best percussion-drill intersection obtained by Pickands Mather and Company International was 4.57 m at 1.22% Ni and 0.43% Cu (Lewis, 1966). Gossans have also been found in ultramafic rocks of the Tickalara Metamorphics to the north-northeast of Alice Downs at prospects known as Cabernet (639), Shiraz (2868), and Malbec (2869). At Cabernet, gossan samples assayed up to 7.23% Ni and 3.7% Cu and the best intersection was 16 m at 0.9% Ni

and 1.4% Cu, including 4 m at 2.38% Ni and 1.3% Cu (Mazzoni, 1976).

Nickel–copper mineralization has also been recorded from the Group IV Wild Dog Creek Gabbro (e.g. Mabel Hill (Ni–Cu) 1; 5285) where the best drill intersection was 3.8 m at 1.43% Ni and 0.3% Cu (Shaw, 1978).

In the Western zone of the Lamboo Complex, a mineralized norite with up to 40% matrix sulfides (predominantly pyrrhotite with minor chalcopyrite and pentlandite) was intersected in a percussion drillhole at Billymac Yard D (5288; Western Mining Corporation Limited, 1980). This unit has been mapped as part of the Paperbark supersuite on McIntosh; the unit is referred to as the Corridor gabbro by Hoatson (1997). Minor nickel–copper-bearing gossan has been found in gabbro and norite of the Paperbark supersuite at Springvale 7 (4758).

Steel-industry metal — chromium (PGE)

Most chromium occurrences are in the Central zone of the Lamboo Complex, in the Group I Panton intrusions (e.g. Panton 2, 5272; Mini M, 819; Big Ben, 817), and in the younger Lamboo intrusions (e.g. Lamboo (Cr 1), 831). The chromite typically occurs as stratiform chromitite layers ranging in width from a few centimetres to 2.4 m. In the Western zone, chromium is also present in the Group II Springvale intrusion (e.g. Springvale 1, 814) and Wilagee intrusion (e.g. Toby South, 812). Where significant PGE grades have been reported in addition to chromite, the occurrences have been classified as precious metal (see **Precious metal — PGE (gold, nickel, chromium)** below).

Steel-industry metal — titanium

Ilmenite-rich bands and lenses up to 150 m long and 30 m wide have been reported from gabbro and norite of the ?Sally Malay intrusions in the Ord Crossing area (838, 5198, 5199, 5200; El-Ansary, 1971). Lenses of ilmenite are also present in one of the Sally Malay intrusions at Frog Hollow North Lens (842) and Frog Hollow South Lens (843; Hoatson, 1995).

Precious metal — PGE (gold, nickel, chromium)

Occurrences of PGE are almost entirely restricted to the Central zone of the Lamboo Complex. Of these, many PGE occurrences are associated with chromitite layers in Group I Panton intrusions. The most significant of these occurrences is the Panton deposit (825) in the Panton intrusion. This intrusion consists of a lower ultramafic unit and an upper gabbroic unit that includes gabbro, gabbro, and anorthosite. There are five stratiform chromitite layers, each with a strike length of greater than 4.5 km, within the lower ultramafic unit of the intrusion. The uppermost of these layers, designated the A-seam, has an indicated and inferred resource of 2.0 Mt at 6.02 g/t PGE–Au and 0.28% Ni (Perring and Vogt, 1991). Other examples of PGE mineralization associated with chromitites in Group I Panton intrusions include Mini M (Cr–PGE) 2 (12), Togo 5A (828), Alice Downs 4 (3571), and Melon Patch 7 (3561).

Chromitite seams in the strongly folded c. 1810–1830 Ma Lamboo intrusion also contain PGE mineralization (e.g. Lamboo (PGE 1), 6142) where PGE minerals (including sudburyite, sperrylite, geversite, and stibio-palladinite) are enclosed within violarite and silicates in chromitite layers adjacent to, or within, coarse-textured olivine orthocumulate (O'Shea, 1987).

Significant PGE mineralization has also been recorded from chromitite horizons within the Eastman Bore intrusion (e.g. 835, 1044, and 5705; Union Corporation (Australia) Proprietary Limited, 1975; Barrett, 1982).

In the Western zone of Lamboo Complex, minor PGE mineralization is associated with chromitites in leuco-troctolite and anorthosite in the Group II Springvale intrusion (e.g. Springvale 2, 3517).

Base metal — copper (nickel, PGE, chromium)

In the Central zone of the Lamboo Complex, copper-rich gossans, which probably represent remobilized magmatic sulfides, are associated with Group I Panton intrusions (e.g. Mini M (Cu), 3511; MI-2b, 5370; Melon Patch 1, 663); Group V Sally Malay intrusions (e.g. McKenzies Spring 2, 4737) and Spring Creek intrusion (e.g. Coolumbooloo Yard 2, 5268); and the Group VI McIntosh intrusion (e.g. McIntosh A, 766).

In the Western zone of the Lamboo Complex, copper-bearing gossans are present in the Group II Springvale intrusion (e.g. Springvale 5, 4756) and in younger gabbro of the Paperbark supersuite, referred to as the Corridor gabbro by Hoatson (1997; e.g. Billymac Yard A, 806).

Industrial mineral — asbestos

Chrysotile veins occur in serpentinized peridotite of the Lamboo intrusion in the Central zone of the Lamboo Complex. An order of magnitude estimate of the resources at the Lamboo prospects (Lamboo (asbestos 2), 5486; Lamboo (asbestos 3), 5487; and Lamboo (asbestos 4), 5488) is 15 to 20 Mt containing 7–10 volume percent chrysotile (Betts, 1976; but note that this does not conform to the present JORC code). An asbestos seam has also been reported from a carbonated ultramafic (?Sally Malay) intrusion at Ord Crossing (asbestos) (5201).

Industrial mineral — talc

Pickands Mather and Company International (1965) reported 'reasonably high grade' talc in a shear zone in the Lamboo intrusion at Lamboo (talc) (5103).

Construction material

There is black dimension stone, reported to have characteristics superior to the Black Hill Norite in South Australia, in the Springvale intrusion at McIntosh B (2765). The rock is a magnetite gabbro and has a uniform grain size: about 1–1.5 mm for feldspar and 3 mm for pyroxene (Temby, 1994).

Undivided mineralization

Occurrences in the East Kimberley area that are associated with mafic dykes and intrusions, which are not obviously layered, have been classified as ‘orthomagmatic mafic and ultramafic — undivided’. A total of 11 occurrences have been identified including seven base metal occurrences, three steel-industry metal occurrences, and one construction material occurrence.

Base metal — copper, zinc, lead (silver)

In the Central zone of the Lamboo Complex, copper–nickel–PGE mineralization has been intersected in amphibolite of the Tickalara Metamorphics at Eileen Bore (764); the mineralization has probably been mobilized during metamorphism. A gossan with small amounts of malachite and visible gold is associated with amphibolite at Louisa Downs 3 (6312). At Emull (1147), base metal mineralization occurs in serpentinite close to its contact with the Koongie Park Formation. It is not certain whether the sulfides are primary in origin or whether they represent assimilated volcanogenic mineralization (Sanders, 1999).

In the Red Rock Basin, minor malachite and azurite has been reported from basalt or dolerite of the Red Rock Formation at Fish Hole Waterhole (2483) and Osmond Range West (2485; Dow and Gemuts, 1969).

Steel-industry metal — nickel, cobalt (copper)

A nickel–copper-bearing gossan containing malachite and secondary copper sulfides is associated with gabbro and pyroxenite of the Sally Downs supersuite at Osmond Valley (Ni–Cu) (810). Anomalous cobalt is associated with small veinlets of pyrrhotite in a fine-grained mafic dyke that intrudes gabbro and norite of the Sally Downs supersuite at McKenzies Spring 10 (5204).

Construction material

Gabbro suitable for building stone is present at Halls Creek North (6474).

Disseminated and stockwork mineralization in plutonic intrusions

Base metal — copper (molybdenum, silver, gold)

A total of seven mineral occurrences have been classified as ‘disseminated and stockwork in plutonic intrusions’. These occurrences are all base metal occurrences in the Central zone of the Lamboo Complex. It is not certain whether these occurrences are porphyry copper deposits or whether they should be classified as ‘vein and hydrothermal — undivided’.

At Mount Angelo Porphyry (1098), mineralization is hosted by sheared Angelo Microgranite that shows

evidence of silicic, sericitic, and chloritic alteration. The microgranite is probably the subvolcanic equivalent of felsic volcanic rocks in the Koongie Park Formation (Sheppard et al., 1995). Chalcopyrite occurs as disseminations and veinlets with minor bornite, tetrahedrite, pyrite, sphalerite, and molybdenite in the primary zone and malachite, chrysocolla, and minor native copper in the oxidized zone (Blain, 1969). The deposit has an indicated resource of 144 Mt at greater than or equal to 0.2% Cu (Sewell, 1999). Anomalous molybdenum and silver levels are also present, with intersections up to 3 m at 50.4 g/t Ag and 3.05 m at 0.25% Mo reported (Pickands Mather and Company International, 1967b). Gold values are subeconomic; the best intersection is 8 m at 0.2 g/t (Sewell, 1999).

The McHale Granodiorite of the Sally Downs supersuite hosts the McHales (1093) and Killarney occurrences (e.g. Killarney 1, 1087). At these occurrences, mineralization occurs as vein and fracture fillings and as disseminated copper mineralization in granitoid. The granitoid shows evidence of propylitic, sericitic, potassic, silicic, and argillic alteration and a number of magmatic–hydrothermal breccia dykes have been recognized (Witt and Sanders, 1996). The mineralization has been interpreted as porphyry copper by Stratin Minerals Proprietary Limited (1972) and Witt and Sanders (1996). An alternative interpretation (Scott, 1973) is that it is related to volcanic rocks unconformably overlying the granite (Texas Downs Formation).

Pegmatitic mineralization

A total of seven pegmatitic occurrences have been recorded. They are all at the southern end of the Eastern zone of the Lamboo Complex and include six speciality metal occurrences and one steel-industry metal occurrence.

Speciality metal — tin, tantalum, niobium, lithium, beryl

A pegmatite at Columbian Creek (5754) contains scattered crystals of spodumene, beryl, columbite, and cassiterite. Cassiterite has also been recorded from small pegmatites at Mount Heartbreak F (926), Mount Cross (Sn 1) (969), Mount Cross 4 (6380), Junda 3 (5914), and Minneroo Pool 1 (976). An imprecise SHRIMP date of 1740 ± 40 Ma has been obtained on one of these pegmatites (AGSO OZCHRON database).

Steel-industry metal — tungsten

At Junda 1 (961), there are moderately coarse crystals (4 to 20 mm across) of scheelite in a pegmatitic vein (300 × 3 m) within mica schist; the scheelite is restricted to a zone about 15 m in length.

Skarn mineralization

A total of 22 occurrences have been classified as skarn, but the genesis of many of these occurrences as ‘typical

skarns' is doubtful (see **Mineral controls and exploration potential**, p. 49). The occurrences have been subdivided according to their major commodity into base metal (15), steel-industry metal (4), speciality metal (2), and industrial mineral (1).

Base metal — copper, zinc, lead, silver (molybdenum, tin, tungsten)

Most of the base metal occurrences classified as skarn (12) are within calc-silicate horizons of the Tickalara Metamorphics, in the Central zone of the Lamboo Complex. Examples are Melon Patch Well A (1050), White Rock Well (1085), and Wills Creek B1 (1063). There is no obvious association with granitic rocks except at Wills Creek B6 (1078), where Western Mining Corporation Limited (1979) reported malachite and chalcopryrite in quartz–garnet–scapolite–diopside skarn, containing up to 16.5% Cu and 80 ppm Ag, associated with quartz veins and slivers of foliated tonalite (?Monkey Yard Tonalite).

Chadwick (1984) reported zinc and lead mineralization in weathered epidote–garnet–diopside–calcite skarn within the Koongie Park Formation at BB Hill (3298). Gossan has also been reported from limestone of the Winnama Formation near its contact with dolerite at Osmond Valley (Zn 7) (5530; Codner, 1973).

In the Eastern zone, epidote–calcite–titanite–diopside–vesuvianite–andradite–grossular skarn with anomalous copper, silver, and molybdenum values has been reported from the Biscay Formation in the Taylor Lookout area at Skarn A (5743; Mazzoni, 1979).

Steel-industry metal — tungsten (tin)

Two scheelite-bearing skarns have been reported from the Koongie Park Formation, in the Central zone of the Lamboo Complex, at Central Skarn (3299) and Southern Skarn (3300).

In the Eastern zone of the Lamboo Complex, scheelite- and cassiterite-bearing skarn occurs in the Butchers Gully Formation close to a pegmatite at Mount Heartbreak A (887). Scheelite-bearing calc-silicate rock and carbonated mafic tuff of the Ding Dong Downs Volcanics at Frog Creek (930) have been interpreted as skarn mineralization by Sanders (1999).

Speciality metal — tin (silver)

In the Eastern zone of the Lamboo Complex, cassiterite-bearing skarns have been reported from the Butchers Gully Member at Mount Cross (Sn 2) (972) and from the Biscay Formation at Taylor Lookout 7 (5748).

Industrial mineral — garnet

Garnet-bearing skarn or calc-silicate rock has been reported from the Tickalara Metamorphics at Garnet Hills (5613) in the Central zone of the Lamboo Complex.

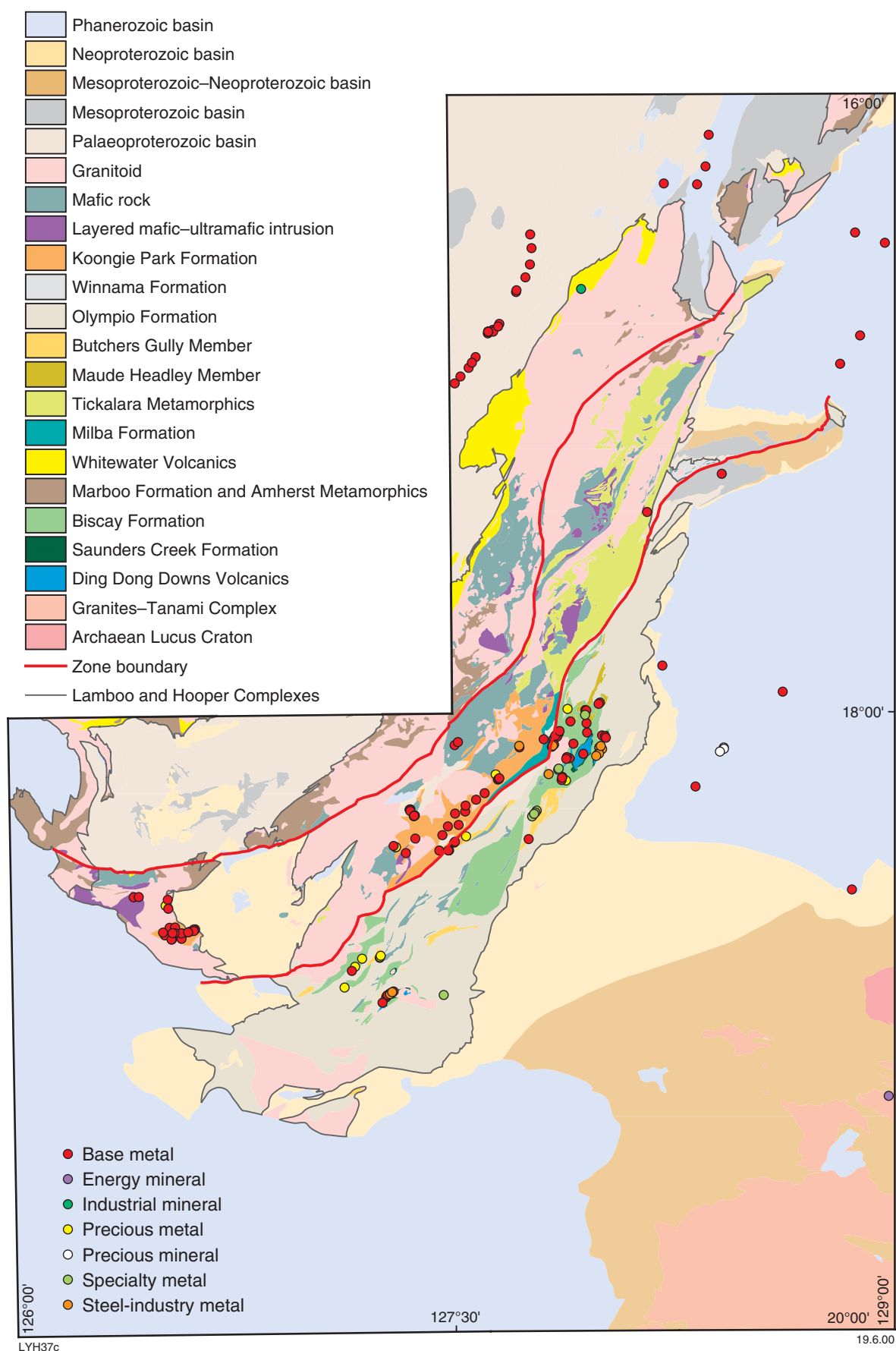
Stratabound volcanic and sedimentary mineralization

In the east Kimberley area, occurrences included under 'stratabound volcanic and sedimentary' include: VMS deposits associated with proximal felsic volcanic rocks; volcanic exhalative deposits associated with distal volcanoclastic sediments; REO and tantalum mineralization associated with alkaline volcanic rocks; native copper mineralization in mafic volcanic rocks; and mineralization associated with interflow sedimentary rocks and tuffs in mafic volcanic units. In the database, the deposits in this category have been subdivided into volcanic-hosted sulfide, sedimentary-hosted sulfide, volcanic oxide, and undivided. However, this classification has been somewhat arbitrary because many of the deposits are hosted by tuffaceous sedimentary rocks or a mixture of rock types. Consequently, 104 of the 171 occurrences in this category have been classified as undivided, but many of these are genetically related to those classified as volcanic-hosted sulfide or sedimentary-hosted sulfide. For the purposes of discussion, all of the occurrences are considered under the main heading of 'stratabound volcanic and sedimentary' and the subheadings of the main commodity types present. There is a total of 131 base metal, 14 steel-industry metal, 13 precious metal, eight speciality metal, one energy mineral, one industrial mineral, and three precious mineral occurrences in the stratabound volcanic and sedimentary category (Fig. 9).

Base metal — copper, lead, zinc, silver

A significant number (53) of the base metal occurrences in the stratabound volcanic and sedimentary category are within the Koongie Park Formation in the Central zone of the Lamboo Complex. The Koongie Park Formation in the Koongie Park area has been informally divided into three members by Sewell (1999): Coolibar tuff member; Camp shale member; and Weldon Creek lava member. The basal part of the Camp shale member, termed the 'Mimosa interval' by Sewell (1999), consists of intercalated chert, chloritic schist, banded magnetite or pyrrhotite–chlorite rock, black shale, impure dolomite, and andesitic lava and volcanoclastic rocks. The 'Mimosa interval' hosts the Sandiego (1125) and Mount Angelo North (1137) deposits and many other smaller VMS prospects in the Koongie Park area.

At Sandiego, zinc–lead–copper–silver(–gold) mineralization occurs along the western limb of a fold over a strike length of 350 m. There is a crude zonation at Sandiego from massive sphalerite, pyrrhotite, and accessory galena at the base of the deposit to a zone of stringers of chalcopryrite and pyrite at the top of the deposit. The inferred resource at Sandiego includes 0.335 Mt at 6.71% Cu and 288 g/t Ag in the supergene zone, and 4.3 Mt at 0.51% Cu, 0.83% Pb, 7.9% Zn, and 31 g/t Ag in the primary sulfide zone (Sewell, 1999). At Mount Angelo North, massive copper–silver mineralization is hosted by shales within a tuffaceous horizon in the 'Mimosa interval' and disseminated and stringer copper mineralization is present in the overlying Weldon Creek lava member. The



inferred resource for Mount Angelo North includes: 0.076 Mt at 1.0% Cu and 9.6 g/t Ag in the gossan zone; 0.035 Mt at 1.1% Cu and 45.0 g/t Ag in the Ag-rich supergene zone; 0.011 Mt at 7.7% Cu and 19.1 g/t Ag in the Cu-rich supergene zone; and 0.2 Mt at 2.0% Cu in the primary sulfide zone (Sewell, 1999).

In the Louisa Downs area, subeconomic base metal mineralization is associated with distal volcanoclastic sediments, chert, and tuffs of the Koongie Park Formation (e.g. Landrigan Creek, 1139; Eastman Yard 1, 5398; and Bullock Bore 1, 1140). At Landrigan Creek, a copper–zinc–silver–gold zone is underlain by a lead–molybdenum zone (Platts, 1983).

In the Tickalara Metamorphics, the only stratabound volcanic and sedimentary base metal occurrence reported is the Tickalara occurrence (1252) where malachite-bearing gossan occurs in quartz–muscovite–feldspar schist interpreted by Hoyle (1976) as a felsic volcanic or pyroclastic rock.

At Morgans Grave (1300 and 5959), to the northwest of Louisa Downs, gossans are associated with quartz–chlorite–sericite schist within an area mapped as amphibolite with intercalated metasedimentary rocks on MOUNT RAMSAY. Here, massive pyrrhotite was intersected in a diamond drillhole. The mineralization was interpreted as volcanogenic by Menzies and Ross (1982) but as a mineralized shear by Sanders (1999).

Most of the stratabound volcanic and sedimentary base metal occurrences in the Eastern zone of the Lamboo Complex are in the Biscay Formation (31 occurrences) with a few occurrences in the Ding Dong Downs Volcanics (four occurrences), the Butchers Gully Member (two occurrences), and the Maude Headley Member (two occurrences).

The most important occurrences are Ilmars Central (1180) and Little Mount Isa (1188). At Ilmars Central, zinc–lead–copper–silver–gold mineralization is in the Biscay Formation at the contact between chlorite–quartz schist and calc-silicate rocks close to felsic volcanic rocks (Sewell, 1999). Little Mount Isa is possibly on the same stratigraphic horizon as Ilmars Central. At Little Mount Isa, two parallel zones of gossan are separated by 6 m of chlorite schist. The gossan horizons change to pyritic black shale at depth (Sewell, 1999). Hemimorphite and cerussite-rich gossans are associated with meta-rhyolite and chert 4 km south along strike from Ilmars Central at Twin Hills (1192, 5728; Sewell, 1999). Gossans in chloritic and graphitic schists and dolomite of the Biscay Formation at Gentle Annie B (1202) are similar in style to Ilmars (Harris, 1967). At Taylor Lookout 6 (5745), there are small gossanous stringers in felsic volcanic rock within the Biscay Formation.

Gossans with minor amounts of base metals have been reported in association with felsic volcanic rocks of the Maude Headley Member (e.g. Palm Creek, 5377) and Butchers Gully Member (e.g. Ruby Queen – Brockman 2, 5074).

Native copper occurs in thin quartz veins and quartz–calcite amygdaloids within basalt of the Ding Dong

Downs Volcanics at Ding Dong Downs (Cu) (2429). Minor base-metal gossans are associated with the Ding Dong Downs Volcanics in felsic volcanoclastic rocks at Castle Creek 15 (5051) and in black shales at Sophie Downs (Zn) 3 (4918).

In the Kimberley Basin, widespread, but patchy, copper mineralization in interflow sediments and tuffaceous flow tops of basalt of the Carson Volcanics has been interpreted as volcanogenic by Planet Management and Research Proprietary Limited (1971). The most significant of these occurrences is Carson Volcanics 9 (2440) where mineralization is concentrated along a northeasterly and a northwesterly trending fault, and also in the interflow sedimentary rocks. An inferred resource of 1.8 Mt at 0.51% Cu has been estimated for this deposit (Australian Anglo American Limited, 1971).

In the Osmond Basin, a felsic lithic tuff near the base of the Bungle Bungle Formation at Paradise Creek (Cu B) (2517) was reported to contain 0.57 to 0.61% Cu and 4 ppm Ag (de Angelis, 1984). The copper occurs as chrysocolla and sky blue turquoise.

In the Ord Basin, rounded pieces of native copper up to 2.5 cm long have been found in creek beds and in vesicular basalt of the Cambrian Antrim Plateau Volcanics at Cuddys Find (2522). The native copper is sheathed in cuprite and has been interpreted by Erskine et al. (1970) to be Michigan-type copper — native copper in vesicular basalt and agglomerate resulting from remobilization of primary copper in a thick sequence of basalt into permeable layers during metamorphism (White, 1968). Native copper, cuprite, and chrysocolla segregations in basaltic agglomerate of the Antrim Plateau Volcanics can be traced over a strike length of 5 km at Panton-Elvire (Cu) (2505). Boulders of prehnite up to 1 m in diameter that have weathered out of basalt of the Antrim Plateau Volcanics at Dunham Crossing (2487) also contain native copper (Harms, 1959). Traces of chalcopyrite are associated with calcite in siliceous sinter of the Antrim Plateau Volcanics at Mount Wittenoom (5422; Clarke and Blockley, 1961).

Steel-industry metal — tungsten

In the Taylor Lookout area, at Taylor Lookout 1 (942), discontinuous scheelite mineralization occurs over a strike length of over 1 km within a calc-silicate horizon in the Biscay Formation in the Eastern zone of the Lamboo Complex. This mineralization has been classified as stratabound volcanic and sedimentary in this study, but Sanders (1999) interpreted it as skarn or hydrothermal mineralization related to the San Sou Monzogranite. Scheelite also occurs in the Ding Dong Downs Formation at Berthas Butt in the Eastern zone, where it is hosted by mafic tuff (e.g. Castle Creek 33, 5872), felsic tuff (e.g. Glory Rock Hole 8, 5846), and epidote-bearing chert (e.g. Castle Creek 22, 5837). This mineralization is almost certainly related to scheelite mineralization occurring in epidote-quartzite rock (of controversial origin) in the Berthas Butt area, and may be volcanic exhalative in origin (see **Stratabound sedimentary mineralization**, p. 31).

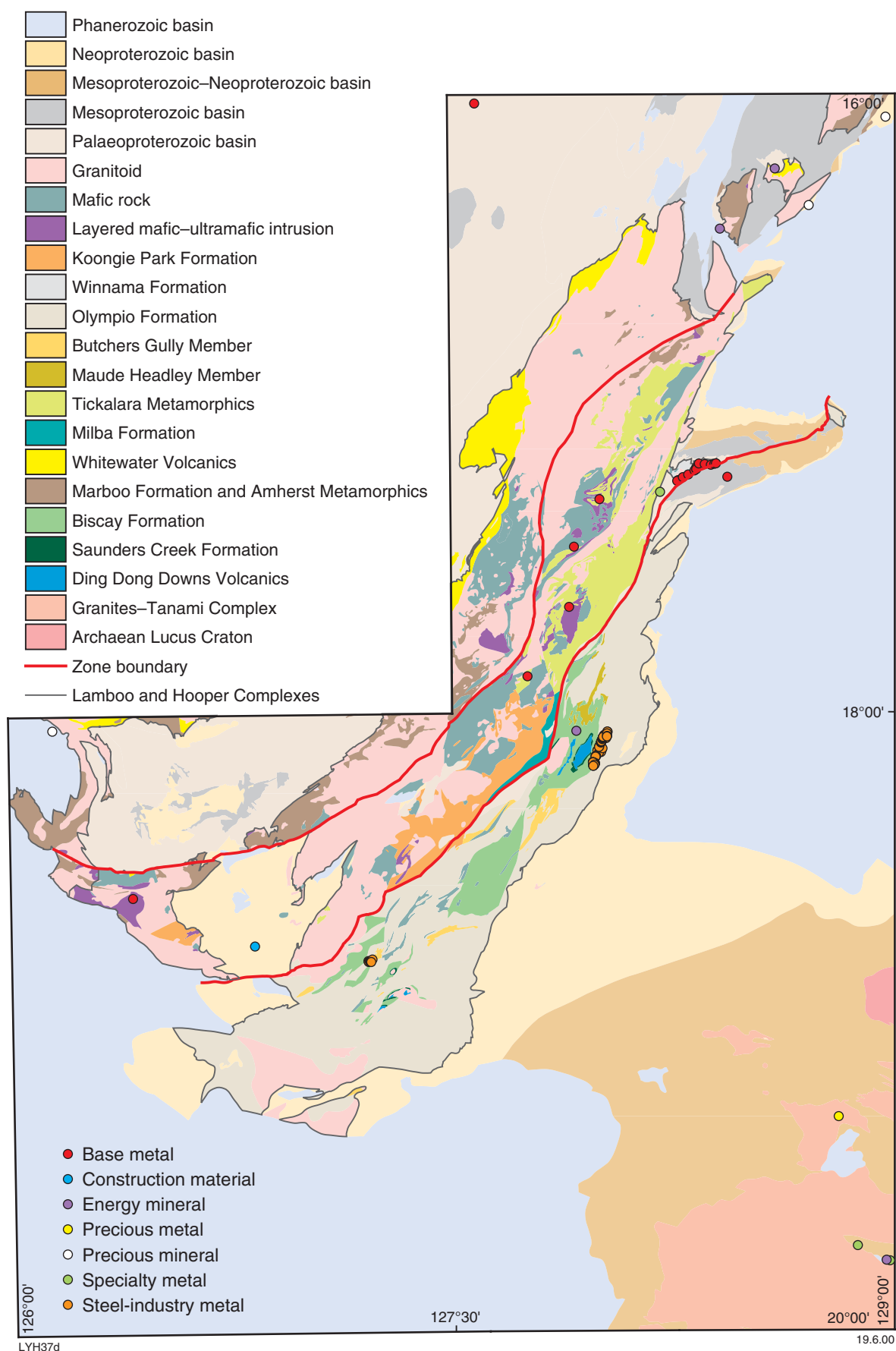


Figure 10. Tectonic sketch showing distribution of clastic-hosted stratabound sedimentary deposits. Scale — 1 cm : 20 km (approximately)

Precious metal — gold

There are several gold occurrences that are possibly volcanogenic within the Biscay Formation in the Eastern zone of the Lamboo Complex. Gold is associated with galena and malachite in a rock described as microdacite at Lulu (6440; Dam Gold NL, 1997). Fine-grained gold is associated with pyrite in BIF at Freda (6437; Dam Gold NL, 1997). At Mount Carmel East 2 (6094), gossanous siliceous basalt or basaltic andesite exposed in a pit contains anomalous gold, silver, and base metals (Robinson, 1989).

At Dockrell East 2 (6393), brecciated felsic tuff with fragments of quartz and feldspar in a green chloritic matrix is overlain by chert; both tuff and chert contain gold mineralization. This area was mapped as Olympio Formation on DOCKRELL (Tyler and Griffin, 1994) but has been reinterpreted as Butchers Gully Member on Plate 1. Minor gold mineralization is associated with black shale in the Butchers Gully Member at Koongie Rockholes 2 (6066).

Speciality metal — REE, niobium, tantalum, beryllium, gallium

There is an unusual style of REE–niobium–tantalum mineralization in the Eastern zone of the Lamboo Complex that is hosted by alkaline volcanoclastic rocks of the Butchers Gully Member, Maude Headley Member, and Biscay Formation. The most significant occurrence of this type is the Brockman deposit (1102, 6111, 6112, 6113) in the informally named ‘Niobium tuff’ of the Butchers Gully Member. The Niobium tuff has been interpreted as an ashflow tuff (Ramsden et al., 1993) or tuffaceous volcanoclastic deposit (Taylor et al., 1995). Taylor et al. (1995) suggests that the volcanic sequence at Brockman was erupted from a small volcanic shield complex probably in a rift-related basin in a shallow-marine setting. The Niobium tuff outcrops along the western flank of a southwest-plunging syncline over a strike length of 3.5 km and varies in thickness from 5 to 35 m (Ramsden et al., 1993). The mineralization is very fine grained (< 20 µm). Ore minerals include gel-zircon (which includes heavy REE and some niobium), thorite (which includes heavy REE), altered columbite, complex Y–REE niobates, fluorocarbonates (including bastnaesite, parisite, and synchisite that contain light REE), fluorite, bertrandite, gallium-bearing K-mica, and base metal sulfides (principally sphalerite; Ramsden et al., 1993). The Brockman deposit has a measured resource of 4.29 Mt at 0.44% Nb₂O₅, 0.027% Ta₂O₅, 0.124% Y₂O₃, 1.04% ZrO₂, 0.011% Ga, 0.035% HfO₂, and 0.09% REO, and an indicated and inferred resource of 45 Mt at similar grades (Chalmers, 1990).

Anomalous niobium is present in felsic volcanic rocks of the Maude Headley Member at Ding Dong Downs (Nb) (5380). Altered trachyte within mafic tuffs that have been mapped as Biscay Formation on HALLS CREEK contains anomalous niobium, fluorine, and zirconium at Sophie Downs 1 (1103) and Sophie Downs 2 (5916).

Industrial mineral — fluorite

At Archie Creek (2391), fluorite is present as an accessory mineral in rhyolite, as large dark-purple sheared crystals

in rhyolitic tuff, and in volcanogenic sedimentary rocks of the Whitewater Volcanics in the Western zone of the Lamboo Complex (Harrop, 1977). Fluorite is also present as an accessory mineral at Brockman (see above).

Precious mineral — prehnite

At the Valley occurrence (6406), prehnite fragments, which have been weathered from scoriaceous basalt of the Cambrian Antrim Plateau Volcanics, are concentrated on the surface along a north-northwesterly trending zone 150 m long and 40 m wide. The colour of the prehnite varies from green to golden yellow to white to clear. About 500 kg of good-quality material, suitable for cutting as a semiprecious stone or for carving as a jade substitute, has been produced from a small pit. A resource estimate of 174 t (not recognized under the JORC code) has been made (Mullumby, 1994). Similar occurrences are present at the Mound (6407), where approximately 15 kg of prehnite has been produced and a resource estimate of 138 t (not recognized under the JORC code) has been made (Mullumby, 1994), and at New Deposit (6409).

Energy mineral — uranium (copper)

At Mount Junction (6262), ferruginous chert with anomalous uranium and copper is associated with altered volcanic rocks and shale of the Killi Killi ‘Beds’ in the Granites–Tanami Complex (Pearcey et al., 1988). This is possibly volcanogenic in origin.

Stratabound sedimentary mineralization

Clastic-hosted mineralization

A total of 104 occurrences have been classified as ‘stratabound sedimentary — clastic-hosted’. These occurrences have been further subdivided according to their main commodity into steel-industry metal (73 occurrences), base metal (19 occurrences), energy mineral (four occurrences), speciality metal (three occurrences), precious mineral (three occurrences), precious metal (one occurrence), and construction material (one occurrence; Fig. 10).

Steel-industry metal — tungsten (molybdenum)

In the Eastern zone of the Lamboo Complex, in the Berthas Butt area, mineralization outcrops over a strike length of 20 km around the Castle Creek Anticline (Buckovic et al., 1982; Fig. 11). The mineralization is mainly hosted by epidote quartzite of the Saunders Creek Formation (informally called the ‘Upper quartzite’), such as at Glory Rock Hole 1 (868), Rock Pile 1 (849), and Snake Gully North (873). Some mineralization is also present in epidote quartzite within the Ding Dong Downs Volcanics (informally called the ‘Lower quartzite’), such as at Late Night (876) and Castle Creek 12 (5047). The mineralized epidote quartzites in both the Upper and Lower quartzites grade laterally into feldspathic sand-

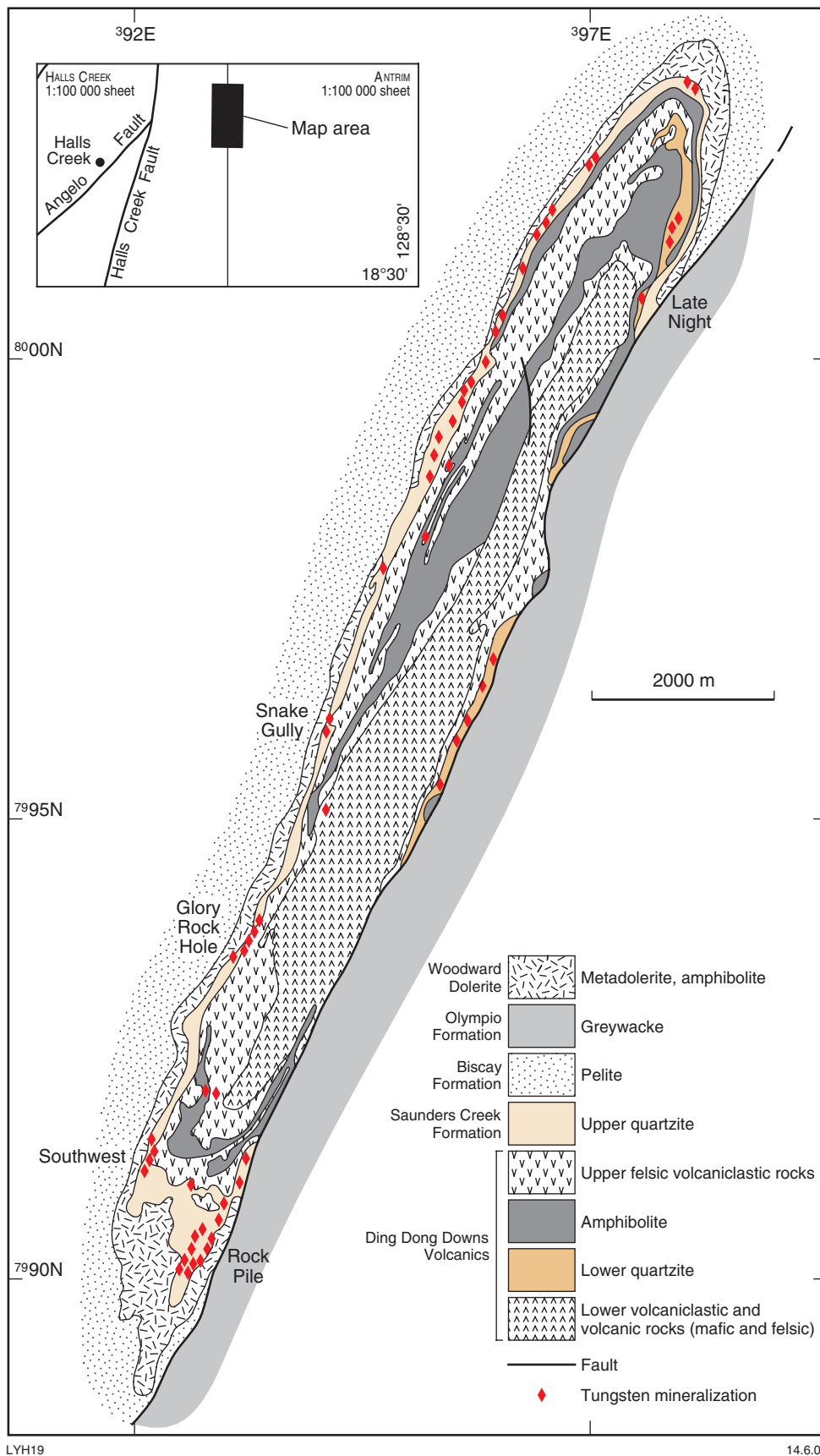


Figure 11. Geological map of the Castle Creek anticline showing tungsten occurrences near Halls Creek (after Buckovic, 1982; Blight, 1983; Ririe, 1989; and Sanders, 1999). Minor differences between the geology shown on this map and that on HALLS CREEK (Blake et al., 1999a) and ANTRIM (Blake and Warren, 1996) may be due to differences in the scale and interpretation of the mapping

stones and felsic crystal tuffs (Buckovic et al., 1982). The scheelite is disseminated within the quartzites as broad tabular zones; within the zones, the scheelite mineralization generally conforms well to stratification within the quartzites, although some mineralization is cross-cutting (Buckovic et al., 1982). Surface grades are up to 6.8% WO_3 (Buckovic et al., 1982), but the highest drill intersections were 0.5 m at 2.38% WO_3 and 1.4 m at 0.34% WO_3 from a diamond drillhole at Glory Rock Hole 2 (5840; Blight, 1983). Powellite (CaMoO_4) is associated with scheelite at Late Night (Blight, 1983).

Elsewhere in the Eastern zone, in the South Dockrell area, stratabound scheelite is associated with multiple thin beds of epidote quartzite and tourmalinites (e.g. South Dockrell 1, 884). The epidote quartzites are similar to those at Berthas Butt, but are thinner and much finer grained. The tourmalinites consist of interlaminated fine-grained quartz and fine euhedral tourmaline laths. The epidote quartzites and tourmalinites are interbedded with chloritic tuffaceous siltstones and minor mafic volcanoclastic rocks and amphibolites (Goldner, 1984). The rocks were interpreted as Biscay Formation by Goldner (1984), but are now considered to be part of the Olympio Formation (Tyler et al., 1998a). They probably belong to the Butchers Gully Member (Tyler, I. M., 1999, pers. comm.).

Base metal — copper, lead, zinc (silver)

In the Central zone of the Lamboo Complex, a number of base-metal bearing gossans have been reported from feldspathic sandstones of the Winnama Formation (e.g. Osmond Valley (Ag–Pb–Zn), 5531). There are also a few scattered occurrences of base-metal bearing gossans in metasedimentary rocks of the Tickalara Metamorphics (e.g. Little Panton (Cu), 5343).

In the Kimberley Basin, at Menuairs Dome 3 (4543), there is extensive malachite staining in shale at the base of the Middle Pentecost Sandstone. This copper enrichment continues over a strike length of 200 km but most of the dome is outside the east Kimberley area. No ore-grade values were intersected in exploration drilling, suggesting surface enrichment of weakly anomalous syngenetic copper (Australian Anglo American Limited, 1971).

In the Osmond Basin, native copper is hosted by clean, porous oxidized sandstone and siltstone of the Bungle Dolomite at Paradise Creek (Cu A) (2508). The best drilling intersection by Samin Australia was 25 m at 0.34% Cu (de Angelis, 1984).

Energy mineral — uranium

At Rustlers (5797), in the Eastern zone of the Lamboo Complex, anomalous uranium occurs within the Biscay Formation in ferruginous pods in chloritic sandstone close to its contact with carbonaceous schist.

In the Bonaparte Basin, sandstone of the Devonian Galloping Creek Formation hosts secondary uranium minerals at Horse 7 (5341).

In the Birrindudu Basin at Killi Killi 2 (5403), abundant uranium ochres are exposed in a pit within the basal conglomerate of the Gardiner Sandstone, but these extend only to a depth of 60 to 90 cm, suggesting that there has been surface enrichment (i.e. they are probably not typical sandstone-hosted uranium deposits). Uranium is also associated with xenotime in grit and conglomerate beds at the base of the Gardiner Sandstone at Killi Killi 1 (5316) and Killi Killi West (5317).

Speciality metal — REE

At Junior (5315), fine-grained sandstone of the Palaeoproterozoic Red Rock Formation in the Red Rock Basin contains abundant heavy minerals. Rock-chip samples assayed up to 3.9% total REE and 0.4% Th. X-ray fluorescence (XRF) scans showed a predominance of light REE (Ce and La) suggesting that the ore mineral is monazite (Campana et al., 1973).

Rare earth element mineralization is also present in xenotime associated with uranium in basal grits and conglomerate of the Gardiner Sandstone in the Birrindudu Basin at Killi Killi 1 and Killi Killi West (as mentioned above).

Precious mineral — diamond

Gem-quality diamonds are said to have been recovered by Bulletin Trading from the basal Barramundi Conglomerate at Barramundi Range in the Canning Basin (Davies, 1990); however, later work failed to find any diamonds (Sabminco NL, 1993).

Precious mineral — zebra stone

Zebra stone occurrences in the Ranford Formation of the Wolfe Creek Basin have been worked at Lake Argyle 1 (5611) and Lake Argyle 2 (6497). The stone is carved and sold to tourists in Kununurra, Wyndham, and Perth as ('precious') ornamental items.

Construction material — sandstone

Small quantities of sandstone of the Neoproterozoic Lubbock Formation in the Louisa Basin have been quarried at Berrangi Creek (6471) and used locally.

Carbonate-hosted mineralization

A total of 37 occurrences have been classified as 'stratabound sedimentary — carbonate-hosted'. All are base metal occurrences (Fig. 12).

Base metal — lead, zinc, copper (silver, gold, molybdenum, tungsten, tin)

Significant carbonate-hosted base metal mineralization is being mined at Kapok (6472) near the eastern end of the Lennard Shelf. The Kapok mine is part of the Lennard Shelf Project that includes the Cadjebut, Goongewa, and Pillara mines to the west of the study area boundary. The Mississippi Valley-type lead–zinc mineralization at Kapok,

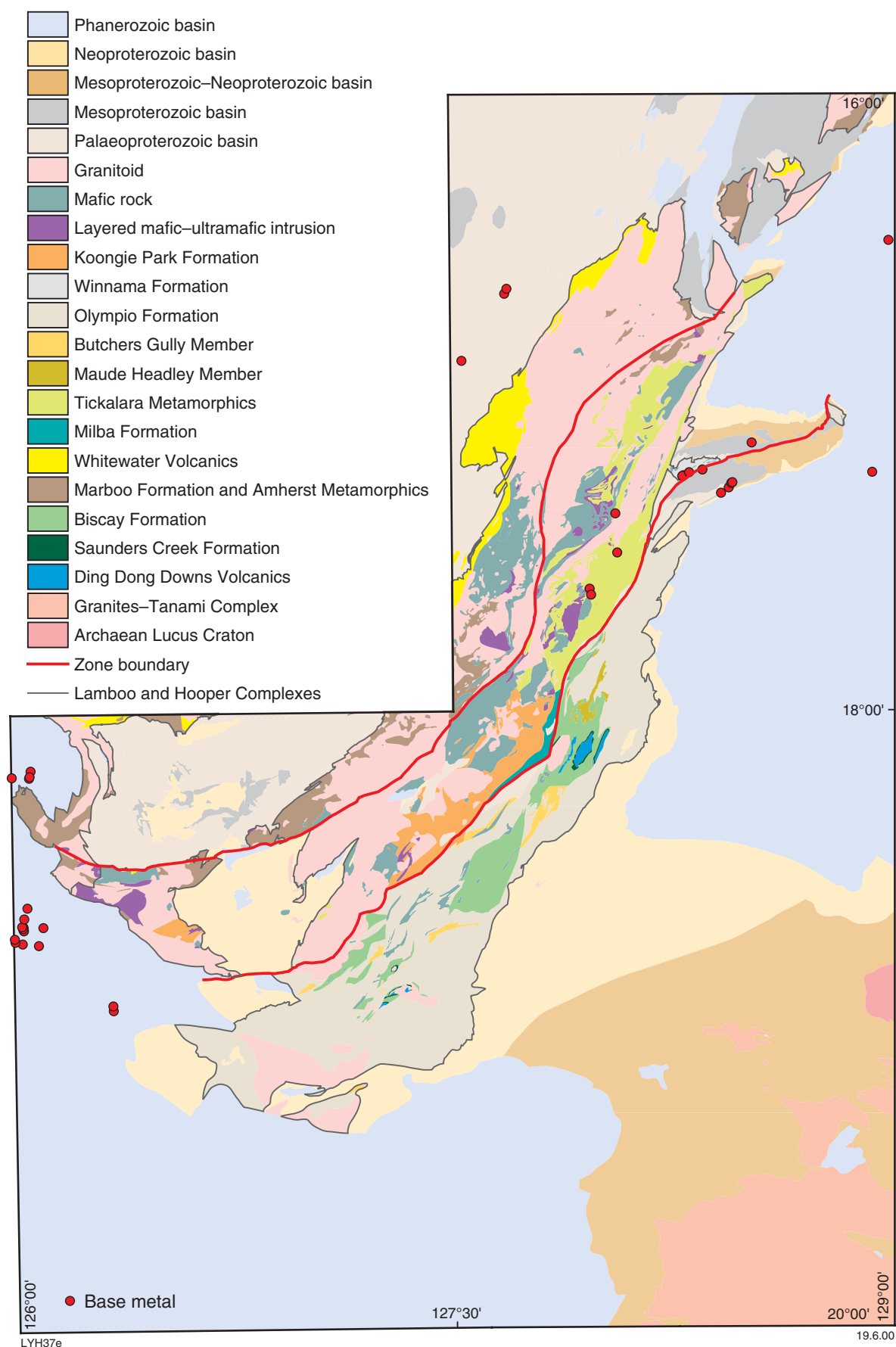


Figure 12. Tectonic sketch showing distribution of carbonate-hosted stratabound sedimentary deposits. Scale — 1 cm : 20 km (approximately)

like that at Cadjebut, is hosted by the Devonian Pillara Limestone in the Emmanuel Ranges (Fig. 13) close to a splay fault of the Cadjebut Fault (Ferguson, 1999). At the end of December 1999, the total measured, indicated, and inferred resources were 5.210 Mt at 8.0% Zn and 7.9% Pb and the total proved and probable reserves were 3.449 Mt at 6.9% Zn and 5.7% Pb (Western Metals Limited, 2000).

At Horse Spring Range 1 (5437), a copper–lead–zinc gossan is associated with black siltstone intercalated with fine-grained crystalline micrite and intraformational calcareous breccia of the Devonian Virgin Hills Formation near the Lindner Fault. Gossan samples contained up to 34% Zn, 6% Cu, 5% Pb, and 192 ppm Ag. The best drill intersection was 3 m (2 m true width) at 9.96% Zn and 0.49% Pb (Urquhart, 1983).

Mississippi Valley-type mineralization has also been reported from the Ross Hill area (e.g. Ross Hill 1, 5647; Findlay Hill 2, 5443; Teichart Hills, 5659; Pinnacles Lead, 6473; and Christmas Creek Reef 5, 5636).

Remobilized syngenetic or Mississippi Valley-type mineralization occurs in the Proterozoic Bungle Bungle Dolomite in the Osmond Basin. At Trend 1 (2518), rock-chip samples of gossan assayed up to 3.0% Pb; the best drill intersection was 12 m at 1.2% Pb (Trend Exploration Proprietary Limited, 1973). At Paradise Creek (Pb) (2507), two thin galena-bearing stromatolitic dolomite beds can be traced over a strike length of 2 km. Rock-chip samples from Paradise Creek (Pb) assayed up to 6.2% Pb (McIntyre, 1981); the best drill intersection was 1 m at

4.81% Pb (de Angelis, 1983). The galena occurs as blebs and veins within fractures and solution lines associated with soft-sediment deformation. Galena is commonly associated with pyrite and chert, and it also occurs in black carbonaceous laminae (de Angelis, 1983). Copper and zinc-bearing gossan is present in Bungle Bungle Dolomite at Trend 2 (5373).

In the Central zone of the Lamboo Complex, there are narrow zones of patchily developed copper(–tungsten–gold–tin) mineralization in calc-silicate rocks of the Tickalara Metamorphics at Alice Downs (1045). The best exploration drilling intercept was 2 m at 1.7% Cu (Chadwick, 1985). Gossans have also been reported from calcareous rocks of the Tickalara Metamorphics (e.g. Sally Downs C, 5030), where a gossan containing anomalous copper and molybdenum occurs in graphitic carbonate rocks. Minor copper-bearing gossans have been reported from limestone of the Winnama Formation (e.g. Osmond Valley (Cu 1), 5522).

In the Kimberley Basin, there is widespread sub-economic syngenetic copper mineralization in dolomite and shale in the uppermost unit of the Teronis member of the Elgee Sandstone (e.g. Teronis 3, 4510; Owen, 1970). At Teronis 1 (4489), there has been some secondary chalcocite enrichment (Hawley, 1970).

In the Ord Basin, there are small low-grade patches of secondary copper mineralization at Mount Elder (6483) in the uppermost limestone of the Cambrian Panton Formation of the Goose Hole Group (Dow and Gemuts, 1969).



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Figure 13. Photograph of the Devonian reef complex in the Emmanuel Ranges near the site of the Kapok mine

Undivided mineralization

This classification of stratabound sedimentary mineralization includes occurrences that are hosted by sedimentary rocks that are neither clastic nor carbonate (e.g. chert), occurrences hosted by a mixture of sedimentary types, and occurrences where the nature of the host rock is uncertain either because of metamorphism or because the host was not reported. A total of nine occurrences have been classified as 'stratabound sedimentary — undivided'. These include seven base metal occurrences and two steel-industry metal occurrences.

Base metal — copper, lead, zinc, silver

Minor base-metal bearing gossans occur in meta-sedimentary rocks of the Tickalara Metamorphics in the Central zone of the Lamboo Complex (e.g. Wills Creek A, 1059).

In the Ord Basin, at Discovery Yard (2506), there is a layer of banded chalcocite and malachite mineralization about 30 cm thick at the contact between limestone and sandstone in the Headleys Limestone, just above its contact with basalt of the Antrim Plateau Volcanics (Ball, 1966).

Steel-industry metal — vanadium, chromium, molybdenum

At Sally Malay Bore 5 (5597) and Sally Malay Bore 6 (5598), vanadium–chromium–molybdenum gossans are associated with BIF in the Tickalara Metamorphics in the Central zone of the Lamboo Complex.

Sedimentary mineralization — banded iron-formation (supergene enriched)

Iron

A total of seven occurrences have been classified as 'sedimentary — banded iron-formation (supergene enriched)'. The most important of these are high-grade supergene hematite and low-grade porous limonite iron ore deposits hosted by the Golden Gate Siltstone in the Carr Boyd Basin. These include the Pompeys Pillar (including Sam, 5344; Central, 5345; Tony, 5346; and Peter, 5349) and Matsu (5348) deposits (Harms, 1959; Brown, 1973). The hematite mineralization grades laterally and vertically into siliceous hematite and ferruginous shale; sedimentary features such as cross-bedding, ripple marks, and mudcracks indicate shallow-water deposition (Harms, 1959). There is an indicated resource at Sam of 8 Mt and Tony of 7 Mt at greater than or equal to 60% Fe, an inferred resource at Peter of 1.6 Mt and Central of 0.76 Mt at greater than or equal to 60% Fe, and an inferred resource at Matsu of 177 Mt at greater than or equal to 45% Fe (Brown, 1973).

In the Central zone of the Lamboo Complex, at Osmond Valley (Fe) (3295), BIF containing a large

amount of specular hematite is interbedded with chert within metasandstone of the Winnama Formation (Harder, 1983).

In the Granites–Tanami Complex, a small deposit of siliceous iron ore at Black Hill (5419) has been interpreted as a surficial concentration in metasedimentary rocks of the Killi Killi 'Beds' (Clarke and Blockley, 1961)

Vein and hydrothermal mineralization

Undivided mineralization

A total of 548 mineral occurrences in the east Kimberley area have been classified as 'vein and hydrothermal — undivided'. For convenience of description, these have been further subdivided on the basis of the main commodity present into precious metal (262 occurrences), base metal (222 occurrences), industrial mineral (23 occurrences), steel-industry metal (18 occurrences), energy mineral (17 occurrences), and speciality metal (six occurrences; Fig. 14); however, many of these occurrences are polymetallic so that the subdivision has often been partly arbitrary. The reader is referred to the list of commodities in each occurrence next to the WAMIN number in Appendix 1.

Precious metal — gold

In the Lamboo Complex, there is a marked concentration of vein and hydrothermal gold occurrences in the Eastern zone. Many of the occurrences define a northeasterly trend roughly parallel to the boundary with the Central zone, 7–12 km east of the boundary (Figs 14 and 15). Most of these (112 occurrences) are within the Biscay Formation. A significant number also occur in the Olympio Formation (52 occurrences), typically close to its contact with either the Biscay Formation or the Butchers Gully Member and within the Butchers Gully Member (21 occurrences). There are a few vein and hydrothermal gold occurrences in the Central zone, mainly in the Koongie Park Formation, Sally Downs supersuite, and Loadstone Monzogranite, and in the Western zone, mainly in the Amhurst Metamorphics. Significant vein and hydrothermal gold deposits are also present in the Granites–Tanami Complex. Examples from each area are briefly described below.

Lamboo Complex — Eastern zone

Brockman – Mount Bradley area

The regional setting of the Palm Springs openpit mine (2223; Fig. 16) together with that of the historical workings in the Brockman – Mount Bradley area including Faugh-A-Ballagh (2232), Golden Crown (2224), Lady Margaret (2233), Deimos (2235), Io (2234), Titan (4284), Europa (4288), and Mount Bradley (2236) is shown on Figure 17. Most of the mines in this area are hosted by trachyandesite of the Butchers Gully Member, which has been structurally repeated by high-amplitude isoclinal folding (D₅; Sanders, 1999). Production figures

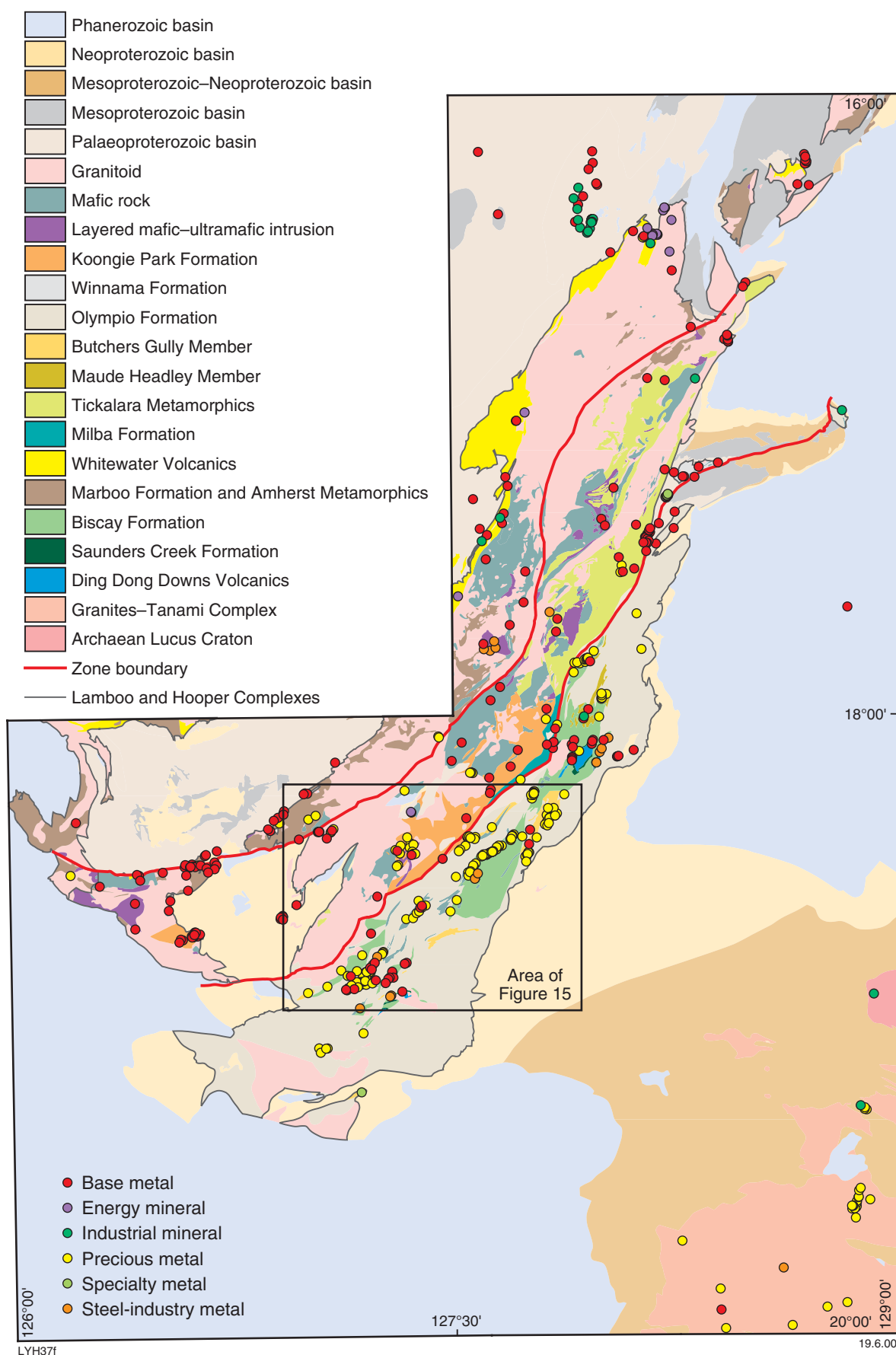


Figure 14. Tectonic sketch showing distribution of undivided vein and hydrothermal deposits. Scale — 1 cm : 20 km (approximately)

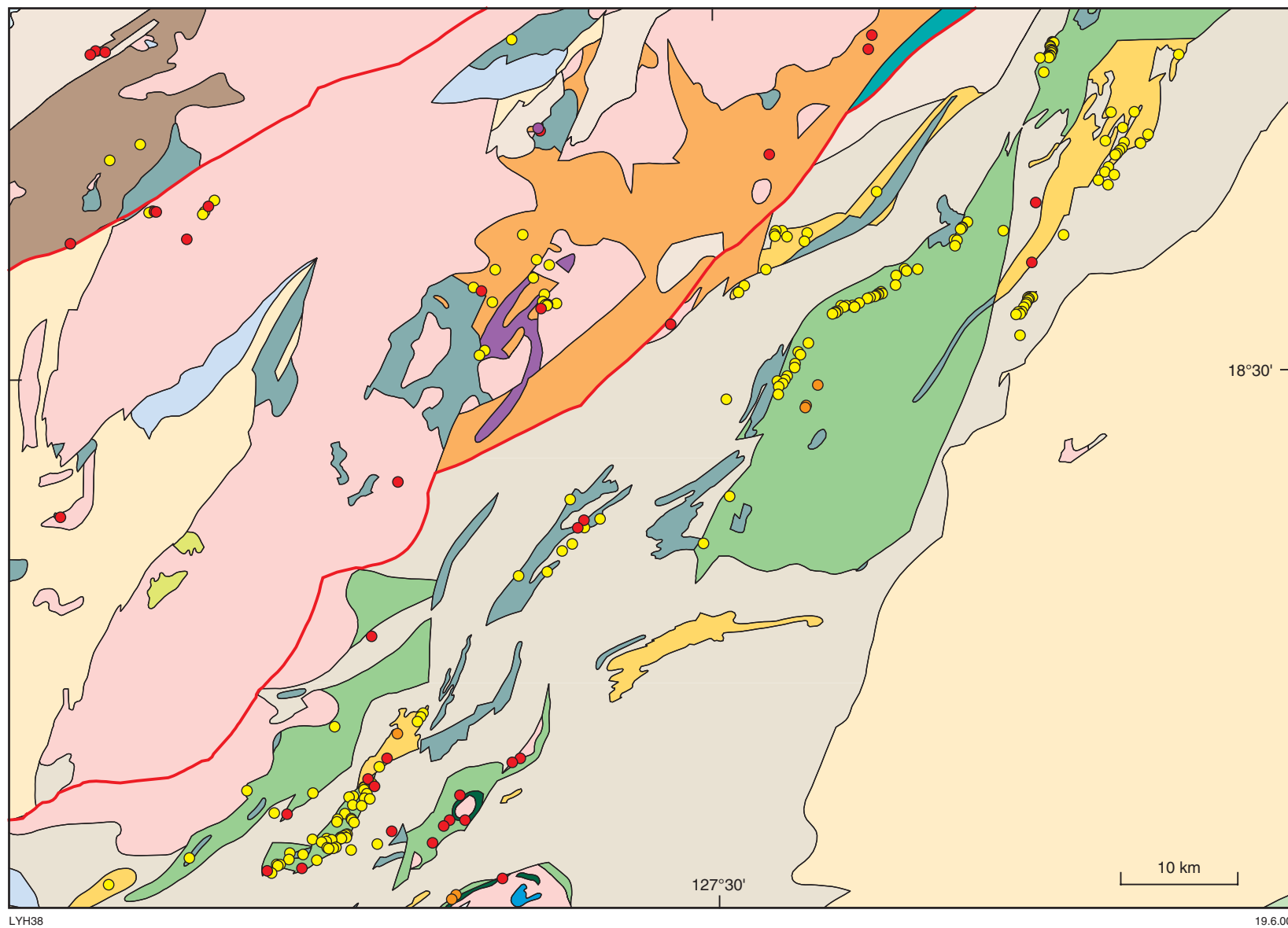


Figure 15. Enlargement of portion of Figure 14 showing the distribution of gold deposits in the southeastern part of the Lamboo Complex. See Figure 14 for key to symbols and colours



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Figure 16. Photograph of the Palm Springs mine

for the historical mines are given in Table 2. Between April 1995 and June 1997, the Palm Springs mine produced 2272 kg of gold (Precious Metals Australia Limited, 1997); this is more than double the total recorded pre-1964 production for the east Kimberley.

The quartz stockwork mineralization at Palm Springs is confined to a 30 m-wide fractured andesite of the Butchers Gully Member in a tightly folded south-southwest plunging anticline (Fig. 18). Numerous quartz veins cross-cut the anticline axis; these are probably tension veins related to the Ruby Queen Shear zone, which runs roughly parallel to the anticlinal axis (Fig. 18). Gold mineralization is associated with pyrite, minor arsenopyrite, and pyrrhotite in the quartz veins and in the altered wallrocks. Intense alteration of the andesite has produced an albite-ankerite-sericite-chlorite-quartz-pyrite assemblage (Sanders, 1999). The high sulfide content of the ore prevented mining of the ore below the zone of oxidation in the original Butchers Gully (Sams) mine in the 1930s (Barnes, 1985a).

Ruby Queen mining centre

Gold mineralization at the Ruby Queen mining centre is associated with a series of narrow, subvertical quartz veins, concordant with the bedding and foliation, in steeply dipping greywacke of the Olympio Formation (Dow and Gemuts, 1969). The veins are related to the north-northeast trending Ruby Queen Shear zone (Sanders, 1999). The Ruby Queen mine (2237) was the largest historical producer of gold in the east Kimberley. Prior to

1908, the Ruby Queen mine produced 9833 t at 19.66 g/t (193.3 kg), and between 1938 and 1964, it produced 3179 t at 17.1 g/t (54.3 kg). The quartz reef at Ruby Queen averaged 0.9 m in width and most of the gold was at the contact between the reef and the wallrocks (Smith, 1898). Other workings in the group include the St Lawrence (2238), West and Left (2239), Goliath (2240), Rising Sun (2243), and Sunny Corner (2244) mines. The mines are described in detail by Finucane (1939a) and Sanders (1999). Production figures for these mines are included in Table 2.

Kerry-Reform-Darcys

An arcuate group of workings lies within the Biscay Formation, and is aligned roughly parallel to its faulted contact with the Olympio Formation but 1–1.5 km east of the contact (Fig. 15). The official recorded production from the largest mine was 6.5 kg from the Reform mine (2259) in 1894, but company records (Eupene, 1983) suggest that production was considerably greater. The mineralization at the Reform mine is related to several shear-parallel quartz veins in a 1.5 m wide shear zone within siltstone and silicified shale (Sanders, 1999). A single stamp battery is said to have produced 300 oz (9.3 kg) from 300 t of ore from the Shane mine (2260; 2.4 km southwest of the Reform mine) between 1950 and 1961 (Eupene, 1983). Copies of mint returns were given in support of this claim but there are no localities shown on the mint returns. The gold in the Shane mine appears to have been in small rich shoots in metasedimentary rocks and tuffs of the Biscay Formation (Eupene, 1983).

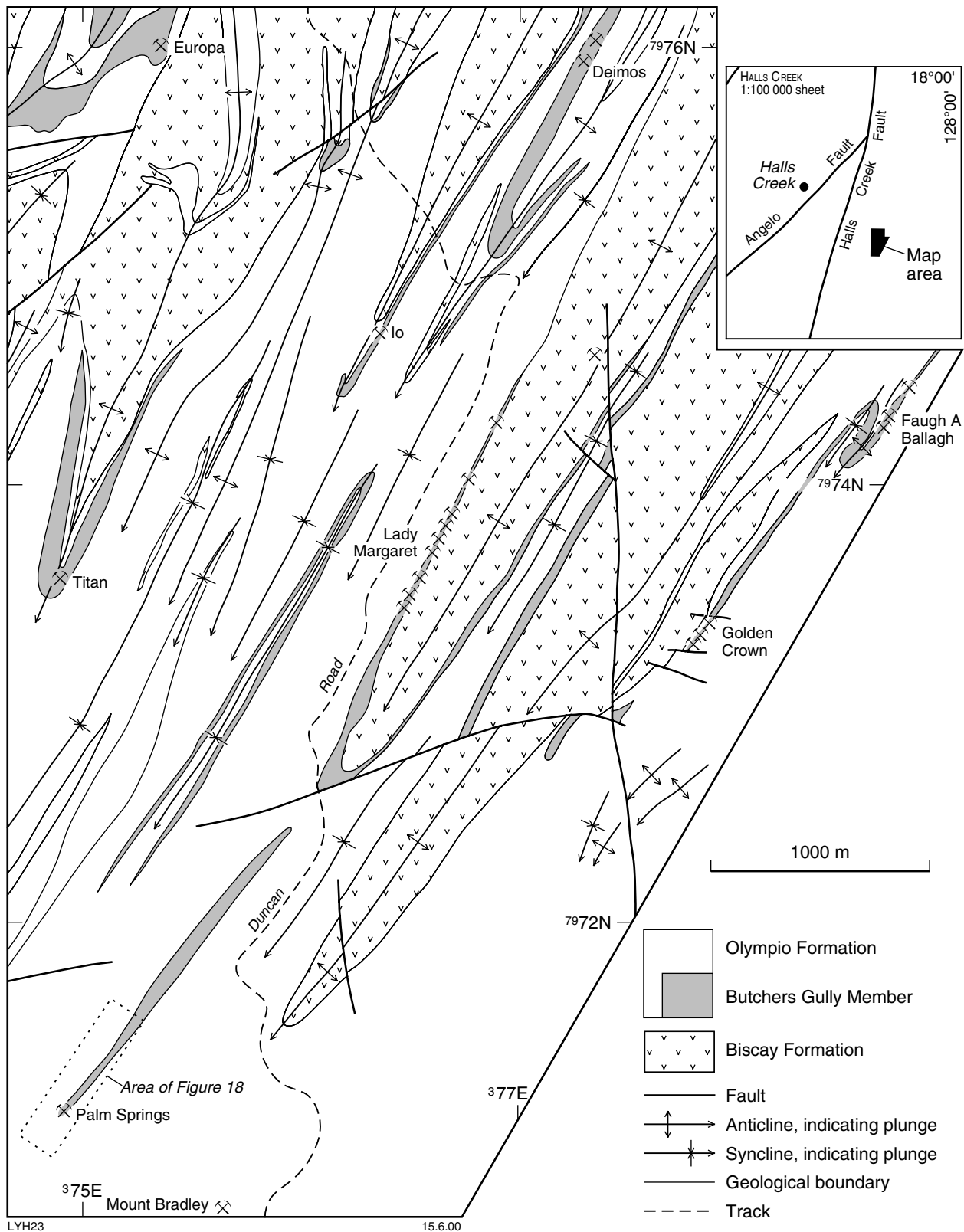


Figure 17. Brockman – Mount Bradley area — geology and location plan, 1995 (after Sanders, 1999). Differences between the geology shown on this map and that on HALLS CREEK (Blake et al., 1999a) may be due in part to differences in the scale and interpretation of the mapping

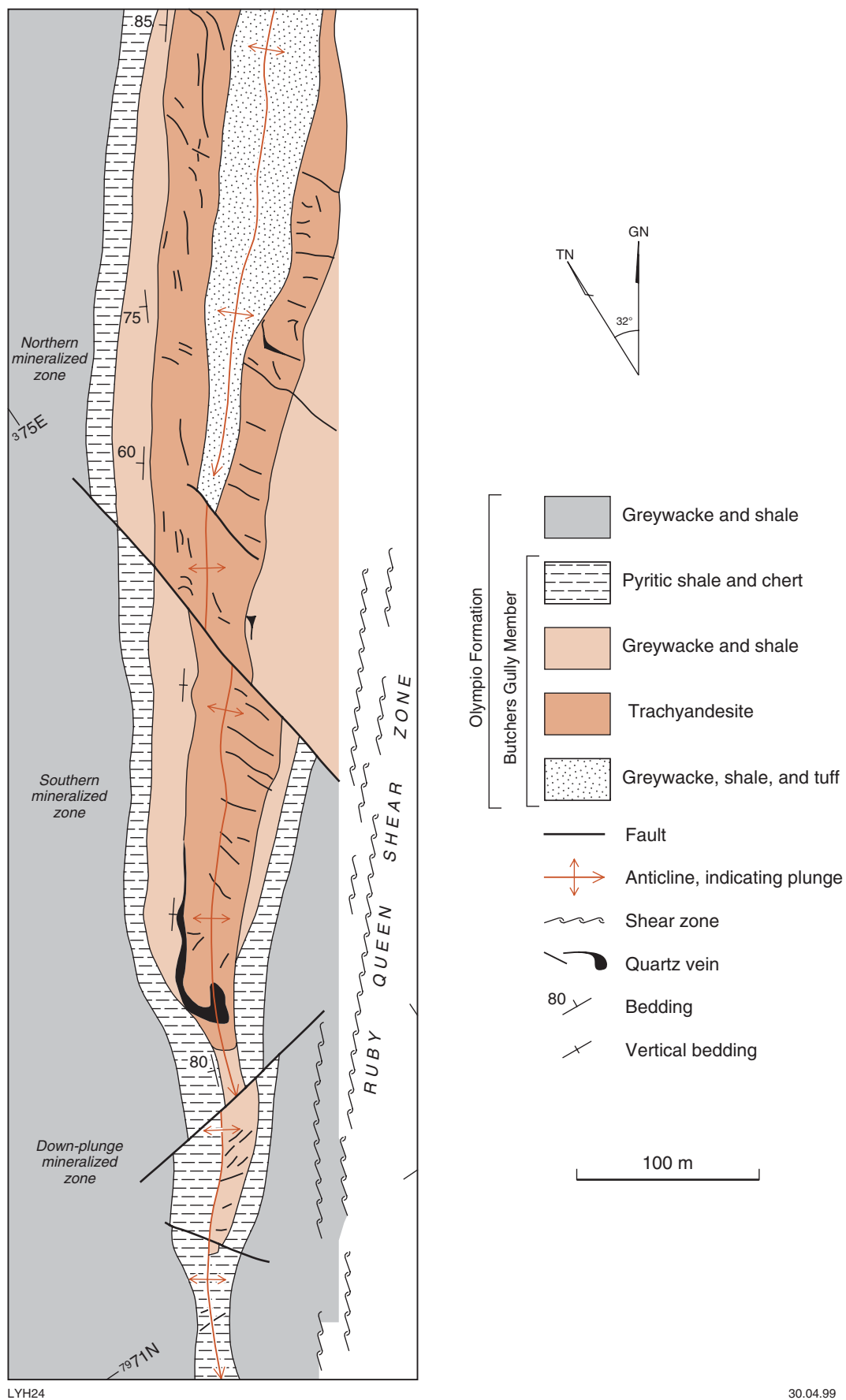


Figure 18. Simplified geology — Palm Springs mine area (after Sanders, 1999)

Table 2. Historical gold production from the east Kimberley (up to 1964)

<i>Mine name</i>	<i>Year</i>	<i>Alluvial and dollied gold (kg)</i>	<i>Ore treated (t)</i>	<i>Lode gold (kg)</i>
Mary River Goldfield				
Reform	Pre-1897	–	405	6.5
Newlook ^(a)	Pre-1897	32.1	–	–
PA99 ^(a)	1940–1941	–	48	1.7
PA151 ^(a)	1950	0.4	–	–
Total		32.5	453	8.2
Halls Creek Goldfield				
Jubilee	Pre-1897	–	23	14.0
Lady Broome	Pre-1897	–	406	0.9
Jacksons ^(b)	Pre-1891	–	78	1.6
Gladstone ^(b)	Pre-1891	–	2	0.1
Black Mount ^(b)	Pre-1891	–	8.5	0.1
unnamed ^(a)	Pre-1928	–	96	1.9
PA84 ^(a)	1938–1939	–	10	0.4
PA112 ^(a)	1941	–	9	0.5
PA118 ^(a)	1942	–	10	0.5
PA108 ^(a)	1942	0.9	–	–
PA115 ^(a)	1943	–	82	1.6
PA158 ^(a)	1953	–	9	0.5
PA161 ^(a)	1953	–	3	0.1
Total		0.9	736.5	22.2
Brockman – Mount Bradley				
Afgan	Pre-1897	–	96	2.1
Brockman King	Pre-1897	–	87	1.1
Faugh-A-Ballagh	Pre-1897	–	15	0.5
Golden Crown	Pre-1897	–	1 172	39.8
Southern Cross	Pre-1897	–	5	0.1
Mount Bradley ^(a)	Pre-1928	–	2 500	56.4
Mount Bradley	1940–1942	–	196	1.6
Lady Margaret ^(a)	Pre-1938	–	98.5	7.2
Lady Margaret No. 1 North ^(a)	Pre-1938	–	20.8	0.5
Lady Margaret No. 4 North ^(a)	Pre-1938	–	6.6	0.3
PA78 ^(a)	1938–1939	0.2	–	–
PA93 ^(a)	1939	–	–	0.1
PA97 ^(a)	1940	0.2	7	0.4
Total		0.4	4 203.9	110.1
Grants Creek area				
Lone Hand ^(c)	Pre-1898	–	–	8.1
Comet	Pre-1897	–	29	4.2
Star of Kimberley	Pre-1897	–	1	0.1
Perseverance ^(c)	Pre-1898	–	5.1	0.9
unnamed ^(a)	Pre-1923	–	3	0.5
PA96 ^(a)	1940	–	3	0.1
PA162 ^(a)	1955	0.2	–	–
Total		0.2	41.1	13.9
Mount Dockrell Goldfield				
Lady Hopetoun	Pre-1897	–	41	10.0
Victoria	Pre-1897	–	4	3.8
PA59 ^(a)	1935	–	37	0.8
unnamed ^(a)	1935	–	–	0.8
Mount Minaird ^(a)	1936	–	187	5.3
Western Lead	1935–1942	–	219	2.3
Irish Lass	1937–1943	0.7	346	8.3
Old Mac	1938–1940	–	239	5.6
Erin-Go-Bragh	1939–1941	–	62	1.7
PA75 ^(a)	1938	0.6	–	–
PA102 ^(a)	1940	–	26	0.4
PA101 ^(a)	1940–1941	–	83	0.7
unnamed ^(a)	1941	–	16	0.1
Old Golden Dream	1941–1942	–	93	0.8
Sundry Claims ^(a)	–	–	163	2.8
Total		1.3	1 516	43.4

Table 2. (continued)

<i>Mine name</i>	<i>Year</i>	<i>Alluvial and dollied gold (kg)</i>	<i>Ore treated (t)</i>	<i>Lode gold (kg)</i>
Ruby Creek Goldfield				
No. 1 North Ruby Queen	Pre-1897	—	194	5.8
Rising Sun	Pre-1897	—	586	15.3
Sunny Corner	Pre-1897	—	22	1.8
West and Left	Pre-1897	—	690	29.5
West and Left	1938–1941	—	10	0.35
St Lawrence	Pre-1902	—	1 510	47.7
St Lawrence	1938	—	10	0.2
Ruby Queen (claim 61)	Pre-1908	—	9 833	193.3
Ruby Queen (claim 97)	1938–1964	0.1	3 179	54.3
unnamed ^(a)	Pre-1928	—	153	3.9
PA79 ^(a)	1938	—	93	0.7
Goliath ^(a)	1938–1942	—	123	3.2
Goliath ^(a)	1940	—	7	0.2
PA91 ^(a)	1939	—	120	nk
PA93 ^(a)	1939	0.4	—	—
PA96 ^(a)	1939	—	12	0.1
unnamed ^(a)	1939	0.1	—	—
Shorts Hope	1940	—	51	0.1
Darcys mine	1941–1942	0.5	79	1.9
PA124 ^(a)	1947	0.4	8	0.2
Total		1.5	16 680	358.6

SOURCE: Geological Survey of Western Australia gold production statistics database (GOLDSTATS) except where indicated

NOTES: (a) Dow and Gemuts (1969)
(b) Woodward (1891)
(c) Smith (1898)
nk not known

Production from the Kerry mine (2262), at the southern end of belt, in the same period is reputed to be about 2000 oz (62.2 kg; Eupene, 1983), although no official production is recorded. The Darcys (2253) workings, at the northern end of the trend, are on quartz veins along a shear zone. Recorded production from Darcys is 1.905 kg plus 0.499 kg of dollied gold. Mullumby (1986) reported that selected dump samples contained up to 28.8% Pb, 1.14 ppm Ag, and 1.42% W.

Old Halls Creek area

The Old Halls Creek area is where the first gold rush occurred in Western Australia, and the first gold mines opened here. Mines included Jubilee (2246), Lady Broome (2250), Jacksons Reef (2251), Black Mount (6131), and Gladstone (2249). Most gold from this area probably went unreported (Playford and Ruddock, 1985); published production figures are included in Table 2. The mines are hosted by the Biscay Formation. At the Jubilee mine, gold was mined from cross-cutting quartz veins that contained abundant argentiferous galena and copper oxide (Woodward, 1891; Smith, 1898). At Jacksons Reef, gold mineralization was associated with a quartz reef and quartz stringers parallel to the schistosity.

Mount Dockrell area

Most of the workings in the Mount Dockrell area are on quartz veins along shears in the Biscay Formation or Butchers Gully Member. Many of the quartz veins contain

significant galena and high silver values in addition to gold (Woodward, 1891). At McNeils (2321), rich specimen gold was obtained at the intersection of shears. Other significant producers were the Victoria (2318), Irish Lass (2324), and Old Mac (2332) mines (see Table 2; note that Lady Hopetoun (2322) was an amalgamation of the Victoria, McNeils, Hard Labour, and Black Prince leases; Woodward, 1891).

Grants Creek area

Most gold in the Grants Creek area was mined from quartz veins in the Biscay Formation. The most significant producers were the Lone Hand (2221), Comet (2218), Perseverance (2212), and Star of Kimberley (2211) mines (see Table 2). Galena and silver was associated with the gold mineralization (Smith, 1898).

Christmas Creek

At Christmas Creek 1 (2300) there is an auriferous quartz–sulfide stockwork in sheared dolerite; the dolerite here differs from the Woodward Dolerite in that it is highly magnetic (Barnes, 1985b).

Lambo Complex — Central zone

Nicolsons Find

At Nicolsons Find (2340), mineralized quartz veins within felsic volcanic rocks of the Koongie Park Formation are related to the Nicolson Fault; anomalous As and Pb values

are associated with the gold mineralization (Sanders, 1999). Trial mining of the deposit was undertaken by Precious Metals Australia in 1996; 15 000 t at 5.08 g/t (76.2 kg) was treated at the Palm Springs plant (Precious Metals Australia Limited, 1996). The inferred resource at June 1997 was 278 000 t at 5.54 g/t (Precious Metals Australia Limited, 1997).

Lady Helen

At Lady Helen (also known as Moola Bulla, 2339), there are shallow workings in gossanous vuggy quartz veins in a shear in mafic rock of the Sally Downs supersuite close to the contact with the Loadstone Granite. Silver and arsenic are associated with the gold mineralization (Mullumby, 1989; Sanders, 1999). On the basis of panning, Mullumby (1989) estimated that the maximum resource was about 1000 t at 10 g/t, but note that this is an order of magnitude estimate only and is not a formal resource under the JORC code.

Lamboo Complex — Western zone

Minor gold mineralization is associated with silver and base metal mineralization in quartz veins within the Amherst Metamorphics at Me No Savvy C (1275). At the Jailhouse Creek prospect (2345 and 2347), low-grade gold and silver are associated with base metals in quartz veins in felsic lithic crystal tuff of the Whitewater Volcanics. Hunter Resources Limited (1991) interpreted the veins in the area to be epithermal because some have bladed

textures. There are no official records of gold production from the Western zone.

Granites–Tanami Complex

Gold was discovered in the Granites–Tanami Complex at the Tanami mine in the Northern Territory in 1900. Additional gold mineralization was discovered at Dead Bullock Soak and Callie in the Northern Territory in 1988 and 1991 respectively. The Granites–Tanami Complex in Western Australia was largely neglected until the discovery of the Kookaburra deposit in 1995 followed by the discovery of Coyote in 1999, with its spectacular, high-grade intercepts. The area of the Granites–Tanami Complex is now almost completely under tenure.

Larranganni area

Quartz samples containing anomalous gold were first reported from the Larranganni area by Talbot (1910), but it was not until 1983 that any significant effort was made to follow up Talbot's discovery (Doust, 1997). Glengarry Resources NL has reported significant gold mineralization over a 15 km strike length within folded mafic volcanic and sedimentary rocks that have been informally named the Bald Hill sequence (Glengarry Resources, 1999, written comm.; Quinn, 1999; Fig. 19). Carbonate, sericite, quartz, iron, and chlorite alteration is associated with the gold mineralization. Some gold mineralization occurs 800 m to the west in a similar sequence informally named the Quartz Ridge sequence. These sequences have been

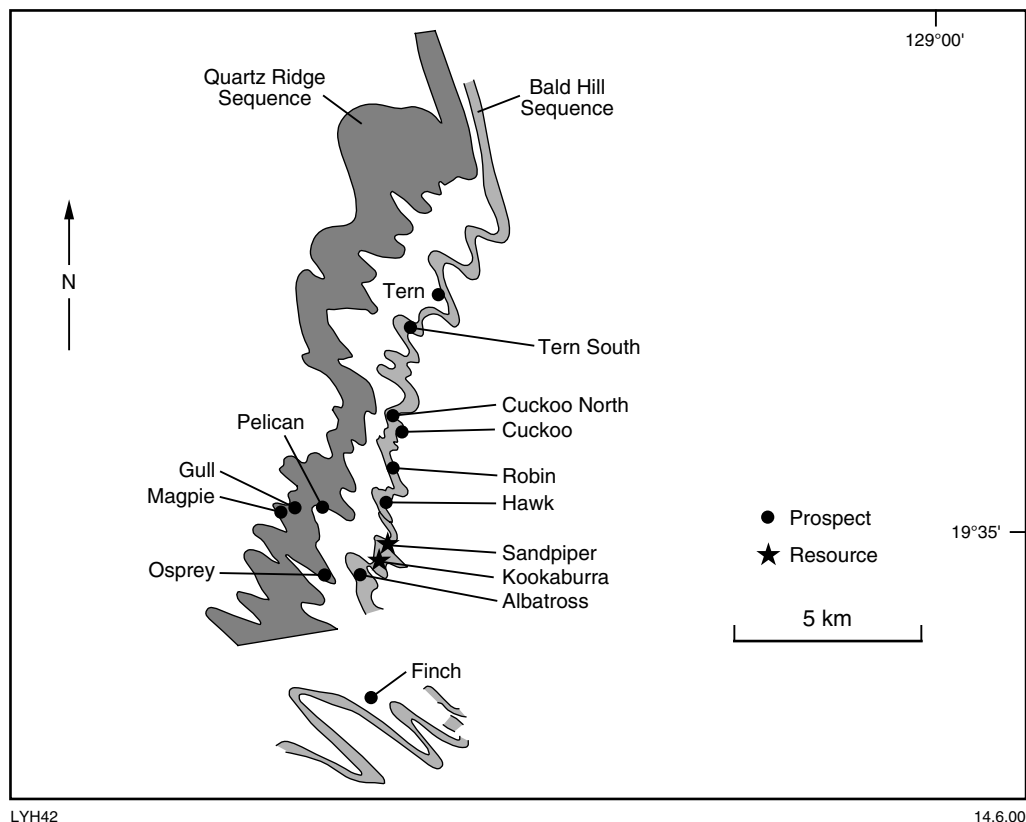


Figure 19. Distribution of gold deposits in the Larranganni area (after Quinn, 1999)

correlated with the Mount Charles beds in the Northern Territory (Glengarry Resources NL, 1999, written comm.). The Kookaburra deposit (6074) is situated in the Bald Hill sequence to the south of a west-northwesterly trending fault (Fig. 19) and contains multiple quartz veins in a brecciated ferruginous unit (Doust, 1997). The Kookaburra deposit is estimated to have an indicated resource of 1.5 Mt at 2.1 g/t (100 000 oz; 3110 kg) and an inferred resource of 0.2 Mt at 2.1 g/t (15 000 oz; 467 kg) using a cutoff grade of 0.5 g/t (Glengarry Resources NL, 1996). The Sandpiper deposit (6075), which lies just to the north of the west-northwesterly trending fault, contains an additional indicated and inferred resource of 0.58 Mt at 3.22 g/t (60 000 oz; 1866 kg; Glengarry Resources NL, 1998). Other prospects spread over a 15 km strike length (including Hawk, Finch, Osprey, Albatross, and Cuckoo North) have given significant drill intersections (Glengarry Resources NL, 1999a).

Coyote

Spectacular high-grade gold mineralization has been intersected over an area of 800 × 200 m at the Coyote prospect (6466). Better intersections include 8 m at 458 g/t gold from a depth of 105 m, and 13 m at 39 g/t gold from a depth of 174 m. Oxidation extends to a depth of greater than 100 m. Free gold is commonly visible but sulfides are rare, even in fresh rock (Acacia Resources Limited, 1999). The mineralization is said to be hosted by nonmagnetic sedimentary rocks of the Killi Killi Formation, informally referred to as the Madigan beds (Tanami Gold NL, 1999).

Other areas

At the Camel prospect (6467), there is a sheeted vein system that is axial planar to a major northwesterly plunging fold in metadolerite. Veins also occur along the contact between metadolerite and metasedimentary rocks. The associated gold–arsenic anomaly extends for 5 km (Glengarry Resources NL, 1999b).

Gold occurs in narrow, steeply dipping, northerly trending anastomosing quartz veins (associated with shear zones) and in minor saddle reefs and shallow-dipping quartz veins at the Venus prospect (5809, 5810, and 5811). The veins are in altered and deformed Killi Killi 'Beds' close to its unconformity with the Gardiner Sandstone (Stocklemayer and Rothnie, 1987). Gold is also associated with uranium at the Don prospect (5793), 2.8 km west-northwest of the Venus prospect (see **Unconformity-related mineralization**, p. 48).

Base metal — copper, lead, zinc, silver

The distribution of base metal vein and hydrothermal occurrences is shown on Figure 14. In addition to the occurrences classified as base metal occurrences, many of the gold occurrences also contain significant base metals. Base metal vein and hydrothermal occurrences fall within a wide range of stratigraphic units in all three zones of the Lamboo Complex and in the overlying younger basins. Only one occurrence has been recorded from the

Granites–Tanami Complex. A few examples from different stratigraphic settings are briefly described below.

Lamboo Complex — Eastern zone

At Horrys Copper Show (2330), in the Mount Dockrell area, Cu–Au–Ag mineralization is hosted by a gossanous quartz vein within a shear zone at the contact between the Olympio and Biscay Formations. The high-grade zone is about 2.4 m wide and extends 60 m along the shear zone, but a zone of sporadic mineralization is about 7.4 m wide and extends 900 m along the shear zone. At the Hanging Drum Copper Show (2331), which is also near Mount Dockrell, there is abundant malachite in a quartz-filled shear in black shale of the Olympio Formation (Barber, 1971; Sanders, 1999). Small high-grade lenses of copper mineralization are associated with quartz and sheared greywacke of the Olympio Formation on a narrow, near-vertical, northeasterly trending fault at Mount Kinahan (1211, 5078, and 5083). Minor copper mineralization occurs in quartz veins in the Ding Dong Downs Volcanics at Saunders Creek Anticline (5100, 5101, and 5102) and in the Biscay Formation at Sophie Downs (Cu) 1 (4822). At Ding Dong Downs (Pb–F) (1104), there is coarse-grained galena and purple fluorite in a breccia zone within the Maude Headley Member. The breccia has been interpreted as a breccia pipe by Fletcher (1981), but Johnson (1981) has interpreted it as tectonic breccia.

Lamboo Complex — Central zone

At the Calico Spring prospect (e.g. 1222, 5541, 5542), gossanous and malachite-bearing quartz veins, up to 300 m long, occur over a strike length of 3 km. The veins are aligned parallel to the stratigraphy and a north-northeasterly trending splay off the Halls Creek Fault, and they are hosted by the Tickalara Metamorphics and the McHale Granodiorite (Australian Anglo American Limited, 1973; Sanders, 1999). The Eastman Bore intrusion hosts possible remobilized primary copper mineralization in shears (e.g. Eastman Bore (Cu 3), 6212). At Bullock Bore 3 (4299) and Bullock Bore 4 (4301), in the same general area, copper mineralization is associated with shear zones in tuffs of the Koongie Park Formation. Koongie Park 2 (5725) and Ten Minute (Cu 2) (2530) are other examples of mineralization hosted by the Koongie Park Formation. Possible epithermal lead–barite mineralization occurs in granite of the Sally Downs supersuite at Boxers (1221) in Lake Argyle. The mineralized vein contains brecciated country rock suggesting it is a filling of a minor fault (Harms, 1959). The vein trends 020°, which is parallel to the Carr Boyd Fault.

Lamboo Complex — Western zone

The only recorded production of base metals from the 'vein and hydrothermal — undivided' style of deposit in the east Kimberley area is from Black and Glidden (1257; 6.24 t of lead and 1045.7 g of silver from 9.41 t of ore; Blockley, 1971). At Black and Glidden, mineralized quartz veins are hosted by the Amherst Metamorphics and are related to the Lubbock Range Fault. Sporadic mineralization occurs over a strike length of 2200 m and individual

veins are up to 20 m long and 3 m wide. Widespread quartz vein mineralization at Me No Savvy (1272, 1274, and 1277) is also hosted by the Amherst Metamorphics and related to the Glidden Fault. Copper mineralization in the Whitewater Volcanics is associated with shear zones at the Base Metal Prospect (1340) and Mount Remarkable 2 (1334), and with epithermal quartz veins at Jailhouse Creek 2 (2346).

Speewah Basin

Lead–copper–silver(–gold–antimony) quartz vein mineralization in the Speewah Basin is hosted by the granophyric phase of the Hart Dolerite (termed the Yilingbun Granophyre by Alvin, 1993). Examples include Martins (1500), Calamondah (1501), Speewah IV (1503), Speewah V (1504), and Speewah VI (1505). Some base metals are also associated with epithermal fluorite veins (e.g. Martins (fluorite), 2428).

Kimberley Basin

At Campbellmerry (2966), brecciated copper mineralization occurs within narrow shears parallel to major northwesterly trending faults in the Pentecost Sandstone (Harms, 1959).

Bastion Basin

At Plants (6488), there are minor veins of copper carbonates, oxides, and chalcocite in siltstone of the Mendena Formation (Harms, 1959).

Carr Boyd Basin

Disseminated galena and lead oxides occur in a silicified breccia zone within sandstone of the Golden Gate Siltstone at the Carr Boyd prospects (e.g. 2520; Pickands Mather and Company International, 1966)

Ord Basin

Minor veins with malachite in the Cambrian Headleys Limestone are exposed in a road cutting at Brook Creek (5250).

Granites–Tanami Complex

Copper mineralization is associated with ferruginous quartz along an east-southeasterly trending fault in the Killi Killi ‘Beds’ at Errol Bore (5720; Pratt, 1979).

Industrial mineral — fluorite

The Speewah Basin hosts Western Australia’s largest fluorite deposit (Fig. 20). In the Main zone, there are at least nine fluorite vein sets over a strike length of 7.5 km that are related to a north-northeasterly trending splay fault of the Greenvale Fault. The largest of these is B Vein (5352), which is 700 m long and up to 10 m wide. The total published resource for A Vein (2397), B Vein (5352), and C Vein (5354) is 1.89 Mt measured, 0.41 Mt indicated, and 1.59 Mt inferred at 25% CaF₂ using a 13% CaF₂ cutoff grade (Rogers, 1998). Additional veins outcrop over an area of 30 × 8 km. The veins are hosted by granophyre



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Figure 20. Photograph of the A vein at Speewah

and dolerite of the Hart Dolerite, arkose of the Lansdowne Arkose, siltstone and mudstone of the Valentine Siltstone, and a tectonic and hydrothermal breccia informally termed the Doon Doon breccia (Rogers, 1998). The veins have epithermal textures, and fluid-inclusion studies indicate that they have trapping temperatures of about 150 to 170°C (Rogers, 1998). Samarium–neodymium dating of fluorite in the veins gave an age of 120 Ma (Rogers, 1998). A possible carbonatite dyke is associated with fluorite mineralization at West Ridge (2399; Alvin, 1993; Rogers, 1998; see **Mineralization in carbonatite and alkaline igneous intrusions**, p. 22).

Minor fluorite is present at the Frog prospect in the Whitewater Volcanics in the Western zone of the Lamboo Complex. It occurs both as separate veins (e.g. Frog 8; 5334) and in breccia zones in association with uranium mineralization (e.g. Frog 4; 2370).

Purple fluorite is also associated with brecciated cross-cutting quartz veins in felsic volcanic rocks of the Maude Headley Member at Ding Dong Downs (F) (5379) in the Eastern zone of the Lamboo Complex.

Industrial mineral — barite

Barite is associated with some of the fluorite veins in the Speewah Basin and some separate barite veins also occur in this area (e.g. Main Shear North, 2423). Boulders of coarsely crystalline barite have also weathered from a vein in the Mabel Downs Tonalite at Mabel Downs (Ba) (6494) in the Central zone of the Lamboo Complex.

Steel-industry metal — tungsten

All known tungsten vein deposits in the study area are in the Eastern zone of the Lamboo Complex. Scheelite occurs in quartz veins in Biscay Formation at Ruby Plains (5900, 5901, 5902) and in the Woodward Dolerite at Balara (959). Tungsten mineralization is also in quartz veins in the Berthas Butt (e.g. Castle Creek 9, 5044), Taylor Lookout (e.g. Taylor Lookout 4, 945), and Mount Dockrell (North Dockrell, 882) areas, but at these places it is associated with stratabound scheelite mineralization (see **Stratabound sedimentary mineralization**, p. 31) and has probably been remobilized from this style of mineralization.

Steel-industry metal — nickel

A gossan with anomalous nickel and copper occurs on the faulted contact between the Springvale intrusion and the Airfield Granodiorite at Springvale 8 (5187) and Springvale 10 (5189). Gossan with anomalous nickel and copper also occurs on shears within the Springvale intrusion (e.g. Springvale 9, 5188). Pickands Mather and Company International (1968) considered this mineralization to be remobilized primary sulfides.

Steel-industry metal — molybdenum

Pickands Mather and Company International (1967a) reported rumours of a molybdenite occurrence associated

with calc-silicate rocks at Mabel Hill (5169); they found anomalous Mo in creeks draining Mabel Hill but did not follow it up.

Energy mineral — uranium

All but one of the 17 uranium occurrences (classified as ‘vein and hydrothermal — undivided’) are in the Western zone of the Lamboo Complex. At Antares (2380), secondary uranium mineralization occurs in fractures and joints in felsic tuff in the Whitewater Volcanics; there are several mineralized zones up to 20 m long and 10 m wide. At the Frog prospect, secondary uranium mineralization also occurs within the Whitewater Volcanics on joints and shears (e.g. Frog 1, 2348) and in breccia zones associated with fluorite (e.g. Frog 4, 2370). At Dunham River B (2378), torbernite occurs in a brecciated dolerite dyke along the Dunham Fault zone and in mineralized quartz veins in sheared Crooked Creek Granite adjacent to the dyke (Fraser, 1955; Dale, 1976). At Donkey Creek (4804), secondary uranium mineralization occurs in siltstone of the Golden Gate Siltstone of the Carr Boyd Basin in the Dunham Fault zone (Dale, 1976).

The only other known uranium occurrence is John Galt (U) (5350), which is in the Red Rock Formation of the Red Rock Basin. At this locality, there are specks of carnotite in silicified sandstone boulders; boulders of mylonitic breccia also occur with autunite and torbernite on cleavage planes.

Speciality metal — REE

At the John Galt (REE 1) prospect (1105), xenotime occurs as veinlets or segregations up to 30 cm thick in quartz veins, which follow joints and small faults in sandstone of the Red Rock Formation of the Red Rock Basin. Some xenotime is also present as impregnations in the sandstone. The mineralization is exposed on the faulted east wall of a mesa. Mascot Mining estimated an indicated resource of 51 250 t at 0.41% total REE (pre-JORC code; Campana, 1973; Reisgys, 1987). Additional xenotime-bearing veins are present at John Galt (REE 2) (5318) and John Galt (REE 3) (5319).

Xenotime has also been reported from the Lucas Craton and the Granites–Tanami Complex. At Browns Range Dome (6260), there is a vein of xenotime, 10–30 cm thick and 15 m long, and also a series of smaller xenotime veins associated with vuggy quartz and hematite alteration in the Archaean Browns Range Metamorphics of the Lucas Craton. At Whites Beach (6141), minor xenotime, of possible hydrothermal origin, occurs in kaolinitic shale of the Killi Killi ‘Beds’ in the Granites–Tanami Complex.

Speciality metal — bismuth

At Shiddi Creek (5673), anomalous bismuth and gold mineralization are present in quartz–plagioclase–sericite schist and fine-grained dolomitic carbonate rock that has been interpreted by Fraser and Doe (1975) as Biscay Formation, but which could be Butchers Gully Member (Tyler, I. M., 1999, pers. comm.).

Unconformity-related mineralization

Energy mineral — uranium

There are six unconformity-related uranium occurrences within the study area, and all are in the Gardiner Range area at, or close to, the unconformable contact between the Killi Killi 'Beds' of the Granites–Tanami Complex and the overlying Gardiner Sandstone of the Birrindudu Basin. At the Don prospect (5793), uraninite and secondary uranium minerals (and gold) occur in siltstones and carbonaceous shales of the Killi Killi 'Beds'. They occur along an east-northeasterly trending fault, close to the unconformity between the Killi Killi 'Beds' and Gardiner Sandstone. The best intersection obtained from exploration drilling was 0.11 m at 4.65% U and 2.0 ppm Au in BIR1 (Laporte, 1981). At Mount Mansbridge 1 (5805), there is a zone of kaolinized black shale with pyritic quartz veins in the Killi Killi 'Beds' marked by high background spectrometer counts and a strong biochemical anomaly. Samples from pits assayed up to 980 ppm U and 0.16 ppm Au; xenotime is also present in the pits (Sutherland, 1983).

Regolith mineralization

Alluvial to beach placer mineralization

A total of 162 occurrences have been classified as 'regolith — alluvial to beach placers'. These have been further subdivided according to their major commodity into precious metal (107 occurrences), precious mineral (32 occurrences), speciality metal (13 occurrences), and steel-industry metal (10 occurrences).

Precious metal — gold

Most of the alluvial gold occurrences in the study area are in the Eastern zone of the Lamboo Complex. They are concentrated in creeks and rivers draining primary vein and hydrothermal deposits. Alluvial workings in the Old Halls Creek area (e.g. Nuggety Gully 1, 2730) probably produced significant amounts of gold in the early years, but most of the early production is unrecorded. More recently, alluvial gold was mined from the Mary River area (e.g. Mary River 1, 6169). Between 1988 and 1990, 72.02 kg of gold was produced from the Mary River and its tributaries (Roebuck Resources NL, 1990).

Alluvial gold has been reported from the Gardiner Range area near the boundary of the Granites–Tanami Complex and the Birrindudu Basin at Lewis Creek (5826; De Havelland, 1986) and Slaty Creek (5417; Clarke and Blockley, 1961). Harms (1959) also reports that alluvial gold has been found in Prospect Creek (6486), which drains the Carr Boyd Basin.

Precious mineral — diamond

Significant quantities of gem-quality alluvial diamonds, derived from the Argyle AK1 pipe, are being mined from Upper Smoke Creek (1108), Lower Smoke Creek (1109), Limestone Creek (1107), and Bow River (alluvials) (1113). The Argyle alluvials (Smoke Creek and Limestone Creek)

had a total measured, indicated, and inferred resource of 45 Mt at 0.2 ct/t and a reserve of 5.8 Mt at 0.24 ct/t at 31 December 1999 (Ashton Mining Limited, 2000).

At Barramundi Range 2 (6387), gem-quality diamonds are said to have been recovered from the bank of Middle Creek, which may be a palaeodrainage channel of the Leopold River (Davies, 1990).

Small diamonds or microdiamonds have also been found in the vicinity of the Durack Range (e.g. Lightning Creek, 5376), near Gordon Downs Homestead (e.g. Gordon Downs 1, 6028), and in the northwest corner of MOUNT BANNERMAN at Salt Creek (6420) and Wattle Spring Creek (6422).

Precious mineral — topaz

Abundant clear topaz (up to 55% of heavy mineral concentrate) was found in a stream-sediment concentrate from Argyle Dykes 3 (3004) during diamond exploration (Brown, 1992). It is not known if there is any topaz of significant size in the area. The topaz was possibly derived from a lamproite dyke, but the source was not traced.

Speciality metals — tin, tantalum, niobium (gold, tungsten)

In the southern part of the Halls Creek Orogen, alluvial tin has been reported from Mount Cross 3 (6379). Alluvial tin, tantalum, niobium, and tungsten occur at Junda 2 (965) and Junda 4 (5915); alluvial tin and gold occur at Minneroo Pool 3 (5822). Alluvial tin has also been recorded in the northern part of the east Kimberley area (e.g. Neil Creek 1, 966).

Speciality metal — heavy minerals, REE

Xenotime and monazite have been recorded from recent alluvium and raised cobble wash at Corkwood Yard (3005).

Steel-industry metal — tungsten, titanium, vanadium

Alluvial scheelite occurrences include Foal Creek (968) in the northern part of the Halls Creek Orogen and Hangmans (W) (5909) in the southern part of the Orogen. Alluvial vanadiferous ilmenite has been reported from Speewah Valley (Ti) (844).

Residual to eluvial placer mineralization

A total of 37 occurrences have been classified as 'regolith — residual to eluvial placers'. These include 33 precious metal occurrences, two base metal occurrences, one steel-industry metal occurrence, and one speciality metal occurrence.

Precious metal — gold

Eluvial gold occurs in close association with primary gold. Examples include Top of Bucket (6185) near Mount

Dockrell, Comet Hill (3213) in the Grants Patch area, and Christmas Creek 6 (6071) at the southern end of the Halls Creek Orogen. Early prospectors worked the eluvial gold by dry-blowing, but more recently, metal detectors have been used to locate eluvial gold. Unfortunately, some of the recent operators have used bulldozers (apparently without approval from DME) to remove surface soil, and this has caused serious environmental degradation.

Base metal — copper

At Copper Flats (2756) in the Ord Basin, there are abundant pebble-sized fragments of secondary copper minerals and copper-stained shale in skeletal soil over an area 1000 × 600 m (Matheson, 1996). Several agate boulders containing abundant chalcopyrite were reported from Eastman (Cu) A (790) near the southern contact of the Eastman Bore intrusion (Lewis, 1968).

Steel-industry metal — tungsten

There is a colluvial scheelite concentration in surface rubble at Sophie Downs North (933).

Residual and supergene mineralization

A total of eight occurrences have been classified as 'regolith — residual and supergene'. These include four steel-industry metal occurrences, two energy mineral occurrences, one base metal occurrence, and one industrial mineral occurrence.

Steel-industry metal — nickel

Lateritic nickel mineralization is associated with the Lamboo intrusion. An intersection of 18 m at 1.0% Ni was obtained in diamond drilling at Lamboo (Ni 1) (6155; Chadwick, 1986).

Energy mineral — uranium

At Spinifex (2727), secondary uranium mineralization has resulted from supergene enrichment of uranium on joints and breccia zones adjacent to the Revolver Fault. Up to 13.5% U₃O₈ was obtained in selected surface samples, but uranium values decrease rapidly with depth in trench sampling (Bianconi, 1973). Secondary uranium minerals have also been reported in association with the Esaw Monzogranite at A1 (6194).

Base metal — copper, silver (gold)

The oxide zone of the Sandiego deposit (Sandiego (oxide), 1128) has been classified in the 'regolith — residual and supergene' category (see **Stratabound volcanic and sedimentary mineralization**, p. 27).

Industrial mineral — magnesite

There is residual magnesite on the Lamboo intrusion at Lamboo (magnesite) (5163).

Mineral controls and exploration potential

Mineralization in kimberlite and lamproite intrusions

Precious mineral — diamond

The Argyle AK1 pipe is situated near the eastern margin of the Halls Creek Orogen, on a splay of the Halls Creek Fault that is inferred to be the zone of weakness along which the olivine lamproite magma ascended (Boxer et al., 1989). Interaction of the magma with groundwater in the partly consolidated sediments of the Carr Boyd Group into which the diatreme was intruded caused phreatomagmatic explosions resulting in mixing of sand grains with the magma to form sandy tuffs (Boxer et al., 1989). Inclusions within the diamonds indicate that they formed in both mantle eclogite and mantle peridotite, and are thus regarded as xenocrysts in their olivine lamproite host rock (Boxer and Jaques, 1990).

The Lissadell Road Dykes were probably intruded into an echelon tension gashes associated with the same deep-seated fracture as the Argyle diatreme, 6 km to the north (Boxer et al., 1989).

There is potential for further discoveries of lamproitic and kimberlitic diamond occurrences associated with the many deep-seated faults in the east Kimberley area. On the basis of the distribution of kimberlitic indicator minerals in stream gravels, Fielding (1983) suggested that there could be a further eight or more kimberlitic dykes in the Maude Creek area and five or more in the Devils Elbow area. There are also several alluvial diamond occurrences away from the palaeodrainages of the Argyle AK1 deposit that are worthy of follow-up for lamproitic or kimberlitic diamond deposits.

Mineralization in carbonatite and alkaline igneous intrusions

According to Andrew (1990), the Cummins Range (REE) carbonatite deposit lies at the Junction of the Halls Creek and King Leopold Orogens. This is not obvious on Plate 1, but there has been no recent mapping on MOUNT BANNERMAN. On the basis of Nd and Sr isotope data, Sun et al. (1986) suggested that the carbonatite was derived from an ocean island basalt-type source within a depleted mantle. Weathering processes including leaching, dissolution, and silicification of the carbonatite have led to enrichment in resistate minerals and concentration of REE, niobium, and phosphorus. Weathering has also caused hydration of biotite to form hydrobiotite and vermiculite (Andrew, 1990).

The Cummins Range (REE) deposit is a significant unexploited REE, niobium, and phosphorus resource. The presence of minor carbonatites at Copperhead and in the Speewah Valley, together with the presence of the Bow Hill lamprophyre dykes, suggests that the east

Kimberley may be an alkali igneous province. There is potential for other carbonatite deposits to be found along major tectonic boundaries such as the Halls Creek Fault zone.

Orthomagmatic mafic and ultramafic mineralization

Steel-industry metal — nickel (copper, cobalt)

Layered mafic intrusions of Hoatson's (1997, in press) Group V are hosts to most of the nickel occurrences in the Lamboo Complex. At the Sally Malay deposit, sulfides are concentrated at the base of the intrusion. In some deposits (e.g. Corkwood), there may have been some remobilization of sulfides as there is irregular mineralization at different levels within the intrusion. Hoatson (1997, in press) suggested that the feeder zone of Group V intrusions is particularly prospective for deposits similar to the Voiseys Bay Ni–Cu–Co deposit in Labrador, Canada (Naldrett et al., 1996).

Intrusions other than those of Hoatson's Group V are also prospective. Nickel mineralization occurs in ultramafic rocks within the Tickalara Metamorphics at Bow River and Cabernet, in Group IV (Wild Dog Creek Gabbro) at Mabel Hill within the Central zone of the Lamboo Complex, and in the Paperbark supersuite at Billymac Yard D in the Western zone of Lamboo Complex.

Steel-industry metal — PGE, chromium

Group I Panton intrusions host many of the PGE and chromium occurrences in the Central zone of the Lamboo Complex. According to Perring and Vogt (1991), chromitite layers in the Panton Sill resulted from chromite saturation when silica-contaminated resident magma mixed with new pulses of primitive magma. Platinum group element grades peak towards the base of the A-seam at the Panton deposit. Perring and Vogt (1991) suggests that the magma-mixing event induced sulfur saturation and that the immiscible sulfide phase rapidly scavenged PGE, Ni, and Cu from the liquid before concentrating onto the floor of the magma chamber. Platinum group element-bearing chromitite horizons in the Lamboo and Eastman Bore intrusions in the Central zone, and the Group II Springvale intrusion in the Western zone, probably have a similar origin. Hoatson (1997) suggested that the upper parts of the Group VI McIntosh intrusions have potential for remobilized Au–Pd–Cu mineralization similar to the Platinova reef of the Skaergaard intrusion in East Greenland (Anderson et al., 1998).

Steel-industry metal — titanium, vanadium

Ilmenite is hosted by upper gabbroic parts of the Group V Sally Malay intrusions indicating potential for economic titanium and possibly vanadium. Hoatson (1997) also suggested that the upper parts of the Group VI McIntosh intrusions have potential for stratabound titaniferous magnetite layers.

Base metal — copper (nickel, PGE, chromium)

Remobilized magmatic sulfides are associated with Group I Panton intrusions, Group V Sally Malay intrusions and the Spring Creek intrusion, and the Group VI McIntosh intrusion in the Central zone; and the Group II Springvale intrusion in the Western zone. All of the layered intrusions in the East Kimberley have potential to host this style of deposit.

Disseminated and stockwork mineralization in plutonic intrusions

Base metal — copper (molybdenum, silver, gold)

On the basis of the presence of magmatic–hydrothermal breccia dykes and styles of alteration typically associated with porphyry copper systems, Witt and Sanders (1996) concluded that the copper occurrences in the McHales Granodiorite of the Sally Downs supersuite were of porphyry copper style. There is also a possibility that the occurrences are related to hydrothermal activity along the Halls Creek and Angelo Fault zones (Witt and Sanders, 1996) or have been derived from the volcanic rocks unconformably overlying the granitoid (Scott, 1973). Sheppard et al. (1997c) and Tyler et al. (in prep.) suggest that older granitoids (1835–1820) of the Sally Downs supersuite were emplaced into a convergent continental margin, above a northwesterly dipping subduction zone. Therefore, there is the potential for porphyry-copper style mineralization in other granitoids of this age.

Pegmatitic mineralization

Sanders (1999) has suggested that the tin–tantalum–niobium–tungsten pegmatites in the Dockrell area are related to the waning stages of the 1788 Ma San Sou Monzogranite, as the imprecise SHRIMP date of 1740 ± 40 Ma obtained on one of these pegmatites is almost within error of the age of the San Sou Monzogranite. However, the San Sou Monzogranite is not an S-type granite (the type of granite usually associated with tin deposits).

All pegmatites discovered to date have been small, and the potential for an economic deposit seems to be low.

Skarn mineralization

Most of the occurrences described by companies as skarns have no obvious association with granitoids. However, tonalite has been reported in association with the Wills Creek B6 occurrence. It is possible that the Monkey Yard Tonalite is more widespread than indicated on DIXON and that it has played a role in forming the skarns in the Tickalara Metamorphics in the Central zone. However, the Tickalara Metamorphics contains medium- to high-grade rocks and therefore all calc-silicate horizons will be dominated by a 'skarn-like' mineralogy. The skarn-like

mineralogy is thus probably completely unrelated to the Monkey Yard Tonalite. The San Sou Monzogranite or pegmatites may have played a role in the Taylor Lookout and Mount Heartbreak areas in the southern part of the Eastern zone as suggested by Sanders (1999). An alternative explanation is that the calc-silicate horizons contained syngenetic or volcanic exhalative mineralization that was recrystallized during metamorphism (see **Stratabound volcanic and sedimentary mineralization** below, and **Stratabound sedimentary mineralization**, p. 52).

Stratabound volcanic and sedimentary mineralization

Base metal — copper, lead, zinc, silver

In the Central zone of the Lamboo Complex, the stratabound volcanic and sedimentary base-metal occurrences within the Koongie Park Formation are related to both proximal and distal felsic volcanism. Sulfides appear to have been mobilized during late folding. At Sandiego, rich shoots are aligned parallel to the fold axes on tight folds (Sewell, 1999). The ‘Mimosa interval’ of the Camp shale member hosts most of the occurrences in the Koongie Park area and can be regarded as highly prospective, especially in fold hinges where sulfides may have been concentrated.

In the Eastern zone of the Lamboo Complex, base metal mineralization in the Ilmars – Little Mount Isa – Twin Hills area is in distal volcanoclastic sedimentary rocks associated with felsic volcanic rocks within the Biscay Formation. At Ilmars Central, mineralization has been concentrated near cross-cutting east–west faults (Sewell, 1999). Little is known about the other occurrences in the Biscay Formation, but most of them are associated with interflow sedimentary rocks or tuffs and probably formed in a distal environment. The Biscay Formation has potential for low-grade base metal mineralization with higher grades in the vicinity of felsic volcanic rocks or where mineralization has been concentrated by folding or faulting. There is also potential for base metal mineralization in felsic volcanic rocks of the Maude Headley Member, Butchers Gully Member, and Ding Dong Downs Volcanics.

In the Kimberley Basin, the presence of copper occurrences in the Carson Volcanics over a strike length of greater than 70 km, together with the concentration of copper in chert, sandy interflow sediments, and tuffaceous flow tops of basalt of the Carson Volcanics, supports the interpretation by Planet Management and Research Proprietary Limited (1971) that the mineralization is volcanogenic in origin. Faulting appears to have been significant in concentrating the mineralization. The faults at Carson Volcanics 9 appear to be syndepositional because there is a change in stratigraphy to the north of this deposit (Planet Management and Research Proprietary Limited, 1971; Sanders, 1999). It is possible that other concentrations of copper may have formed where there are syndepositional faults.

In the Osmond Basin, the possible volcanogenic copper occurrence in felsic lithic vitric tuffs near the base

of the Bungle Bungle Dolomite at Paradise Creek (Cu B) (de Angelis, 1984) was not followed up by Samin Australia, but is worthy of further investigation.

In the Ord Basin, native copper in vesicular basalt of the Cambrian Antrim Plateau Volcanics is possibly of Michigan type, but there is no evidence that it has been concentrated by metamorphic processes as in the Michigan area (White, 1968). Harms (1959) interpreted the native copper in boulders of prehnite at Dunham Crossing as steam-hole fillings by late magmatic volatiles during solidification of the lavas. No economic concentrations of this style of mineralization have yet been found, but there is potential for new finds.

Steel-industry metal — tungsten

The stratabound scheelite mineralization in the Taylor Lookout area is confined to a calc-silicate horizon within the Biscay Formation. The mineralization is possibly related to the San Sou Monzogranite that outcrops 1–1.5 km away as suggested by Sanders (1999), but this would not explain the stratabound nature of the mineralization. The granite may, however, be responsible for remobilizing tungsten into the later cross-cutting veins in the area (see **Vein and hydrothermal mineralization**, p. 57). The origin of mineralization at Berthas Butt is highly controversial, but a volcanic exhalative model is possible (see **Stratabound sedimentary mineralization**, p. 52).

Precious metal — gold

Some gold occurrences within the Biscay Formation in the Eastern zone of the Lamboo Complex (e.g. Lulu, Freda, and Mount Carmel East 2) are possibly volcanogenic in origin. According to Kinnane (1983), gold in the Dockrell area is located around fossil volcanic fumarolic vents in the Biscay Formation, but no specific locations were given. The Butchers Gully Member also hosts possible volcanogenic gold mineralization. There is the possibility that further volcanogenic mineralization will be found, although most gold mineralization appears to have been remobilized into later structures (see **Vein and hydrothermal mineralization** below).

Speciality metal — REE, niobium, tantalum, beryllium, gallium

Ramsden et al., (1993) suggested that the ‘Niobium tuff’ of the Butchers Gully Member at the Brockman deposit was originally enriched in incompatible elements and fluorine, and that the mineralization is the result of alteration and remobilization of magmatic minerals such as columbite and zircon by F-rich deuteric solutions that were retained in the tuff. The presence of anomalous niobium in alkaline volcanic rocks of the Maude Headley Member and in trachyte within mafic tuffs that have been mapped as Biscay Formation suggests that these rocks may also have potential for REE – rare metal deposits, especially where the presence of fluorite indicates that F-rich solutions could have been available to transport REE, beryllium, tantalum, and niobium, and deposit them in economic concentrations.

Precious mineral — prehnite, agate

Prehnite has formed as a vesicle filling in the Cambrian Antrim Plateau Volcanics. There is potential for further deposits of this type within the Antrim Plateau Volcanics. There is also potential for agate deposits formed in a similar way.

Stratabound sedimentary mineralization

Clastic-hosted mineralization

Steel-industry metal — tungsten (molybdenum)

The origin of the stratabound scheelite mineralization in the Berthas Butt and South Dockrell areas is controversial. For the purposes of this study, the mineralization type is classified as 'stratabound sedimentary — clastic-hosted', but an alternative classification could be 'stratabound volcanic and sedimentary — volcanic oxide'. On the basis of the presence of gypsum casts and quartz layers that possibly replaced anhydrite, Ririe (1989) proposed that tungsten mineralization in the Berthas Butt area formed in a continental-sabkha playa basin, similar to Searles Lake in Canada where brines contain up to 70 ppm WO_3 and the playa muds contain up to 118 ppm WO_3 . Trudu (1990) questioned Ririe's (1989) interpretation and suggested that the mineralization was related to the Sophie Downs Granophyre Member, and that the tungsten-bearing solutions followed the most porous units in the stratigraphy, that is, the quartzites. However, recent studies (Blake et al., 1999b) show that the c.1910 Ma Sophie Downs Granophyre Member is unconformably overlain by the c. 1880 Saunders Creek Formation (see Table 1), so there can be no genetic relationship between the granophyre and mineralization in the Saunders Creek Formation.

Sanders (1999) suggested that the stratabound tungsten mineralization in both the Berthas Butt area and the Dockrell area was related to post-tectonic, granite-derived fluids that were focused into anticlinal domes and along zones of permeability. Sanders (1999) proposed that the fluids were related to 1740 ± 40 Ma tin-bearing pegmatites, which in turn may have been related to the waning stages of the c. 1780 Ma San Sou Monzogranite. However, the nearest outcrop of San Sou Monzogranite to the South Dockrell area is 13.5 km away, and the Berthas Butt area is 50 km from the nearest outcrop of granitoid of the San Sou suite. Therefore, there is unlikely to be any genetic relationship between the tungsten mineralization and the San Sou suite.

Textural and fluid-inclusion data (fluid inclusions within the scheelite have been annealed) indicate that the scheelite at Berthas Butt pre-dates the last main phase of deformation of the rocks (? M_4 ; Blight, 1983; Ririe, 1989). There is also textural evidence that scheelite in epidote quartzites and tourmalinites in the South Dockrell area pre-dates the main phase of regional metamorphism (? M_4 ; Goldner, 1984).

An alternative origin for the tungsten mineralization is that it may be related to volcanic exhalative activity or

hot-spring activity. For comparison, the Felbertal stratabound scheelite deposit in Austria, which is one of the largest tungsten deposits in the world, is hosted in a metabasalt sequence and is interpreted to have formed as hot-spring sinters in a narrow rift graben during the initial stage of island-arc evolution (Plimer, 1987). A similar genesis has been proposed for the amphibolite- and tourmalinite-hosted stratabound scheelite deposits in the Broken Hill area, New South Wales, and the tourmalinite-hosted tungsten mineralization in West Greenland (Plimer, 1987). In both the South Dockrell area and the Berthas Butt area, tungsten-bearing epidote quartzites are closely associated with volcanoclastic sedimentary rocks, suggesting a link with volcanic exhalative activity. The close association between epidote and scheelite could be explained by the replacement of feldspars in the quartzite by epidote resulting from exhalative activity. The tourmalinites at South Dockrell could be volcanic exhalative in origin, similar to that suggested by Plimer (1987) for scheelite-bearing tourmalinites elsewhere. The presence of tourmalinites would also be consistent with metamorphism of a boron-rich evaporite deposit. It is possible that volcanic exhalations were concentrated in a playa environment.

There is stratabound tungsten mineralization in a range of formations of different ages in the Eastern zone including the Ding Dong Downs Volcanics and the Saunders Creek Formation at Berthas Butt and, possibly, the Butchers Gully Member at South Dockrell. The calc-silicate rock-hosted tungsten and base metal mineralization in the Biscay Formation in the Taylor Lookout area is probably of similar origin. Although no economic mineralization of this style has been found in the study area to date, this type of mineralization has the potential to be economic, as illustrated by the Felbertal stratabound scheelite deposit in Austria. There is also potential for other exhalative-type mineralization (e.g. gold and base metal) to be associated with the tungsten. The presence of epidote quartzite, tourmalinite, exhalative chert, or calc-silicate rock could be used as an exploration guide.

Base metal — copper, lead, zinc

The base-metal bearing gossans in feldspathic sandstones of the Winnama Formation in the Central zone of the Lamboo Complex are inferred to be syngenetic in origin, although there is little detailed information about them. The base-metal bearing gossans hosted by graphitic schists of the Tickalara Metamorphics are possibly also syngenetic. However, as the host rocks have been strongly metamorphosed, it is not certain whether these occurrences are syngenetic or hydrothermal.

In the Kimberley Basin, the extensive malachite staining in shale at the base of the middle Pentecost Sandstone, over a strike length of 200 km around Menuairs Dome in the northwestern part of the study area, suggests a stratigraphic control. However, exploration drilling has shown that the grade drops off with depth, suggesting probable surface enrichment. This unit could be worth exploring in the vicinity of faults, where it is possible that economic grades of mineralization are concentrated.

In the Osmond Basin, native copper at Paradise Creek (Cu A) fills pore spaces in the hosting sandstone and siltstone of the Bungle Bungle Dolomite, but the source of the copper is unknown. It could be volcanic exhalative in origin because copper mineralization has been reported in association with tuffs at Paradise Creek (Cu A). It could also be of hydrothermal origin and related to the nearby fault.

Energy mineral — uranium

The controls on the uranium mineralization at Rustlers (in the Biscay Formation in the Eastern zone of the Lamboo Complex) are unknown. Horse 7 (in the Bonaparte Basin) and the Killi Killi deposits (in the Birrindudu Basin) are possibly sandstone-uranium type occurrences, but surface enrichment has probably played a significant role in concentrating uranium at these occurrences.

There has been considerable exploration for uranium both in the Birrindudu Basin and in the King Leopold Sandstone at the base of the Kimberley Basin, but no significant uranium mineralization has been found; most radiometric anomalies have been due to thorium. This tends to downgrade the potential of the area.

Speciality metal — REE

Fluviatile processes have probably played a significant role in concentrating heavy minerals at Junior where monazite is concentrated in heavy mineral bands in fine-grained sandstone of the Palaeoproterozoic Red Rock Formation. Xenotime in basal grits and conglomerate of the Gardiner Sandstone in the Birrindudu Basin at Killi Killi 1 and Killi Killi West could be detrital, or the product of surface enrichment.

Precious mineral — diamond

The gem-quality diamonds said to have been recovered from the basal Barramundi Conglomerate of the Canning Basin at Barramundi have kimberlitic rather than lamproitic characteristics and are inferred to have been transported from the Kimberley Craton to the east and deposited in the Devonian palaeoplacer (Davies, 1990). Since no further diamonds have been found in the area despite intensive drilling, there is some doubt about the authenticity of the reported occurrence and hence the prospectivity of the area may be downgraded.

Carbonate-hosted mineralization

Base metal — lead, zinc, copper (silver, gold, molybdenum, tungsten, tin)

A large number of occurrences of Mississippi Valley-type mineralization have been found in the Devonian limestones of the Lennard Shelf area of the Canning Basin, and it remains a highly prospective area.

The Kapok deposit probably has similar ore controls to the nearby Cadjebut mine. At Cadjebut, rhythmically banded ore is hosted by evaporites that Tompkins et al. (1994a, 1997) interpreted as the source of sulfur. A later phase of mineralization is represented by sulfides in vugs

and cavities. The second phase of mineralization is interpreted by Tompkins et al. (1994b, 1997) to be related to a period of reefal platform emergence and minor uplift of the Lennard Shelf coinciding with the waning stages of the Alice Springs Orogeny. The Cadjebut and Kapok orebodies are situated close to the Cadjebut Fault and a splay of this fault. The Cadjebut Fault is a synsedimentary (Late Devonian), normal fault, which in turn is a splay of the Pinnacle Fault. The Pinnacle Fault was probably the main basin-tapping fluid conduit for the mineralizing brines for both mineralizing events (Tompkins et al., 1997). Oil and gas generation in the basin, triggered by a fall in sea level, may have been a contributing factor to the generation of overpressured fluid that deposited the breccia ore (Tompkins et al., 1997).

The Lindner Fault probably played a role in the localization of mineralization at Horse Spring Range, 59 km north of Kapok.

In the Osmond Basin, the stratabound nature of the galena in carbonates of the Bungle Bungle Formation suggests that it is syngenetic in origin. Soft-sediment compaction and dewatering may have resulted in mobilization and concentration of the sulfides as suggested by de Angelis (1983). Organic matter may have played a role in precipitating or concentrating the base metals. It is also possible that the ore metals are volcanic exhalative in origin because chert is associated with the lead mineralization at Paradise Creek (Pb), and de Angelis (1984) has reported copper mineralization associated with felsic, lithic, vitric tuff at Paradise Creek (Cu B). This area is considered to be prospective, especially because most of the exploration holes drilled during the 1980s failed to meet their target or ended prematurely in mineralization due to difficult drilling conditions.

Sanders (1999) proposed a granitoid-related skarn origin for the minor copper(–tungsten–gold–tin) mineralization in calc-silicate rocks and marbles of the Tickalara Metamorphics in the Central zone at Alice Downs, but Chadwick (1985) considered that they could be primary exhalative in origin.

Although syngenetic copper mineralization in dolomite and shale in the upper part of the Teronis member of the Elgee Sandstone is widespread, no economic mineralization has yet been found in the Kimberley Basin. It is possibly worth exploring in the vicinity of faults where there may have been some greater concentration of copper.

In the Ord Basin, low-grade copper mineralization at Mount Elder, in limestone of the Cambrian Panton Formation of the Goose Hole Group, has been interpreted by Dow and Gemuts (1969) to be sedimentary in origin, but they did not consider the mineralization to be economic.

Undivided mineralization

Base metal — copper, lead, zinc, silver

In the Lamboo Complex, the origin of the base metal mineralization in the Tickalara Metamorphics and the

Biscay Formation is unknown, but the mineralization could be volcanic exhalative in origin. Copper mineralization in the Headleys Limestone of the Ord Basin at Discovery Yard was possibly derived by secondary enrichment from the underlying basalt of the Antrim Plateau Volcanics (Ball, 1966).

Steel-industry metal — vanadium, chromium, molybdenum

In the Central zone of the Lamboo Complex, at Sally Malay Bore 5 and Sally Malay Bore 6, the association of the vanadium–chromium–molybdenum gossans with BIF of the Tickalara Metamorphics suggests that mineralization could be volcanogenic exhalative in origin.

Precious metal — gold

The gold in silicified black shale and dolomite of the Biscay Formation at Woodward Creek (Au) is possibly syngenetic or volcanic exhalative in origin.

Sedimentary mineralization — banded iron-formation (supergene enriched)

Iron

Field evidence, including the presence of cross-bedding visible in the hematite ore in the Pompeys Pillar area, indicates that the ore has been little changed since deposition, apart from the leaching of silica, and that the iron is sedimentary in origin (Harms, 1959). The subsequent discovery of enormous resources of iron ore in the Pilbara has made it unlikely that iron ore in the Pompeys Pillar area will be mined in the foreseeable future.

Vein and hydrothermal mineralization

Undivided mineralization

Many of the vein- and hydrothermal-related mineral occurrences in the east Kimberley are polymetallic, but they have been subdivided according to the main commodity (see **Mineralization**, p. 19). Sanders (1999) suggested that there is a sequential pattern of metal zonation (i.e. Au-only to Au–Pb–As–Cu(–Zn) to Cu–Au–Ag to Cu–Mo–Ag–Sn–W to W(–Sn) to Sn–Ta–Nb) that shows a positive correlation with the M_4 metamorphic grade of the host rock in the Eastern zone of the Lamboo Complex. According to Sanders, gold-only occurrences are mainly hosted by rocks of low metamorphic grade (greenschist facies); base-metal bearing gold occurrences are mainly hosted by rocks metamorphosed to upper greenschist to lower amphibolite facies; and W(–Sn) or Sn(–W) mineralization is hosted by rocks metamorphosed at low to mid-amphibolite facies. Sanders (1999) states that there was an absence of gold mineralization in all

areas that contain tin or tungsten mineralization, and hypothesizes that ‘the gold mineralization may be part of a larger metal suite (of possible granite association) that was deposited at a time when the geothermal gradient was still relatively high, but post-peak metamorphic in timing’.

Dow and Gemuts (1969) suggested that gold and base metal mineralization in the Lamboo Complex may be related to the Sophie Downs suite, but these granitoid rocks have been dated at c. 1910 Ma (Blake et al., 1999b) and thus pre-date the gold and base metal mineralization hosted by the c. 1880 Ma Biscay Formation. The only granitoid with an age that could be related to mineralization in the Eastern zone is the c. 1790 Ma San Sou Monzogranite. Sanders (1999) suggests that the San Sou Monzogranite is more extensive in the subsurface, but there is no evidence for this.

Alternative controls on vein and hydrothermal mineralization are given below.

Precious metal — gold

Lamboo Complex — Eastern zone

Dow and Gemuts (1969) were the first to recognize the strong stratigraphic control on gold mineralization in this area. They pointed out that all of the economic mines are located close to the boundary between the Biscay and Olympio Formations. This study has provided further evidence of strong stratigraphic control on mineralization. As stated in the **Mineralization** section (p. 19), 112 of the total of 262 vein and hydrothermal gold occurrences within the whole of the east Kimberley area are within the Biscay Formation in the Eastern zone of the Lamboo Complex. Another 52 occurrences are in the Olympio Formation, generally close to its contact with either the Biscay Formation or the Butchers Gully Member, and 21 occurrences are within the Butchers Gully Member, near the base of the Olympio Formation. There are hardly any gold occurrences in the Olympio Formation away from its contact with the Biscay Formation.

Warren (1994b) suggested that the Butchers Gully Member, which contains high concentrations of incompatible elements such as zirconium, yttrium, and niobium, was the most likely source-rock for both gold and chemically active fluids. The mafic volcanic rocks, volcanoclastic rocks, and interflow sedimentary rocks of the Biscay Formation are another possible source of the gold. Kinnane (1983) suggested that the gold in the Dockrell area was located around fossil volcanic fumarolic vents in the Biscay Formation.

Gold has been mobilized into steeply dipping, northeasterly trending shears, conformable quartz veins, and cross-cutting quartz veins. Dow and Gemuts (1969) stated that the quartz reefs in the Ruby Queen area have been folded, but Sanders (1999) could find no evidence of this. According to Griffin and Tyler (1992) and Warren (1997), the mineralization at Ruby Queen occurs along sinistral strike-slip faults and antithetic tension veins that are related to the fourth deformation in the Eastern zone (D_6). This is related to the Palaeozoic Alice Springs Orogeny and therefore the gold deposits are unlikely to

be related to the M_4 isograds as suggested by Sanders (1999). At Palm Springs, the quartz stockwork mineralization in trachyandesite of the Butchers Gully Member is a series of tension veins, related to the same shear system as that at Ruby Queen (Sanders, 1999; Fig. 18). Most of the other workings in the Brockman – Mount Bradley area are also in the trachyandesite of the Butchers Gully Member, which is structurally repeated by high-amplitude isoclinal folding (D_5) of the Mesoproterozoic Yampi Orogeny, suggesting that this unit fractured more readily during sinistral faulting (D_6 ; Sanders, 1999). At Old Halls Creek, mineralization occurs along the Halls Creek Fault, in shears parallel to this fault, and in related tension veins. The Kerry–Reform–Darcys group of workings in the Biscay Formation is related to northwesterly to north-northwesterly trending shears parallel to the arcuate boundary between the Biscay and Olympio Formations. Gold mineralization in the Dockrell and Grants Creek area is also along shears parallel to the regional faults; the richest mineralization in the Dockrell area developed at the intersection of two shears at McNeils.

The main gold-mining centres in the Eastern zone have been intensely prospected for gold. However, in the past, only rich high-grade veins and lodes were mined. If the gold could not be readily extracted by dollying or amalgamation, the area was abandoned. Veins with abundant sulfide, such as those at the original Butchers Gully (now Palm Springs) mine, could not be profitably worked. With new processing technology and open-pit mining techniques, it is probable that some of the old mining areas could be profitably worked. In particular, quartz stockwork mineralization in trachyandesite of the Butchers Gully Member in the Brockman – Mount Bradley area is an obvious exploration target.

Further south in the Eastern zone, where exposure is poorer and there has been less prospecting and exploration activity, there is potential for further discoveries of gold. The mafic rocks in the Christmas Creek area would be worth further investigation. The anomalous bismuth and gold mineralization in schist and dolomitic rock, possibly belonging to the Biscay Formation or Butchers Gully Member at Shiddi Creek, is also worth following up.

Lamboo Complex — Central zone

In comparison to the Eastern zone, there is very little known gold mineralization in the Central zone. The most significant gold occurrence in the Central zone is at Nicolsons Find, where auriferous quartz veins in the Koongie Park Formation are related to the northerly trending, steeply dipping Nicolsons Fault (Sanders, 1999). Other shear-related gold mineralization in the Nicolsons Find area is hosted by the Loadstone Monzogranite and the Lamboo intrusion. The gold could have been derived from volcanogenic mineralization in the Koongie Park Formation or from primary mineralization in the Lamboo intrusion.

An epithermal origin for the gold at Lady Helen is suggested by the widespread occurrence of hydraulic fractures and brecciation and the vuggy nature of the veins (Pirajno et al., 1994). This evidence also suggests that the veins formed in late brittle structures that post-dated the

Loadstone Monzogranite and were probably initiated at about 1300–1100 Ma (Pirajno et al., 1994; Sanders, 1999).

There is potential for further gold mineralization in the Central zone along the Springvale Fault and other late faults, especially where they cut the Koongie Park Formation.

Lamboo Complex — Western zone

In the Western zone, the Glidden Fault and Lubbock Range Fault have played a role in localizing gold and base metals at Me No Savvy and at Black and Glidden (also at Soda Springs in the Central zone to the south of the Lubbock Range Fault). The Amherst Metamorphics is a possible source of the gold. Low-grade gold–silver mineralization in epithermal veins in felsic tuffs of the Whitewater Volcanics at Jailhouse Creek is localized on a west-northwesterly trending fault. It is possibly related to the same epithermal system that produced the fluorite and base metal mineralization in the Speewah Basin. There is potential for further epithermal mineralization, particularly in the vicinity of the Dunham and Greenvale Faults.

Speewah Basin

To date, only minor gold mineralization has been found in association with base-metal and fluorite veins in the Speewah Basin, but Rogers (1998) suggested that there could be rich epithermal gold mineralization at depth.

Granites–Tanami Complex

Gold mineralization in the Larranganni area shows a particularly strong stratigraphic control with the occurrences being confined to two sequences of folded mafic volcanic and sedimentary rocks that have been informally named the Bald Hill sequence and the Quartz Ridge sequence (Glengarry Resources NL, 1999, written comm.; Quinn, 1999; Fig. 19). Glengarry Resources NL (1999, written comm.) has correlated these sequences with the Mount Charles Beds at the Tanami and Granites–Callie mines in the Northern Territory. The company also highlighted other similarities between the Larranganni mineralization and that of the Tanami and Granites–Callie mines: similar style of alteration (carbonate, sericite, quartz, iron, and chlorite), folded rocks along a northerly trend, major west-northwesterly cross-cutting structures (Trans-Tanami Fault System), and a strong regional gravity anomaly.

The Kookaburra and Sandpiper occurrences are localized on a west-northwesterly fault suggesting that these cross-cutting structures were important channelways for ore fluids. The widespread occurrence of large, randomly orientated muscovite flakes in felsic arenite, together with the gravity low, suggests the presence of a granitoid body beneath the Larranganni area (Glengarry Resources NL, 1999, written comm.). The granitoid could have been the source of the ore fluids; however, the strong stratigraphic control on the gold occurrences suggests that either the mafic rocks or volcanogenic exhalative chert, or both, were the source of the gold and that the granitoid acted as a heat source to mobilize gold-bearing fluids into

the west-northwesterly trending fractures. Alternatively, these fault systems could be very young (e.g. Alice Springs Orogeny) and mineralization completely unrelated to the granite. As exploration continues in this area, more definitive evidence on the controls on mineralization may be revealed. Very little is known about the ore controls of the newly discovered Coyote deposit, but Glengarry Resources NL (1999b) regards the entire area between Coyote and the Larranganni area to be prospective, in addition to areas further west (e.g. around the Camel prospect).

Gold at the Venus prospect appears to have been localized in shears, saddle reefs, and veins related to the tight folding of the Killi Killi 'Beds'. As the mineralization is close to the unconformity between the Killi Killi 'Beds' and the Gardiner Sandstone of the Birrindudu Group, the unconformity may have played a role in the localization of the gold. Other shears and veins in the vicinity of the unconformity are worth exploring for gold (see also **Unconformity-related mineralization**, p. 57).

Base metal — copper, lead, zinc (silver, gold, fluorite, barite)

In the Eastern zone of the Lamboo Complex, base metals are associated with many of the auriferous veins and, in these cases, would obviously have similar controls to the gold mineralization. The copper and zinc occurrences in the Saunders Creek and Sophie Downs areas are related to east-northeasterly and northerly trending faults. The copper may have been derived from primary copper in the Ding Dong Downs Volcanics or the Biscay Formation, whereas the zinc may have been derived from black shales in the Biscay Formation. The Mount Kinahan copper occurrences are in greywacke of the Olympio Formation along an east-northeasterly trending fault that cuts through Neoproterozoic sedimentary rocks of the Albert Edward Group. It is not known whether the mineralization was on a pre-existing fault, which was reactivated during the Neoproterozoic, or whether the mineralization is Neoproterozoic or younger in age (Sanders, 1999). The age and controls on coarse-grained galena and purple fluorite in a breccia zone within the Maude Headley Member at Ding Dong Downs (Pb–F) is also uncertain, with the breccia being interpreted by Fletcher (1981) as a breccia pipe but by Johnson (1981) as a tectonic breccia.

Although there are many base-metal vein occurrences in the Eastern zone, most are minor in extent and, to date, there is no evidence that any of them would be economic to work.

In the Central zone, copper-bearing quartz veins hosted by the Tickalara Metamorphics and the McHale Granodiorite at the Calico Spring prospect are related to a north-northeasterly trending splay off the Halls Creek Fault. Copper mineralization in shears in the Eastman Bore intrusion is possibly remobilized primary mineralization. Base metal mineralization associated with shear zones in tuffs of the Koongie Park Formation could be remobilized volcanogenic mineralization. Lead–barite mineralization at Boxers is possibly related to the Carr Boyd Fault.

In the Western zone, mineralization at Black and Glidden and Me No Savvy is hosted by the Amherst Metamorphics and related to the Lubbock Range and Glidden Fault system. The mineralized quartz veins are typically massive and probably mesothermal in origin (Sanders, 1999); however, at both of these localities there is also carbonate-rich Pb–Zn–Ag–F–Ba mineralization that Sanders (1999) has interpreted as being either closely related to the quartz vein mineralization (and of possible granitic origin) or as a later phase of mineralization that has followed pre-existing mineralized structures. The latter theory is favoured here because the faults post-date the granitoids. Epithermal copper mineralization in the Whitewater Volcanics is associated with the Greenvale Fault and other late faults.

In the Speewah Basin, Rogers (1998) has suggested that the mineralization is epithermal and related to caldera collapse and later movement of hydrothermal solutions along fractures (see section immediately below on fluorite). Rogers (1998) has also suggested that there could be vertical zonation, with various levels prospective for gold and base metals.

In the Kimberley Basin, copper mineralization at Campbellmerry is related to narrow shears parallel to major northwesterly trending faults in the Pentecost Sandstone (Harms, 1959).

In the Bastion Basin, minor veins of copper mineralization in siltstone of the Mendena Formation at Plants are possibly remobilized syngenetic sulfides.

In the Carr Boyd Basin, disseminated galena and lead oxides in silicified breccia zones within sandstone of the Golden Gate Siltstone at the Carr Boyd prospects may be related to splays of the Carr Boyd Fault.

In the Ord Basin at Brook Creek, minor veins with malachite in the Cambrian Headleys Limestone are probably related to the nearby west-northwesterly trending fault. The proximity of the mineralization to barite and fluorite veins suggests that the mineralization is epithermal.

In the Granites–Tanami Complex at Errol Bore, copper mineralization is related to an east-southeasterly trending fault in the Killi Killi 'Beds'.

Industrial mineral — fluorite

In the Speewah Basin, fluorite veins in the Speewah Group and Hart Dolerite in the Speewah Valley area have epithermal textures and there is evidence of volcanic pyroclastic vent-like bodies (Rogers, 1998). Fluid-inclusion studies indicate that the veins have trapping temperatures of about 150 to 170°C (Rogers, 1998). Galena from fluorite veins has given model-lead ages between 15 and 131 Ma and direct Sm–Nd dating of the veins gave an age of 120 Ma (Rogers, 1998). If the age dating is correct, then this suggests a previously unrecognized mineralizing event. The fact that one fluorite vein cuts the Antrim Plateau Volcanics indicates that the mineralization is post-Cambrian (Rogers, 1998). A possible carbonatite dyke is associated with fluorite

mineralization at West Ridge and may be genetically related to the mineralization (Alvin, 1993; Rogers, 1998). On the basis of aeromagnetic data and aerial photographs, Rogers (1998) identified circular and curvilinear structures inferred to be ring structures within the Speewah Dome. Rogers (1998) suggested that fractures related to caldera collapse and later faulting may have acted as pathways for hydrothermal fluids.

In the western zone of the Lamboo Complex, minor fluorite associated with uranium at the Frog prospect, in the Whitewater Volcanics, is probably epithermal in origin, as is the fluorite associated with brecciated cross-cutting quartz veins in felsic volcanic rocks of the Maude Headley Member at Ding Dong Downs (F) in the Eastern zone.

Industrial mineral — barite

Barite associated with some of the fluorite veins in the Speewah Basin would have the same ore controls as the fluorite (see above). The barite occurrence in the Mabel Downs Tonalite at Mabel Downs (Ba) in the Eastern zone of the Lamboo Complex is probably related to a fault parallel to the Halls Creek Fault.

Steel-industry metal — tungsten

Scheelite occurrences at Ruby Plains and Balara in the Eastern zone of the Lamboo Complex are possibly related to northwesterly trending faults. Other tungsten vein mineralization in the Eastern zone in the Berthas Butt, Taylor Lookout, and Mount Dockrell areas is associated with stratabound scheelite mineralization (see **Stratabound sedimentary mineralization**, p. 52) and has probably been remobilized from this style of mineralization.

Steel-industry metal — nickel

Nickel–copper mineralization on the faulted contact between the Springvale intrusion and the Airfield Granodiorite is probably remobilized primary sulfides.

Steel-industry metal — molybdenum

The nature of the rumoured molybdenite occurrence at Mabel Hill is not known, but it is worthy of follow-up.

Energy mineral — uranium

Uranium occurrences in the Western zone of the Lamboo Complex are localized along the Dunham Fault zone and other late faults in the Whitewater Volcanics.

Speciality metal — REE

At the John Galt prospect, xenotime mineralization in sandstone of the Red Rock Formation within the Red Rock Basin is probably related to the northeasterly trending fault that forms the eastern wall of the mesa on which the deposit is located. Xenotime from Browns Range Dome in the Archaean Lucas Craton is associated with vuggy quartz and hematite alteration in the Browns

Range Metamorphics, but controls on mineralization are unknown.

Unconformity-related mineralization

Energy mineral — uranium

The main control on the uranium deposits in the Gardiner Range area is the unconformable contact between the Killi Killi ‘Beds’ and the Gardiner Sandstone. Faults and the presence of carbonaceous shales within the Killi Killi ‘Beds’ probably also played a role in the localization of the deposits. Despite extensive exploration, no economic deposits have been discovered.

Regolith mineralization

Precious metal — gold

Alluvial and eluvial gold have been worked since 1885. They continue to provide an attractive target for small companies and individual metal detector operators, but tonnages are typically too low to be of interest to major companies.

Precious mineral — diamond

Gem-quality diamonds in present day and palaeodrainage channels draining the Argyle AK1 pipe will continue to be a high-priority exploration target. Alluvial diamonds elsewhere in the area are significant because they suggest that there are other sources of primary diamonds that have not yet been located.

Steel-industry metal — nickel

Lateritic nickel occurrences associated with the Lamboo intrusion, and possibly other ultramafic intrusions in the Lamboo Complex, may become exploration targets now that the technology for treating lateritic ore has been developed.

Other commodities — tin, tantalum, scheelite, heavy minerals

Minor occurrences of alluvial tin, tantalum, niobium, scheelite, and heavy minerals have been reported from the study area. None of these have yet been proved to be economic, but they are a useful indicator to primary mineralization in the area.

Conclusion

The east Kimberley region hosts a wide range of mineral commodities with a variety of styles. The Argyle AK1 mine is the World’s largest diamond mine and the Kapok mine is part of Western Australia’s second-largest zinc-producing project (Western Metals’ Lennard Shelf project). Significant deposits of base metals, gold, PGE, nickel, REE, niobium, tantalum, tin, and fluorite have also been discovered. The mineralization shows a very strong

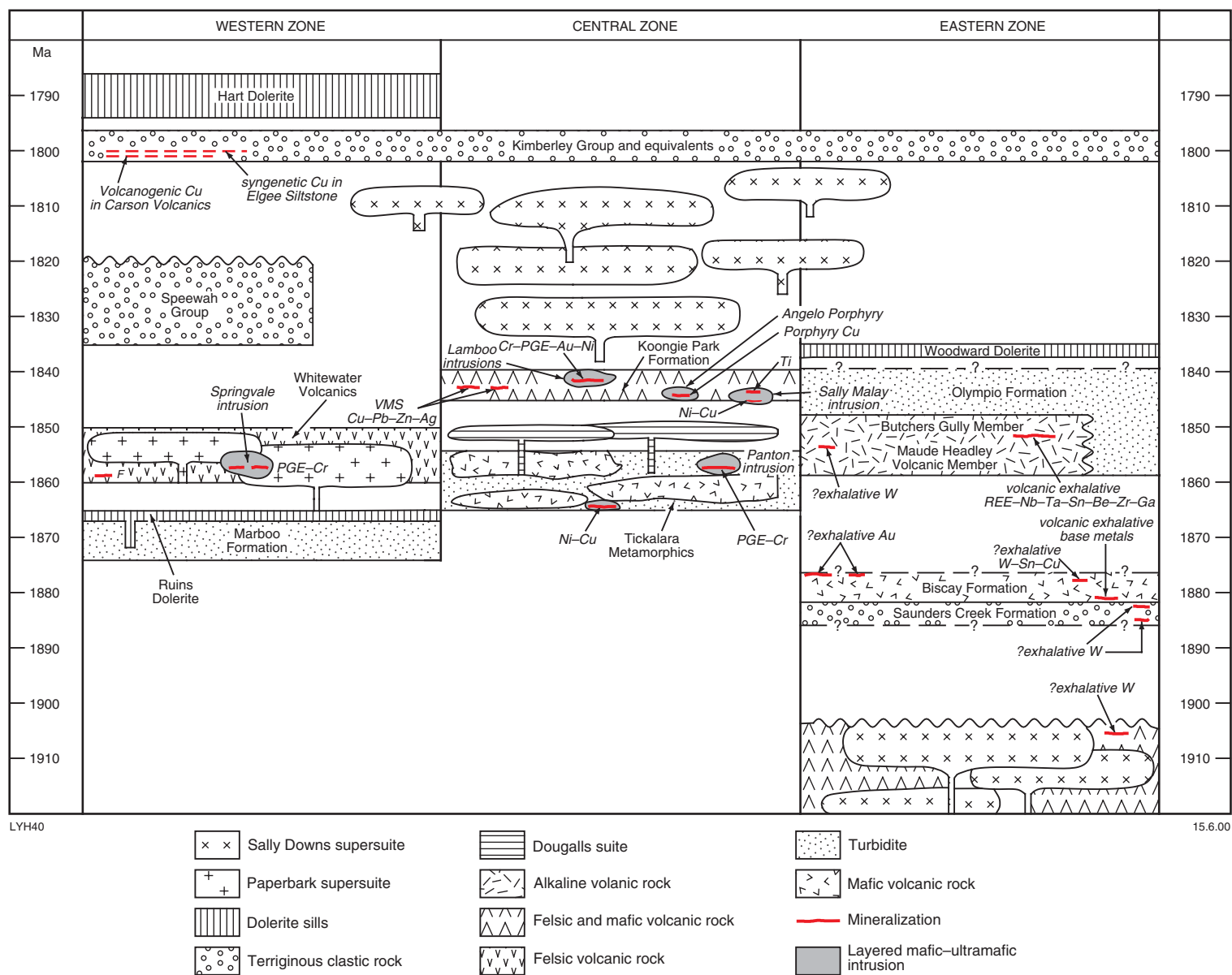


Figure 21. Time-space plot for the three zones of the Lamboo Complex (from Figure 2, this Report) showing mineralization

relationship to the stratigraphic and tectonic development of the area. This relationship should be a useful tool in future exploration of the area.

In the Palaeoproterozoic Lamboo Complex, where many of the occurrences are sited, different types of mineralization occur in the Eastern, Central, and Western zones (Fig. 21).

Most gold occurrences are in the Eastern zone of the Lamboo Complex. Nearly half the gold occurrences in the east Kimberley are in the Biscay Formation, with significant numbers in the Butchers Gully Member and in the Olympio Formation close to its contact with the Biscay Formation. Most of these occurrences are situated on northeasterly trending faults, but these faults are typically barren away from the contact between the Biscay and Olympio Formations. The Biscay Formation, and possibly the Butchers Gully Member, probably contained significant amounts of volcanic-exhalative gold that was later mobilized along the northeasterly trending shears, probably during the Alice Springs Orogeny. Although historical production of gold in the east Kimberley is quite low, a significant amount of gold probably went unrecorded and most of the workings were very shallow. The fact that the open-pit mine at Palm Springs produced 2272 kg gold in its two years of operation, which is more than twice the historical (pre-1964) gold production for the east Kimberley, indicates that there is potential for other old workings to be mined by opencut techniques.

Alkaline volcanic rocks of the Butchers Gully Member in the Eastern zone also host the large Brockman REE–Nb–Ta–Sn–Be–Zr–Ga deposit. This deposit remains undeveloped but is a significant resource. Anomalous niobium values have also been reported from felsic volcanic rocks in the Maude Headley Member and Biscay Formation indicating that these units are also prospective for this unusual style of volcanogenic deposit.

Epidote quartzites in the Ding Dong Downs Volcanics and Saunders Creek Formation host stratabound tungsten mineralization in the Berthas Butt area. Stratabound tungsten is also hosted by epidote quartzites and tourmalinites of the Butchers Gully Member in the North Dockrell area and by calc-silicate rocks within the Biscay Formation in the Taylor Lookout area. This mineralization is interpreted here to be volcanic exhalative in origin. Although the known occurrences are subeconomic, there is potential for a world-class tungsten deposit to be found. The Biscay Formation is also prospective for volcanic exhalative base-metal mineralization: examples include Ilmars and Little Mount Isa.

In the Central zone of the Lamboo Complex, the Koongie Park Formation hosts massive-sulfide mineralization that is probably remobilized VMS-style mineralization. Resources of copper, lead, zinc, and silver have been defined at Sandiego and Onedin in the Koongie Park area and at Mount Angelo North indicating that the Koongie Park Formation is highly prospective. A large low-grade copper deposit of possible porphyry-copper

style mineralization has been defined at Mount Angelo Porphyry in the Angelo Microgranite. In addition, the McHales Granodiorite also contains possible porphyry copper mineralization.

Platinum group element – chromium mineralization in the Central zone is hosted by the c. 1855 Ma Pantom intrusions, and also by the younger Lamboo and Eastman Bore intrusions. Nickel–copper mineralization is at the base of the c. 1845 Ma Sally Malay layered intrusions and includes the significant Sally Malay deposit. Metamorphosed layered mafic intrusions within the Tickalara Metamorphics also host significant nickel–copper mineralization. No significant mineralization has yet been found in the c. 1830 Ma McIntosh intrusion, but because of its large size, it has potential to host vanadiferous magnetite or ilmenite deposits.

In the Western zone, minor PGE–chromium mineralization is hosted by the c. 1855 Ma Springvale intrusions. Minor epithermal gold and fluorite mineralization is also hosted by the Whitewater Volcanics in the Western zone.

The major faults that define the boundaries of the three zones in the Lamboo Complex are deep crustal structures that have moved intermittently from the Palaeoproterozoic to the Phanerozoic, and thus have significant potential to host kimberlitic and lamproitic diamond deposits and carbonatites. The Argyle AK1 lamproite, the largest diamond mine in the world, is located on a splay of the Halls Creek Fault, which defines the boundary between the Eastern zone and the Central zone. The Cummins Range Carbonatite may also be related to the Halls Creek Fault.

The recent discoveries of gold at the Coyote prospect and in the Larranganni area demonstrate that the Western Australian portion of the Granites–Tanami Complex is as prospective as the Granites – Callie – Dead Bullock Soak area in the Northern Territory.

Younger sedimentary basins also host significant mineralization. Mississippi Valley-type base metal mineralization is being mined at Kapok (part of Western Metals' Lennard Shelf project) in the Canning Basin, and the large number of occurrences of this type confirms the potential of the Devonian reef limestones for this type of deposit. Possible Mississippi Valley-type mineralization is also present in the Bungle Bungle Dolomite in the Osmond Basin. The Speewah Basin contains significant epithermal fluorite and base metal mineralization, which is possibly related to caldera collapse during the Cretaceous (Rogers, 1998). The Carson Volcanics in the Palaeoproterozoic Kimberley Basin hosts copper mineralization of possible volcanic-exhalative origin. Stromatolitic dolomites of the Teronis Member of the Elgee Siltstone, and shale at the base of the Middle Pentecost Sandstone in the Kimberley Basin, host subeconomic copper mineralization. Basalts of the Cambrian Antrim Plateau Volcanics have potential for Michigan-style copper mineralization, and also for ornamental stones such as prehnite and agate.

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Appendix 1

List of mineral occurrences

MINERAL OCCURRENCES OF THE EAST KIMBERLEY

* KEY TO OPERATING STATUS

Bold numbers	Operating mine
<i>Bold and italic numbers</i>	Abandoned mine
Plain numbers	Mineral deposit
<i>Italic numbers</i>	Mineral occurrence or prospect

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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PRECIOUS MINERAL

☆	Kimberlite and lamproite intrusions			
1106	Dmd	434452	8152084	Argyle Ak1
1110	Dmd	426237	8147785	Lissadell Road Dykes
1115	Dmd	368433	8148169	Maude Creek Dyke
△	Stratabound volcanic and sedimentary – undivided			
6406	Prh	438879	7997028	Valley
6407	Prh	438935	7996570	Mound
6409	Prh	437483	7995566	New Deposit
□	Stratabound sedimentary – clastic-hosted			
5611	Zebr	469282	8191257	Lake Argyle 1
6386	Dmd	198312	8002677	Barramundi Range 1
6497	Zebr	496776	8222894	Lake Argyle 2
○	Regolith – alluvial to beach placers			
1107	Dmd	436647	8150862	Limestone Creek
1108	Dmd	434447	8153897	Upper Smoke Creek
1109	Dmd	435745	8155561	Lower Smoke Creek
1113	Dmd	455253	8160220	Bow River (alluvials)
2764	Dmd	475733	8103158	Hidden Valley
3002	Dmd	426332	8144958	Argyle Dykes 1
3003	Dmd	425833	8145758	Argyle Dykes 2
3004	Toz	425733	8148058	Argyle Dykes 3
3309	Dmd	450033	8173958	Smoke Creek
3312	Dmd	444033	8169758	Flying Fox Creek
3335	Dmd	426033	8152958	Glen Hill A
3336	Dmd	416333	8158858	Pompey A
3337	Dmd	416733	8168958	Pompey B
3338	Dmd	418433	8178608	Blatchford A
3339	Dmd	418633	8178808	Blatchford B
3340	Dmd	418133	8182708	Blatchford C
3341	Dmd	427632	8186258	Blatchford D
3342	Dmd	418033	8182158	Blatchford E
3343	Dmd	430933	8161358	Glen Hill B
3344	Dmd	430533	8164458	Glen Hill C
5164	Dmd	349437	8187384	Durack River 1
5165	Dmd	347744	8170450	Durack River 2
5166	Dmd	353663	8133558	Durack River 3
5167	Dmd	347132	8124598	Durack River 4
5192	Dmd	341351	8138426	Site 3603
5376	Dmd	345986	8115744	Lightning Creek
6028	Dmd	457503	7935715	Gordon Downs 1
6029	Dmd	457821	7935229	Gordon Downs 2
6030	Dmd	453771	7932270	Gordon Downs 3
6387	Dmd	190216	8000949	Barramundi Range 2
6420	Dmd	189147	7882151	Salt Creek
6422	Dmd	187735	7874775	Wattle Spring Creek

PRECIOUS METAL

+	Orthomagmatic mafic and ultramafic – layered mafic			
12	Pt Pd Cr	379273	8042874	Mini M (Cr-PGE) 2
814	Cr	354583	8035158	Springvale 1
819	PGE Cr	378933	8043558	Mini M

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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PRECIOUS METAL

▲	Stratabound volcanic and sedimentary – undivided			
2338	Au Ag As (Pb Cu)	315634	7921702	Dockrell East 1
5575	Au	357151	7987370	Campbell
6393	Au	315729	7921948	Dockrell East 2
6394	Au	315967	7922448	Green Leader
6437	Au	306793	7918506	Freda
6440	Au Pb Cu	309320	7921377	Lulu
■	Stratabound sedimentary – clastic-hosted			
5719	Au Ag (Cu U)	480045	7865021	Mount Brophy
+	Orthomagmatic mafic and ultramafic – layered mafic			
825	PGE	376963	8037258	Panton
828	Pd Pt Cr V Ti	380965	8043186	Togo 5A
835	PGE Cr Au	245679	7928396	Eastman Bore (PGE 1)
1044	PGE Au Cr	247932	7930258	Eastman Bore (PGE 2)
3189	PGE	386532	8073058	Koondooloo Yard
3517	PGE Cr	352933	8036758	Springvale 2
3536	PGE Cr	357133	8033358	Springvale 3
3539	PGE Cr	386623	8049268	Melon Patch 3
3543	PGE Cr	387233	8049658	Melon Patch 4
3545	PGE Cr	385233	8048658	Melon Patch 5
3546	PGE Cr	380633	8042158	Melon South
3557	Au PGE Cr	387633	8049858	Melon Patch 6
3561	PGE	387833	8049958	Melon Patch 7
3565	PGE Cr	378033	8039558	Panton North 1
3567	PGE Cr	377933	8038558	Panton North 2
3568	PGE Cr	378433	8040558	Panton North 3
3569	PGE Cr	380633	8043358	Alice Downs 3
3571	PGE	382233	8046358	Alice Downs 4
5470	PGE	329233	7963323	Dusty Bore 2
5697	PGE Au Cr	247822	7930225	Eastman Bore (PGE 3)
5698	PGE Au Cr	247681	7930120	Eastman Bore (PGE 4)
5705	PGE Cr	249251	7930174	Eastman Bore (PGE 5)
5951	Cr PGE	250973	7930611	Eastman Bore (PGE 6)
5952	Cr PGE	250680	7930382	Eastman Bore (PGE 7)
5953	PGE Cr	249846	7930286	Eastman Bore (PGE 8)
5954	PGE	248163	7929897	Eastman Bore (PGE 9)
5955	PGE	248480	7930493	Eastman Bore (PGE 10)
5956	PGE Cr	248099	7930373	Eastman Bore (PGE 11)
6142	PGE Cr Ni	323703	7956385	Lamboo (PGE 1)
6143	PGE	323647	7956464	Lamboo (PGE 2)
6150	PGE Au	323600	7956544	Lamboo (PGE 3)
6151	PGE Cr	322670	7957902	Lamboo (PGE 4)
6152	PGE Cr	324092	7957107	Lamboo (PGE 5)
6153	PGE Cr	324822	7960693	Lamboo (PGE 6)
6156	PGE Cr	325129	7960984	Lamboo (PGE 7)
6158	PGE Au Ni	321861	7955654	Lamboo (PGE 8)
6159	PGE	322574	7955332	Lamboo (PGE 9)
6160	PGE Au Ni	323002	7958203	Lamboo (PGE 10)
6161	PGE Cr	322963	7958298	Lamboo (PGE 11)
6162	PGE Cr	323585	7959315	Lamboo (PGE 12)
6163	PGE Au Ni	325563	7961674	Lamboo (PGE 13)
6164	PGE	321498	7955390	Lamboo (PGE 14)
6195	PGE Au Cr	248369	7930416	Eastman Bore (PGE 12)
6209	PGE Cr	248290	7930582	Eastman Bore (PGE 13)
6210	PGE Cr	248179	7930194	Eastman Bore (PGE 14)
6211	PGE Cr	249782	7931115	Eastman Bore (PGE 15)
6221	PGE	250108	7930178	Eastman Bore (PGE 16)
6226	PGE Cr	248068	7930202	Eastman Bore (PGE 17)
6228	PGE	248227	7930305	Eastman Bore (PGE 18)




KEY TO COMMODITY CODES

Minor commodities shown in brackets

Ag	silver	Nb	niobium
Al	aluminium	Ni	nickel
As	arsenic	P	phosphorus
Asbc	asbestos; chrysotile	Pb	lead
Au	gold	Pd	palladium
Bi	bismuth	PGE	platinum group elements
Brl	beryl	Prh	prehnite
Brt	barite	Pt	platinum
Bsbg	building stone; black granite	REE	rare earth elements
Bsdo	building stone; dolerite	Sb	antimony
Cd	cadmium	Sn	tin
Co	cobalt	Sst	sandstone
Cr	chromium	Ta	tantalum
Cu	copper	Ti	titanium
Dmd	diamond	Tlc	talc
Fe	iron	Toz	topaz
Fl	fluorite	U	uranium
Ga	gallium	V	vanadium
Grt	garnet	Vrm	vermiculite
Hf	hafnium	W	tungsten
Ilm	ilmenite	Xen	xenotime
Li	lithium	Zebr	zebra rock
Mag	magnetite	Zn	zinc
Mgs	magnesite	Zr	zirconium
Mo	molybdenum		


NUMBER	COMMODITY	EASTING	NORTHING	NAME
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PRECIOUS METAL



	Stratabound volcanic and sedimentary – volcanic-hosted sulfide			
6094	Au Ag Pb Zn	303021	7911062	Mount Carmel East 2
	Stratabound volcanic and sedimentary – sedimentary-hosted sulfide			
3170	Au	382895	8010833	Woodward Creek (Au)
5385	Ag (Bi Pb Cu Zn)	238916	7940336	Mount Ramsay 3
5802	Ag (Au Cu)	321353	7961270	Emull North 1
5803	Au	321435	7961329	Emull North 2
5804	Au	382337	7985218	Gentle Annie D
6066	Au	346552	7965208	Koongie Rockholes 2
	Vein and hydrothermal – undivided			
1069	Au Ag	401746	8062831	Wills Creek B2
1220	Ag	374726	8007941	Armanda
1275	Au Cu Ag	250586	7955646	Me No Savvy C
1315	Au Pb Ag	293263	7967595	Soda Spring 1
1318	Au Cu Ag	293554	7967712	Soda Spring 2
1322	Au	297853	7967445	Soda Spring 3
1326	Au Cu Ag	298009	7967663	Soda Spring 4
2210	Au	407083	8045708	Slinkey Hill
2211	Au Ag Pb Cu Zn	386805	8027837	Star of Kimberley
2212	Au Ag (Pb Zn)	384890	8028088	Perseverance
2214	Au Ag	385873	8028239	Moodys
2215	Au Ag (Pb Zn)	388965	8029611	Wilson's Reef
2216	Au	388743	8028368	Caledonian
2217	Au	388462	8029188	Scottish Chief
2218	Au	389062	8029188	Comet
2219	Au	389369	8029468	Lady Kimberley
2220	Au (Cu Pb)	389598	8029682	Panton Queen
2221	Au Ag Pb	389574	8029868	Lone Hand
2223	Au	374956	7971075	Palm Springs
2224	Au	377977	7973554	Golden Crown
2232	Au	378645	7974291	Faugh A Ballagh
2233	Au	376598	7973608	Lady Margaret
2234	Au (Pb)	376483	7974868	lo
2235	Au	377474	7976218	Deimos
2236	Au	375755	7970840	Mount Bradley
2237	Au (Pb As)	367721	7958932	Ruby Queen
2238	Au	367555	7959099	St Lawrence
2239	Au	368269	7959909	West and Left
2240	Au	368398	7960115	Goliath
2241	Au	368475	7960226	Ruby Queen Extended North
2242	Au	368373	7960306	Triumph
2243	Au	368761	7960377	Rising Sun
2244	Au	368475	7960393	Sunny Corner
2246	Au Pb Ag (Cu)	369733	7979606	Jubilee
2249	Au	370178	7980860	Gladstone
2250	Au Pb Ag (Cu)	370279	7981855	Lady Broome
2251	Au	370490	7982152	Jacksons Reef
2253	Au Pb Ag W Ti	363211	7966798	Darcys
2254	Au	362148	7964759	Deep Mine
2255	Au	362099	7965292	Three Shafts
2256	Au	362743	7966294	Darcys 2
2257	Au	362774	7966349	Darcys 3
2258	Au	362624	7966183	Darcys 4
2259	Au	349615	7956442	Reform
2260	Au	348440	7954347	Shane
2261	Au	348480	7954680	Sandra
2262	Au	347079	7952706	Kerry
2263	Au	347424	7952973	Big Red
2264	Au	346965	7953193	Danny
2266	Au	347041	7952051	Mays Patch
2268	Au	357065	7961400	Chinamens
2269	Au	358962	7962757	Lynas
2270	Au	355382	7960495	McPhees Creek
2271	Au	358009	7962607	New Prospect
2272	Au	357835	7962789	Sams Creek
2273	Au	357112	7962218	Camp Prospect
2275	Au	353556	7959519	Blue Bar
2276	Au	353961	7959860	Blue Bar Central
2277	Au	354660	7960257	Blue Bar North
2278	Au	353556	7959519	Blue Bar East
2279	Au	354890	7960337	New Prospect 2
2280	Au	351929	7959026	Poverty Creek Central
2281	Au	352651	7959526	Yellow Creek
2282	Au	353383	7959658	Straight Creek
2285	Au	355580	7960662	Kaolin
2286	Au	355826	7960828	Spear Gully
2287	Au Pb	324847	7936526	20 Mile
2289	Au Pb	327292	7936875	Juggler Bow

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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PRECIOUS METAL

	Vein and hydrothermal – undivided			
2291	Au	328570	7938661	Montys
2292	Au	329433	7939258	Montys Extended
2295	Au	331808	7941400	Baxters
2299	Au Cu Pb	346767	7965581	Koongie Rockholes 1
2300	Au	293783	7890373	Christmas Creek 1
2302	Au	296810	7890352	Christmas Creek 2
2305	Au (W)	294362	7888807	Christmas Creek 3
2318	Au	311635	7918414	Victoria
2321	Au Pb Ag	311643	7918485	McNeils
2322	Au	311643	7918239	Lady Hopetoun
2324	Au (Cu Pb)	310072	7914400	Irish Lass
2325	Au	309600	7914110	Erin go Bragh Southwest
2326	Au	309703	7914070	Erin go Bragh Southeast
2327	Au	309324	7915493	Hazletts
2328	Au	309012	7913229	Carol 1
2329	Au	306423	7912699	Western Lead
2332	Au (Pb)	310675	7916932	Old Mac
2333	Au	311429	7916866	Kathleen
2334	Au	312302	7918771	Ross and McIntyre
2335	Au Pb	316680	7924787	Hangmans (Au A)
2336	Au	316481	7924462	Hangmans (Au B)
2337	Au	316204	7924065	Hangmans (Au C)
2339	Au Ag (As)	348150	7988854	Lady Helen
2340	Au	326386	7963567	Nicolsons Find
2341	Au Pb Ag	326108	7962010	Nicholsons Find South
2342	Au	327465	7963113	Burnt Out
2343	Au	325200	7965704	Springvale Fault Prospect
2344	Au As Pb Ag	336288	8001740	Gnewing Bore 1
2345	Au Ag	409566	8179905	Jailhouse Creek 1
2347	Au Pb	409247	8179916	Jailhouse Creek 3
2733	Au	378705	8005972	Duffers
2736	Au	346710	7965749	Big Mac
2739	Au	371423	7965695	Halls Gully 1
2874	Au	401313	8060418	Wills Creek C2
2875	Au	408820	8032989	Black Duck Creek D
3198	Au	384633	8024658	Katy Yard 2
3214	Au	389933	8029058	Comet Hill (primary)
3326	Au	301633	7918158	Mount Dockrell 1
3327	Au	312132	7917458	Mount Dockrell 2
4284	Au	375002	7973738	Titan
4288	Au	375478	7976188	Europa
4802	Au	279563	7970678	Dead Horse 5
4940	Au	394516	8015213	Dry Creek 1
4942	Au	394794	8015715	Kangaroo Hole
5255	Au Ag (Cu)	320988	7961196	Emull West 2
5420	Au	490540	7836506	Larranganni Bluff 1
5421	Au	485725	7835527	Larranganni Bluff 2
5506	Au	349515	7965855	Whopper 1
5507	Au	349275	7965137	Whopper 2
5508	Au	345985	7962727	Whopper 3
5509	Au	344135	7961347	Whopper 4
5510	Au	343645	7960767	Whopper 5
5578	Au	396331	8016932	Dry Creek 2
5581	Au	394568	8015718	Dry Creek 3
5582	Au	394688	8015559	Dry Creek 4
5583	Au	394640	8016440	Dry Creek 5
5584	Au	394545	8017130	Dry Creek 6
5609	Au	386579	7996576	Bertha Peak
5614	Au	329258	7943060	Woodward Range
5800	Au	365657	7986343	Elvire
5809	Au	489245	7868644	Venus 1
5810	Au	488531	7868818	Venus 2
5811	Au	488579	7869367	Venus 3
5821	Au Pb Ag Cu	309611	7895749	Minneroo Pool 2
5945	Au	312941	7920196	South Dockrell 9
5963	Au	348937	7955448	Charlies Pride
5964	Au	348747	7955682	Sarah Barton
5965	Au	347805	7953616	Ginavieve
5966	Au	347607	7953298	Julie
5992	Au	381267	7981123	Poverty Gully 3
6067	Au	346830	7966042	Koongie Rockholes 3
6068	Au	347362	7966050	Koongie Rockholes 4
6069	Au	347791	7965534	Koongie Rockholes 5
6072	Au	296259	7890267	Christmas Creek 7
6074	Au	485500	7833700	Kookaburra
6075	Au	485700	7834100	Sandpiper
6076	Au	485700	7835200	Hawk
6077	Au	485800	7835950	Robin

NUMBER	COMMODITY	EASTING	NORTHING	NAME
PRECIOUS METAL				
	Vein and hydrothermal – undivided			
6078	Au	486319	7837277	Cuckoo
6079	Au	486427	7837428	Cuckoo North
6080	Au	486535	7839757	Tem South
6081	Au	487100	7840600	Tem
6082	Au	484100	7833200	Osprey
6083	Au	485000	7833300	Albatross
6084	Au	485400	7830000	Finch
6085	Au	484468	7834210	Pelican
6086	Au	484269	7834237	Gull
6087	Au	484018	7834250	Magpie
6095	Au	296707	7912399	Mount Carmel 1
6101	Au	289809	7910118	Mount Carmel 2
6107	Au	311882	7917903	Mount Dockrell 8
6119	Au Ag	348553	7988677	Moola Bulla 3
6121	Au	347989	7988900	Moola Bulla 4
6125	Au	370463	7982202	Jacksons Reef North
6126	Au	370590	7982133	Jacksons Reef East
6127	Au Pb Ag	370310	7981764	Lady Broome No. 1 North
6128	Au	370336	7981543	Old Halls Creek 1
6129	Au	370473	7981481	Old Halls Creek 2
6130	Au	370395	7981396	Old Halls Creek 3
6131	Au	370256	7981075	Black Mount
6132	Au	370144	7980933	No. 2 South Black Mount
6133	Au	370369	7981252	Mount Black North
6134	Au	369408	7980829	Old Halls Creek 4
6135	Au	370315	7980998	Old Halls Creek 5
6136	Au	370408	7981416	Old Halls Creek 6
6139	Au	330475	7940661	Head of the Garden 2
6154	Au	322602	7959934	Lamboo (Au 1)
6157	Au	321956	7955805	Lamboo (Au 2)
6165	Au	321498	7955390	Lamboo (Au 3)
6174	Au (Pb W)	310779	7915439	Native Sulphur
6176	Au	308488	7913284	Pete
6177	Au	309235	7913316	Susy
6181	Au	308616	7913197	Carol 2
6183	Au	309639	7914181	Erin go Bragh North
6184	Au	310124	7914141	Mary-anne
6193	Au	310536	7913085	Gossan Zone
6244	Au	362322	7965230	Darcys 5
6255	Au	340647	7939305	Blue Hills South
6256	Au	342883	7943334	Blue Hills North
6257	Au	342600	7951628	Golden Plains 1
6258	Au	366274	7966056	Halls Gully 7
6269	Au	375840	7972543	Mount Bradley 2
6313	Au	326910	7959983	Shiftys Reef
6314	Au	327052	7960602	Paddock Well 1
6315	Au	327124	7959697	Paddock Well 2
6316	Au	327227	7959745	Paddock Well 3
6317	Au	327322	7959753	Paddock Well 4
6318	Au	327282	7959650	Paddock Well 5
6319	Au	328084	7959824	Paddock Well 6
6320	Au	322851	7962711	White Horse
6340	Au	475268	7798193	Billiluna
6342	Au	324251	7982393	Billabong
6343	Au	352405	7959630	Yellow Creek 2
6344	Au	352866	7959638	Yellow Creek 3
6347	Au	355922	7960677	Jocks Find
6348	Au	351929	7959106	Poverty Creek North
6349	Au	351675	7958939	Poverty Creek West
6350	Au	352112	7959169	Poverty Creek East
6351	Au	355446	7969384	Teds Find
6352	Au	355287	7960431	McPhees Creek West
6353	Au	355604	7960574	McPhees Creek East
6369	Au (As)	336510	8001383	Gnewing Bore 2
6373	Au Pb Cu Zn Ag	298853	7968636	Eastern Shear 2
6381	Au Ag	347632	7988788	Moola Bulla 6
6382	Au	375258	7971512	Wolfdale
6383	Au	376417	7973107	Afghan
6384	Au	374401	7970369	Deep Extension
6385	Au	375226	7969980	Phoenix
6390	Au (Cu Pb)	204997	7951928	Mount Fairbairn 2
6403	Au	289900	7972060	Dead Horse 1
6404	Au	292507	7973416	Glidden 2
6424	Au Pb Cu Ag	311676	7917511	Crowter
6427	Au	311921	7918644	Mount Dockrell 4
6428	Au	311762	7918271	Mount Dockrell 5
6429	Au Pb Cu	310595	7915662	Gordons
6430	Au Pb Cu W	310543	7915768	Galena

NUMBER	COMMODITY	EASTING	NORTHING	NAME
PRECIOUS METAL				
	Vein and hydrothermal – undivided			
6431	Au	310000	7916191	Cheryl
6432	Au	310357	7917607	Lyn
6435	Au	309373	7915715	Jack Shaw
6436	Au	307270	7917951	Fred
6445	Au	309889	7914287	Crackerjack
6448	Au	308871	7913956	Ross
6449	Au	308448	7914061	Goddards Find
6450	Au	308407	7913810	Linda 1
6452	Au	308010	7913757	John
6453	Au	307230	7913995	Janette
6454	Au	310722	7917686	Cath
6455	Au	307600	7912209	Martins Find
6458	Au	303737	7911111	Denise
6459	Au	304174	7911852	Boothey
6460	Au (Cu)	304465	7911839	Michelle
6461	Au	304267	7911733	Chris
6462	Au	305219	7912275	Jan
6463	Au	305299	7912818	Helen
6464	Au Pb	303963	7916245	Coxes Find
6465	Au	309122	7923626	Sharon
6466	Au	482385	7799797	Coyote
6467	Au	439058	7790546	Camel
6468	Au	423448	7821746	Afghan 2
6469	Au	437106	7804651	Bandicoot
6470	Au	462845	7791642	Hutches Find
6475	Au	389717	8029820	Brockman King 1
6476	Au	376201	7972913	Brockman King 2
6477	Au	375994	7972659	Southern Cross
6478	Au	367864	7959131	No 1 North Ruby Queen
6479	Au	367364	7958885	St Lawrence Extended South
6480	Au	368023	7959504	Pyramid
6481	Au	368205	7959766	Union
6482	Au	393700	8008834	Mount Coghlan 2
6522	Au	378003	7973529	New Golden Crown
6530	Au	310191	7914443	Old Golden Dream
6534	Au	367703	7957105	Shorts Hope
6537	Au	312772	7913593	Mount Miniard
6556	Au	391835	8032257	Granite
	Regolith – alluvial to beach placers			
2290	Au	327435	7936939	Juggler Bow (eluvial)
2293	Au	330308	7940677	Head of the Garden 1
2301	Au	295624	7890183	Christmas Creek 4
2728	Au	382660	7978775	Twelve Mile
2729	Au	389409	7976550	Elvire River A
2730	Au	377833	7974858	Nuggety Gully 1
2732	Au	388292	8023458	Katy Yard 1
2761	Au	401232	8060358	Wills Creek C1
2766	Au	410533	8040258	Black Duck Creek A
2767	Au	410832	8040058	Black Duck Creek B
2768	Au	411532	8039758	Black Duck Creek C
2876	Au	400234	8021458	Saunders Creek (Au)
2984	Au	408334	8048158	Chestnut Creek
2994	Au	407433	8045458	Slinkey Hill (alluvial)
2996	Au	408534	8048658	Lower O'Driscoll Creek
2997	Au	408134	8050158	Ball Mill Creek
2998	Au	408333	8047258	O'Driscoll Creek
3159	Au	395233	8005458	Coghlan East
3194	Au	414433	8026958	Wilson Yard
3195	Au	414433	8026358	Panton-Elvire (Au)
3196	Au	414333	8028258	Hardmans Bend
3197	Au	421733	8034958	Big Bend
3228	Au	413633	8037658	Sluice Plant
3229	Au	411133	8047858	Steep Creek
3328	Au	312133	7916258	Mount Dockrell 3
3329	Au	311040	7917779	Lois
3330	Au	311216	7917193	Wattle Creek
3356	Au	405033	7988458	Elvire River B
5378	Au	388648	8006914	Ding Dong Downs (Au)
5411	Au	487582	7860956	Gardiner Range 1
5412	Au	491373	7868446	Gardiner Range 2
5416	Au	495647	7875069	Junction Hill
5417	Au	489270	7874699	Slatey Creek
5585	Au	394545	8017130	Dry Creek 7
5610	Au	366036	7967947	White Elvire River
5615	Au	351651	7961083	Koongie Park
5820	Au	293894	7891556	Christmas Creek 5
5826	Au	447496	7877644	Lewis Creek
5827	Au	446980	7880038	Lewis Creek North

PRECIOUS METAL

PRECIOUS METAL

 Regolith – alluvial to beach placers

5935	Au	311163	7920324	Poachers Creek 1
5936	Au W Sn	311591	7918990	Poachers Creek 2
5938	Au	312504	7919887	Horseshoe Creek 1
5941	Au W	312711	7920078	Horseshoe Creek 2
5942	Au Sn	312663	7920204	Horseshoe Creek 3
5943	Au W Sn	312917	7919975	Horseshoe Creek 4
5944	Au	310933	7919847	South Dockrell 8
5946	Au	313243	7920197	South Dockrell 10
5962	Au	348800	7955338	Kerry-Ann
5984	Au	369493	7965480	Halls Gully 3
5985	Au	369121	7964900	Halls Gully 4
5986	Au	368890	7964917	Halls Gully 5
5987	Au	368819	7964766	Halls Gully 6
5990	Au	380609	7979781	Poverty Gully 1
5991	Au	380993	7980831	Poverty Gully 2
5993	Au	381452	7980437	Saunders Creek 2
5994	Au	382183	7981073	Northern Creek
5999	Au	357735	7963516	Red Gully
6000	Au	357759	7962912	Sam Hazlet Creek
6007	Au	357409	7962761	Chinaman Creek
6011	Au	355084	7961444	McPhee Creek 2
6013	Au	355393	7961047	McPhee Creek 3
6014	Au	354965	7961578	Blue Bar Creek
6024	Au	354483	7960491	Ruby Plains Alluvials 1
6025	Au	355039	7961118	Ruby Plains Alluvials 2
6026	Au	355491	7960880	Ruby Plains Alluvials 3
6027	Au	378395	7973917	Nuggety Gully 2
6034	Au	376298	7976028	Upper Spring Creek
6035	Au	376417	7978145	Middle Spring Creek
6036	Au	376047	7979415	No Name Creek
6037	Au	378243	7976386	Eastern Twin Creek
6040	Au	377674	7976928	Western Twin Creek
6046	Au	382238	7972073	Donkey Creek 2
6049	Au	375078	7971602	Butchers Creek
6053	Au	375885	7971761	Brockman Creek
6057	Au	371597	7961718	Ruby Creek
6070	Au	346536	7965685	Koongie Rockholes 6
6091	Au	297003	7912401	8 oz Creek
6102	Au	289423	7910234	Janet Creek
6103	Au	307153	7934694	Syenite Camp
6104	Au	310472	7919075	Lee 1
6105	Au	310433	7918670	Lee 2
6106	Au	309202	7917773	Lee 3
6116	Au W	381908	7984255	Gentle Annie E
6137	Au	369654	7980155	Jubilee Creek
6138	Au	349289	7955790	East Mary River
6169	Au	349281	7956191	Mary River 1
6171	Au	318320	7923807	Fly Creek
6173	Au	311933	7914998	Bucket Creek 1
6185	Au	310648	7915860	Top of Bucket
6186	Au	310520	7914927	Bucket Creek South
6188	Au	310124	7914268	O'Romneys
6192	Au	311870	7914007	Upper Willy Willy Creek
6231	Au	376316	7974895	lo (alluvials)
6252	Au	339827	7953857	Mary River 2
6253	Au	347354	7954075	Thompson Creek
6254	Au	344629	7951866	Six Mile Creek
6266	Au	345216	7950877	Six Mile Creek 2
6276	Au	332118	7941015	Garden Creek 4
6287	Au	350336	7957624	Specimen Creek
6288	Au	350635	7956346	25 oz Creek
6289	Au	346780	7955830	Hopeful Creek
6290	Au	346582	7955305	Little Gold Creek

 Regolith – residual to eluvial placers


2296	Au	331840	7941511	Baxters (eluvial)
2737	Au	371287	7965258	Halls Gully 2
3213	Au	389933	8029058	Comet Hill (eluvial)
6071	Au	295624	7890183	Christmas Creek 6
6189	Au	309195	7913157	No-name Creek
6299	Au	349479	7956861	Bills Patch
6300	Au	346320	7953766	Car Body Creek
6301	Au	346090	7951448	Susan Peter Creek
6302	Au	346852	7951622	Susan Peter
6305	Au	348730	7954000	Fine Gold Creek
6308	Au	346129	7955782	Mary River 3
6310	Au	348902	7957732	Dam Creek
6311	Au	344346	7951152	Golden Plains 2
6336	Au	347939	7956576	Mary River 4

PRECIOUS METAL

PRECIOUS METAL

 Regolith – residual to eluvial placers

6345	Au	354065	7959797	Blue Bar West
6346	Au	354390	7960162	Little Blue Bar Creek
6354	Au	358549	7962574	Golden Fleece - Argonaut
6400	Au	352140	7958212	Poverty Creek 2
6425	Au	311175	7916493	1972
6426	Au	310913	7916297	The Patch
6433	Au	308294	7920311	Kerry 2
6434	Au	309135	7913320	Jackie
6442	Au	308659	7915001	Leonora
6443	Au	308897	7914921	Leonora East
6444	Au	308989	7915212	Shirleys
6446	Au	309730	7915953	Josy
6447	Au	310101	7915424	Sara
6451	Au	308236	7913744	Linda 2
6456	Au	305907	7913241	Alison
6457	Au	305365	7911535	Leo
6486	Au	456130	8194473	Prospect Creek
6514	Au Ag	399835	8006757	Number 1
6546	Au	346651	7951818	New Look
6730	Au	377433	7975458	Rosies Flat

 Regolith – residual and supergene

2222	Au	394489	8016138	Dry Creek (alluvial)
2252	Au	369847	7981472	Olympio (alluvial)
2265	Au	348830	7954624	Ann-Marie
2288	Au	325101	7936287	20 Mile (alluvial)

STEEL INDUSTRY METAL

 Orthomagmatic mafic and ultramafic – layered mafic

5	Ni Cu Co	385047	8074259	Spring Complex 1
7	Cu Ni Co	385699	8075036	Spring Complex 2
10	Ni Cu Co	384286	8072719	Spring Complex 3
13	Cr	380652	8042542	Togo 5B
639	Ni Cu	393034	8047578	Cabernet
640	Ni (Cu)	361734	8023558	West Robin Soak
641	Ni Cu Co	413316	8087584	Corkwood 1
642	Ni Cu Pt	390673	8086758	Keller Creek 1
652	Cu Ni	387904	8068941	Ord Crossing (Ni-Cu)
653	Ni Cu Co	396283	8081558	Sally Malay
659	Ni Cu	423236	8113365	Turkey Creek (Ni-Cu)
666	Cu Ni Au	387283	8054608	Melon Patch 2
667	Cu Ni Ag	429277	8134890	Bow River (Ni-Cu)
762	Ni Cu	403810	8072068	Bulldust Flat
775	Ni Cu	400952	8067949	Wills Creek (Cu-Ni) Alpha
810	Cu Ni	422622	8095686	Osmond Valley (Ni-Cu)
812	Cr (PGE)	361033	8052908	Toby South
817	Cr	376883	8016958	Big Ben
831	Cr	321821	7956163	Lamboo (Cr 1)
838	Ti	389201	8069999	Ord Crossing (Ti)
842	Ti	398632	8090158	Frog Hollow North Lens
843	Ti	398533	8087558	Frog Hollow South Lens
2868	Ni Cu	392690	8046983	Shiraz
2869	Ni Cu	393933	8049682	Malbec
2877	Ni Cu PGE	366733	8030458	Springvale A
3554	Cr (Cu)	391933	8048658	White Rock Well 2
3560	Ni Cu PGE	389632	8049858	White Rock Well 3
4205	Ni Cu	429285	8143523	Salt Lick Creek (Ni-Cu)
4703	Ni Cu	385632	8073858	Coolumbuloo Yard
4739	Ni Cu	380929	8064252	McKenzie Spring 4
4740	Ni Cu	380726	8063938	McKenzie Spring 5
4743	Ni Cu	380357	8063883	McKenzie Spring 6
4745	Ni Cu	380118	8063497	McKenzie Spring 7
4747	Ni Cu	380413	8066665	McKenzie Spring 9
4758	Ni Cu	361033	8038458	Springvale 7
4760	Ni Cu Co	398533	8078158	Tickalara Bore 3
4762	Ni Cu	388282	8054908	Melon Patch 8
5104	Cr (Zn PGE)	322376	7957117	Lamboo (Cr 2)
5105	Cr	325695	7961907	Lamboo (Cr 3)
5110	Cr	325443	7961521	Lamboo (Cr 4)
5113	Cr (PGE)	326032	7961819	Lamboo (Cr 5)
5116	Cr	323738	7959630	Lamboo (Cr 6)
5117	Cr (PGE)	323340	7959061	Lamboo (Cr 7)
5126	Cr	323116	7958294	Lamboo (Cr 8)
5128	Cr	322151	7955529	Lamboo (Cr 9)
5129	Cr	322335	7955714	Lamboo (Cr 10)
5132	Cr	322548	7955833	Lamboo (Cr 11)
5135	Cr	322706	7955780	Lamboo (Cr 12)
5137	Cr	323593	7955410	Lamboo (Cr 13)

NUMBER COMMODITY EASTING NORTHING NAME

STEEL INDUSTRY METAL



Orthomagmatic mafic and ultramafic – layered mafic

5140	Cr	322032	7956614	Lamboo (Cr 14)
5141	Cr	323513	7956402	Lamboo (Cr 15)
5143	Cr	323752	7956574	Lamboo (Cr 16)
5152	Cr	323952	7956826	Lamboo (Cr 17)
5159	Cr	324188	7956985	Lamboo (Cr 18)
5161	Cr	325657	7959140	Lamboo (Cr 19)
5162	Cr	324360	7957341	Lamboo (Cr 20)
5197	Ni Cu Ag Au (Pt)	395610	8075629	Tickalara Bore 4
5198	Ti	385801	8066004	Ord Crossing (Ti) 2
5199	Ti	387599	8068729	Ord Crossing (Ti) 3
5200	Ti	385788	8065448	Ord Crossing (Ti) 4
5202	Ni Cu	395733	8077558	Tickalara Bore 5
5203	Ni Cu Au Co	396205	8076515	Tickalara Bore 6
5262	Ni Cu	382615	8065371	McKenzies Spring 11
5265	Ni	380833	8066158	Donkey Valley 2
5266	Ni Cu Co	382583	8066008	Donovans Reef
5272	Ni	373943	8032931	Panton 2
5273	Cr	374155	8031821	Panton 3
5285	Ni Cu	379200	8046268	Mabel Hill (Ni-Cu) 1
5286	Ni Co	378821	8046247	Mabel Hill (Ni-Cu) 2
5287	Ni Cu	429971	8135214	Bow River G
5288	Ni Cu	363141	8049393	Billymac Yard D
5292	Ni Cu	381676	8064465	McKenzies Spring 14
5301	Ni Cu	387525	8068417	McKenzies East 1
5302	Ni Co Cu	386595	8067893	McKenzies East 2
5304	Ni Cu	384405	8066717	McKenzies East 4
5305	Ni Cu	383795	8065781	McKenzies East 5
5306	Ni Cu	383351	8066220	McKenzies East 6
5310	Ni Cu Co	396018	8078899	Mabel Downs 3
5311	Ni Cu	390858	8086797	Keller Creek 2
5312	Ni Cu	391123	8086718	Keller Creek 3
5313	Ni Cu	390448	8086612	Keller Creek 4
5314	Ni Cu	390792	8086096	Keller Creek 5
5372	Ni	376798	8062690	McKenzie No. 1
5514	Ni (Cu)	413093	8087584	Corkwood 2
5594	Cu Ni Co Cr	409235	8080357	Sally Malay Bore 2
5595	Cu Ni Co	409035	8079357	Sally Malay Bore 3
5596	Cu Ni Co	409535	8082757	Sally Malay Bore 4
5612	Ni Cu Co	396226	8085791	Eskimo Knoll
5787	Cr (PGE)	324781	7958061	Lamboo (Cr 21)
5788	Cr (PGE)	326294	7954201	Lamboo (Cr 22)
5789	Cr (PGE)	325144	7952627	Lamboo (Cr 23)
5790	Cr	324072	7951237	Lamboo (Cr 24)
5791	Cr	323463	7950550	Lamboo (Cr 25)
5792	Cr	323040	7950100	Lamboo (Cr 26)
6140	Cr (PGE)	326767	7955628	Lamboo (Cr 27)



Orthomagmatic mafic and ultramafic – undivided

763	Ni Cu Co	402983	8068308	Cattle Creek
5204	Co	375882	8062487	McKenzies Spring 10



Pegmatitic

961	W	320429	7915784	Junda 1
	Skarn			
887	W (Sn)	314987	7907639	Mount Heartbreak A
930	W Sn Ta Au	316639	7903935	Frog Creek
3299	W (Sn)	368533	8001158	Central Skarn
3300	W	365833	8000358	Southern Skarn



Stratabound volcanic and sedimentary – undivided

934	W	376132	7987481	Sophie Downs South
941	W	377620	7997669	Sophie Downs West
942	W	319739	7909041	Taylor Lookout 1
944	W	319406	7909001	Taylor Lookout 3
5046	W	395275	7996450	Castle Creek 11
5833	W	394320	7997069	Castle Creek 18
5834	W	394273	7996990	Castle Creek 19
5837	W	393979	7994886	Castle Creek 22
5839	W	394177	7994814	Castle Creek 23
5846	W	393153	7994008	Glory Rock Hole 8
5872	W	394799	7997565	Castle Creek 33
5911	W	319334	7908700	Taylor Lookout 10
5912	W	320072	7909406	Taylor Lookout 11
6366	Mo Ag (As)	365585	7997607	Tiger 2



Stratabound sedimentary – clastic-hosted

849	W	392420	7991055	Rock Pile 1
868	W	393105	7994117	Glory Rock Hole 1
873	W	393693	7996133	Snake Gully North

NUMBER COMMODITY EASTING NORTHING NAME

STEEL INDUSTRY METAL



Stratabound sedimentary – clastic-hosted

876	W Mo	396722	8000397	Late Night
884	W Sn Ta	311564	7920544	South Dockrell 1
2633	W	392034	7991938	Southwest 1
5036	W	397251	8002678	Castle Creek 1
5037	W	397132	8002758	Castle Creek 2
5038	W	397064	8001247	Castle Creek 3
5039	W	396071	8001548	Castle Creek 4
5040	W	395307	7999888	Castle Creek 5
5041	W	395204	7999617	Castle Creek 6
5042	W	394680	7998426	Castle Creek 7
5043	W	395921	8001342	Castle Creek 8
5047	W	395140	7996172	Castle Creek 12
5048	W	394990	7995918	Castle Creek 13
5049	W	394862	7995386	Castle Creek 14
5831	W	393574	7995918	Snake Gully South
5836	W	394193	7995267	Castle Creek 21
5840	W	393010	7994016	Glory Rock Hole 2
5841	W	393066	7994007	Glory Rock Hole 3
5842	W	393066	7993968	Glory Rock Hole 4
5843	W	393018	7993904	Glory Rock Hole 5
5844	W	392955	7993817	Glory Rock Hole 6
5845	W	392931	7993762	Glory Rock Hole 7
5847	W	391981	7991814	Southwest 2
5848	W	392076	7991711	Southwest 3
5849	W	392005	7991711	Southwest 4
5850	W	392608	7991521	Castle Creek 24
5851	W	392457	7991385	Castle Creek 25
5852	W	392330	7991401	Castle Creek 26
5853	W	392449	7991259	Castle Creek 27
5854	W	392751	7991402	Castle Creek 28
5855	W	392711	7991243	Castle Creek 29
5856	W	392539	7991118	Castle Creek 30
5857	W	392532	7991071	Castle Creek 31
5858	W	392285	7991039	Rock Pile 2
5859	W	392531	7990952	Rock Pile 3
5860	W	392515	7990856	Rock Pile 4
5861	W	392396	7990746	Rock Pile 5
5862	W Ag (Cu Pb)	392293	7990603	Rock Pile 6
5863	W	392214	7990578	Rock Pile 7
5864	W	392174	7990721	Rock Pile 8
5865	W	392111	7990547	Rock Pile 9
5866	W	392111	7990444	Rock Pile 10
5867	W	392055	7990865	Rock Pile 11
5868	W	392373	7990960	Rock Pile 12
5869	W	392341	7990904	Rock Pile 13
5870	W	392349	7990840	Rock Pile 14
5873	W	394379	7997490	Castle Creek 34
5874	W	395418	7996767	Castle Creek 35
5875	W	394211	7997402	Castle Creek 36
5876	W	394167	7997236	Castle Creek 37
5878	W	394846	7998800	Castle Creek 38
5879	W	394910	7998974	Castle Creek 39
5880	W	394950	7999069	Castle Creek 40
5881	W	395005	7999188	Castle Creek 41
5882	W	394934	7999125	Castle Creek 42
5883	W	395014	7999252	Castle Creek 43
5884	W	395339	8000037	Castle Creek 44
5885	W	395444	8000342	Castle Creek 45
5886	W	395476	8000516	Castle Creek 46
5887	W	395540	8000660	Castle Creek 47
5888	W	395603	8000842	Castle Creek 48
5889	W	397016	8002073	Castle Creek 49
5890	W	397056	8001167	Castle Creek 50
5891	W	397000	8001000	Castle Creek 51
5926	W	311676	7920205	South Dockrell 2
5927	W	312087	7920344	South Dockrell 3
5928	W	312286	7920450	South Dockrell 4
5929	W	312495	7920553	South Dockrell 5
5930	W	313020	7921087	South Dockrell 6
5931	W	312584	7920193	South Dockrell 7



Stratabound sedimentary – undivided

5597	V Mo	411435	8084757	Sally Malay Bore 5
5598	V Cr Mo	411535	8085357	Sally Malay Bore 6



Vein and hydrothermal – undivided

882	W	314487	7923031	North Dockrell
945	W	319500	7909249	Taylor Lookout 4
959	W (Sn)	308356	7904668	Balara

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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STEEL INDUSTRY METAL



Vein and hydrothermal – undivided

5044	W	392679	7992521	Castle Creek 9
5045	W	392743	7992481	Castle Creek 10
5169	Mo	376023	8046164	Mabel Hill
5187	Ni Cu	354781	8032327	Springvale 8
5188	Ni Cu	356726	8033385	Springvale 9
5189	Ni Cu	352557	8032883	Springvale 10
5191	Ni Cu	356223	8035714	Springvale 12
5835	W	393915	7996180	Castle Creek 20
5871	W	394640	7997474	Castle Creek 32
5892	W	397024	8001278	Castle Creek 52
5900	W (Sn)	349424	7951099	Ruby Plains 1
5901	W (Sn)	349350	7950924	Ruby Plains 2
5902	W (Sn)	350418	7952851	Ruby Plains 3
5910	W	319135	7908874	Taylor Lookout 9
6364	W (As Bi)	459635	7812157	Hermes



Regolith – alluvial to beach placers

844	Ilm V	391503	8206338	Speewah Valley (Ti)
968	W	364933	8082658	Foal Creek
2873	Ilm Mag	474333	8150558	Ord (HM)
3489	W	367083	8069908	Ord River Tributary 1
3490	W	367633	8072658	Ord River Tributary 2
3492	W	384233	8058308	Reedy Creek
3494	W	355933	8012408	Little Pantan (W)
3496	W	368433	8062458	Sandy Creek Tributaries 1
3499	W	359733	8058058	Sandy Creek Tributaries 2
5909	W	320618	7929931	Hangmans (W)



Regolith – residual and supergene

5570	Ni (Cr)	323127	7980773	Moola Bulla 2
6155	Ni	324850	7960721	Lamboo (Ni 1)
6166	Ni	325387	7961260	Lamboo (Ni 2)
6167	Ni	324769	7960716	Lamboo (Ni 3)



Regolith – residual to alluvial

933	W	382483	7994308	Sophie Downs North
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SPECIALITY METAL



Carbonatite and alkaline igneous intrusions

1119	REE P Zr Nb Vrm	306820	7867004	Cummins Range (REE)
1122	REE	377371	8041350	Copperhead
5961	REE Nb P	299571	7864392	Cummins Range (REE 2)



Pegmatitic

926	Sn	317617	7909540	Mount Heartbreak F
969	Sn	323546	7917209	Mount Cross (Sn 1)
976	Sn Ta Nb	309915	7895948	Minneroo Pool 1
5754	Sn Nb Ta Li Brl	314765	7907479	Columbian Creek
5914	Ta	320160	7917905	Mount Cross 4
6380	Sn	323289	7917633	Junda 3



Skarn

972	Sn	323973	7917638	Mount Cross (Sn 2)
5748	Sn Ag	319408	7908764	Taylor Lookout 7



Stratabound volcanic and sedimentary – volcanic

1102	Zr Y Nb Ta Hf Ga	371853	7974444	Brockman 1
	REE Al Fl			
1103	Nb Zr Fl	379537	7989043	Sophie Downs 1
5380	Nb	389176	8008710	Ding Dong Downs (Nb)
5916	Nb Zr Fl REE	379727	7989456	Sophie Downs 2
6111	Zr REE Fl Nb	371374	7973752	Brockman 2
6112	REE Zr Nb Fl	370215	7972347	Brockman 3
6113	Nb Sn REE Fl	370924	7973247	Brockman 4



Stratabound volcanic and sedimentary – undivided

5750	Sn Ag	318551	7908391	Taylor Lookout 8
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Stratabound sedimentary – clastic-hosted

5315	REE	416033	8088558	Junior
5316	REE U	498449	7813401	Killi Killi 1
5317	REE U	486910	7818835	Killi Killi West



Vein and hydrothermal – undivided

1105	REE (U)	417730	8087637	John Galt (REE 1)
5318	REE (U)	417929	8087815	John Galt (REE 2)
5319	REE	418206	8088358	John Galt (REE 3)
5673	Bi (Au)	308963	7874679	Shiddi Creek
6141	REE (P U)	487123	7870098	Whites Beach
6260	REE	491828	7909971	Browns Range Dome

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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SPECIALITY METAL



Regolith – alluvial to beach placers

890	Sn Ta	314832	7907758	Mount Heartbreak B
891	Sn Ta	313933	7907908	Mount Heartbreak C
895	Sn Ta	311832	7905408	Mount Heartbreak D
898	Sn Ta	314883	7906058	Mount Heartbreak E
965	Ta Nb Sn W	320124	7917762	Junda 2
966	Sn	374633	8130158	Neil Creek 1
3005	REE	415032	8086158	Corkwood Yard
3376	Sn Ta Nb	351933	8097958	Mount Remarkable 7
3378	Sn	352333	8111408	Mount Remarkable 8
3488	Sn	374933	8130308	Neil Creek 2
5822	Sn Ta Nb Au	310063	7895964	Minneroo Pool 3
5915	Ta	318770	7917585	Junda 4
6379	Sn	323075	7917109	Mount Cross 3



Regolith – residual to elluvial

3156	Sn	390033	8011158	Saunders Creek (Sn)
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BASE METAL



Orthomagmatic mafic and ultramafic – layered mafic

661	Cu Ni	382019	8065031	McKenzie Spring 1
663	Cu Ni PGE	387083	8053408	Melon Patch 1
771	Cu (Ni)	389183	8051908	White Rock Well West
806	Cu Ni	364085	8050038	Billymac Yard A
1037	Cu	362935	8049627	Billymac Yard B
2719	Cu Ni Au	401983	8062808	Wills Creek D
2878	Zn (Ni Cu Cr)	366133	8028958	Springvale B
3511	Cu (Cr)	379733	8042858	Mini M (Cu)
4737	Cu Ni	381647	8064749	McKenzie Spring 2
4738	Cu Ni	381205	8064307	McKenzie Spring 3
4746	Cu Ni	379842	8063404	McKenzie Spring 8
4753	Cu	361533	8036658	Springvale 4
4756	Cu	360432	8035758	Springvale 5
4757	Cu	361033	8033864	Springvale 6
4759	Cu Ni	398833	8078258	Tickalara Bore 2
4952	Cu	427981	8134070	Bow River A
4953	Cu	427742	8134427	Bow River B
4954	Cu	427610	8134533	Bow River C
4955	Cu	427465	8134745	Bow River D
4956	Cu	428166	8134533	Bow River E
4957	Cu	430441	8135816	Bow River F
4964	Cu	250700	7930518	Eastman (Cu) C
5254	Zn Cu Pb Ag	320984	7961025	Emull West 1
5263	Cu	381041	8066336	McKenzie Spring 12
5264	Cu Ni	380253	8066088	Donkey Valley 1; Survey One; Su1
5268	Cu Ni	386733	8074458	Coolumboloo Yard 2
5291	Cu Ni	379426	8063063	McKenzie Spring 13
5293	Cu Ni	396799	8075735	Tickalara Bore 8
5303	Cu	386135	8067217	McKenzie East 3
5368	Cu Ni Co	322573	7964800	Lamboo (Cu-Ni)
5370	Cu	386569	8048568	MI-2b
5566	Cu (Ni)	356397	8071431	Paynes Gossan East
5571	Cu Ni	413117	8087770	Corkwood 3
6213	Cu PGE	250100	7930297	Eastman Bore (Cu 4)
6218	Cu	250068	7930058	Eastman Bore (Cu 5)







Orthomagmatic mafic and ultramafic – undivided



14	Cu Ni	389233	8038958	Eileen Bore 2
764	Ni Cu PGE	389583	8040158	Eileen Bore
766	Cu	387714	8047239	McIntosh A
1147	Zn Cu Pb Ag Au	321408	7961056	Emull
	(Cd Bi)			
2483	Cu	413396	8074317	Fish Hole Waterhole
2485	Cu	418946	8088477	Osmond Range West
6312	Cu (Au Ag)	230412	7939006	Louisa Downs 3



Disseminated and stockwork in plutonic intrusions


1087	Cu Ag	425354	8109657	Killarney 1
1093	Cu	425235	8111162	McHales
1098	Cu Mo Ag	338700	7958663	Mount Angelo Porphyry
5297	Cu	425420	8109949	Killarney 2
5298	Cu	425558	8109869	Killarney 3
5299	Cu	425420	8109446	Killarney 4
5300	Cu	425062	8109499	Killarney 5
	Skarn			
1050	Cu	381751	8051305	Melon Patch Well A
1054	Cu	383233	8056088	Melon Patch Well B
1055	Cu	383733	8054298	Melon Patch Well C
1063	Cu Zn Ag	400383	8066258	Wills Creek B1
1071	Cu	397933	8053058	Wills Creek B4


NUMBER	COMMODITY	EASTING	NORTHING	NAME
BASE METAL				
	Skarn			
1074	Zn Ag	397539	8052431	Wills Creek B5
1078	Cu Ag	397133	8051508	Wills Creek B6
1085	Cu Zn Ag	392188	8049536	White Rock Well
3298	Zn (Pb)	369133	8001158	BB Hill
3306	Cu	398733	8063258	Sally Downs A
3307	Zn	392333	8051358	Sally Downs B
5032	W	393833	8057958	Sally Downs D
5290	Cu	393160	8061216	Wills Creek B7
5530	Zn	430150	8097485	Osmond Valley (Zn 7)
5743	Cu Ag Mo	320075	7909328	Skarn A
	Stratabound volcanic and sedimentary – volcanic-hosted sulfide			
1125	Cu Zn Pb Ag (Au)	340032	7968758	Sandiego (sulphide)
1133	Zn Cu Pb	346223	7974088	Onedin
1135	Zn Pb Cu	346373	7976196	Gosford
1136	Zn Pb Cu	343893	7969288	Atlantis
1167	Cu Ag (Au)	357931	7985518	Golf Course 1
1192	Zn Pb	377528	7997683	Twin Hills 1
5074	Zn Ag	368964	7964212	Ruby Queen - Brockman 2
5377	Zn Ag	389277	8010435	Palm Creek
5472	Cu	358447	7986020	Golf Course 3
5480	Cu (Zn)	342558	7963270	Angelo Valley 2
5572	Zn Pb Cu	389605	8010286	Dry Creek South
5689	Cu	241022	7928204	Eastman Bore East
5696	Zn Pb Cu	249008	7931300	Bullock Bore 10
5728	Zn Pb	377298	7997302	Twin Hills 2
5745	Cu Ag	318441	7908351	Taylor Lookout 6
6279	Pb Zn	350139	7978313	Mother O'Neil
	Stratabound volcanic and sedimentary – sedimentary-hosted sulfide			
1169	Ag Pb Zn	374346	7999969	Duffer Range 1
1180	Cu Zn Pb Au Ag	379320	8001317	Ilmars Central
1188	Zn Pb Cu	377972	8000047	Little Mount Isa
1197	Cu Zn Pb (Sn)	381033	7986508	Gentle Annie A
1202	Cu Pb Zn Ag (Sn)	380773	7985376	Gentle Annie B
1206	Cu Pb Zn Ag Au	389900	8002397	Bullman
1252	Cu (Zn)	411387	8081329	Tickalara
2531	Cu	326322	7974741	Ten Minute (Cu 3)
2871	Zn Pb Cu	342433	7997858	Five Mile B
4805	Zn Pb Cu Ag	379773	8002115	Ilmars North
4812	Zn Cu Ag	379130	8000861	Ilmars Range
4817	Zn Cu	377805	7998517	Discovery Hill
4818	Pb	378067	7998092	Quartz blow
4820	Cu Pb	379591	8002621	Ilmars Syncline
4882	Zn	389689	8005720	Sophie Downs (Zn) 1
4918	Zn	384992	7998378	Sophie Downs (Zn) 3
4923	Zn	382637	7993563	Sophie Downs (Zn) 4
4924	Zn	382439	7992968	Sophie Downs (Zn) 5
4946	Zn	394503	8013096	Panton River (gossan) 1
4947	Zn	394026	8012646	Panton River (gossan) 2
5017	Cu Ag	316617	7905635	Cummins Range (Bm) D
5019	Ag Cu	320609	7909662	Cummins Range (Bm) E
5035	Ag Cu	395254	8001343	Cow Creek Gossan
5088	Cu	380922	7986178	Gentle Annie C
5093	Zn	378474	8001198	Gossan Z
5095	Cu Pb Zn	380036	8002904	Ilmars North 2
5236	Cu Ag	305639	7917006	Black Bishop Hill B
5326	Cu	326369	7974720	Ten Minute (Cu 4)
5327	Cu	326544	7974265	Ten Minute (Cu) 5
5369	Zn	383935	8006357	EM35
5381	Cu	324970	7959199	New Homestead
5384	Pb	239687	7942464	Mount Ramsay 2
5387	Zn Pb	239861	7939267	Mount Ramsay 15A
5471	Zn Pb	358302	7985762	Golf Course 2
5786	Cu	237990	7930729	Landrigan Creek 2
6278	Pb Zn	353196	7980655	Weldens Creek
	Stratabound volcanic and sedimentary – undivided			
943	Cu Mo Ag W	319112	7908652	Taylor Lookout 2
946	Cu Mo Ag	318067	7907899	Taylor Lookout 5
982	Cu	319495	7908891	Cummins Range (Bm) A
988	Cu	317912	7907519	Cummins Range (Bm) B
997	Cu Pb Zn Ag	318308	7908203	Cummins Range (Bm) C
1134	Cu Pb Zn	338073	7965613	Hanging Tree
1137	Cu Ag (Zn)	340478	7960183	Mount Angelo North
1138	Cu Pb Zn Ag	241208	7928301	Eastman Bore (Cu-Pb-Zn)
1139	Cu Zn Pb Ag Au Mo	238659	7929724	Landrigan Creek
1140	Cu Zn Pb	249480	7931895	Bullock Bore 1
1144	Cu Zn Pb Au Bi	248964	7931828	Bullock Bore 6


NUMBER	COMMODITY	EASTING	NORTHING	NAME
BASE METAL				
	Stratabound volcanic and sedimentary – undivided			
1160	Zn Cu Ag Pb	320582	7961758	Location 5
1178	Cu Zn Ag Pb (Au)	328362	7964407	Dusty Bore 1
1214	Zn Cu Pb Ag	244786	7929667	Daves Gossan
1215	Cu Pb Zn Ag	247441	7931007	Eastman Yard East 1
1271	Cu	383632	7993158	Brim Creek
1300	Cu	227456	7943354	Morgans Grave
2429	Cu	388533	7994758	Ding Dong Downs (Cu)
2433	Cu	342467	8127378	Carson Volcanics 4
2434	Cu	344519	8129853	Carson Volcanics 5
2435	Cu	349944	8136808	Carson Volcanics 6
2436	Cu	355948	8146225	Carson Volcanics 7
2438	Cu	356243	8146535	Carson Volcanics 8
2440	Cu	358449	8148701	Carson Volcanics 9
2441	Cu	364447	8160113	Carson Volcanics 10
2442	Cu	364591	8160728	Carson Volcanics 11
2443	Cu	367785	8165352	Carson Volcanics 12
2445	Cu	369424	8169979	Carson Volcanics 13
2452	Cu	369507	8180737	Carson Volcanics 15
2463	Cu	369981	8175822	Carson Volcanics 16
2465	Cu	354326	8145292	Carson Volcanics 17
2466	Cu	347456	8132948	Carson Volcanics 20
2467	Cu	348626	8134801	Carson Volcanics 21
2468	Cu	355967	8146459	Carson Volcanics 22
2469	Cu	354470	8145908	Carson Volcanics 23
2471	Cu	354470	8145908	Carson Volcanics 24
2476	Cu	354470	8145908	Carson Volcanics 26
2477	Cu	356247	8145920	Carson Volcanics 27
2479	Cu	358011	8147776	Carson Volcanics 28
2480	Cu	354470	8145908	Carson Volcanics 33
2481	Cu	354470	8145908	Carson Volcanics 35
2487	Cu	433459	8216415	Dunham Crossing
2488	Cu	485900	8181434	Behn Gorge
2498	Cu	432317	8205091	Dunham Station
2499	Cu	487700	8144562	Wild Dog
2500	Cu	496575	8177751	Rosewood Wall A
2502	Cu	480547	8134409	Spring Hill
2505	Cu	416875	8026351	Panton-Elvire (Cu)
2517	Cu Ag	438133	8095008	Paradise Creek (Cu B)
2522	Cu	459983	8017014	Cuddys Find
2529	Cu	326592	7974868	Ten minute (Cu 1)
2872	Zn	343633	7998958	Five Mile C
4831	Cu Ag Zn	342333	7962958	Angelo Valley 1
4988	Pb Cu Zn	240331	7932460	Margaret River Gossan 1
4989	Pb Zn Ag (Au)	242312	7932534	Margaret River Gossan 2
5051	Cu Ag	396311	8001105	Castle Creek 15
5052	Zn Cu Ag	396616	8000522	Castle Creek 16
5398	Zn (Pb)	244921	7930389	Eastman Yard 1
5399	Zn (Pb)	244246	7929888	Eastman Yard 2
5400	Cu	244573	7929792	Eastman Yard 3
5422	Cu	484771	7946181	Mount Wittenoom
5501	Pb Ag Cu Zn	342610	7973361	Gossan Area 5
5502	Cu (Zn)	336970	7960117	Blue Hills
5778	Ag Cu Pb Zn	244528	7930144	Eastman Yard 4
5780	Cu Ag	244480	7928877	Eastman Yard 5
5781	Cu Ag	244626	7928493	Eastman Yard 6
5782	Pb Zn Ag	244519	7930434	Eastman Yard 7
5783	Pb	246967	7930806	Eastman Yard East 2
5784	Pb Zn	242121	7930523	Eastman Yard 8
5785	Cu	241382	7930515	Eastman Yard 9
5959	Ag Cu Pb	229272	7943454	Morgans Grave 2
6277	Cu	428735	7983057	Coolbah Creek
6324	Cu	328219	7972643	Dim Whiddy (Cu1)
6326	Cu	327822	7972392	Dim Whiddy (Cu2)
6327	Pb	327954	7972564	Dim Whiddy (Pb)
6356	Cu Zn Ag (Au)	340200	7960143	Mount Angelo North 2
6367	Ag Zn	365655	7997067	Tiger 3
6495	Cu	429307	8198672	Cassidys
6496	Cu	417380	8198987	Dunham River (Cu)
	Stratabound sedimentary – carbonate-hosted			
1045	Cu W Au (Sn)	400433	8066258	Alice Downs
2507	Pb	431034	8096058	Paradise Creek (Pb)
2518	Pb	440501	8089634	Trend 1
2519	Cu	448721	8105783	Osmond Range
2523	Cu	497820	8178523	Rosewood Wall B
4489	Cu	344481	8135078	Teronis 1
4506	Cu	359833	8159158	Teronis 2
4510	Cu	360633	8160958	Teronis 3
5030	Cu (Mo)	390533	8053258	Sally Downs C


NUMBER	COMMODITY	EASTING	NORTHING	NAME
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BASE METAL

	Stratabound sedimentary – carbonate-hosted			
5033	Cu Ag	391061	8051128	Alice Downs 5
5193	Zn	399766	8080313	Fletcher Creek
5373	Cu Zn	437722	8087702	Trend 2
5374	Zn	441413	8091254	Trend 3
5375	Zn	441803	8091499	Trend 4
5437	Zn Cu Pb Ag	189200	7984995	Horse Spring Range 1
5440	Pb Zn	189742	7987523	Horse Spring Range 2
5443	Pb Zn	182982	7985155	Findlay Hill 2
5461	Zn	189398	7985511	Horse Spring Range 3
5462	Zn Pb	219635	7901507	Black Spur 1
5466	Zn Pb	219485	7903207	Black Spur 2
5522	Cu	424196	8093992	Osmond Valley (Cu 1)
5523	Cu	423787	8093754	Osmond Valley (Cu 2)
5525	Cu	426247	8095090	Osmond Valley (Cu 3)
5636	Zn Pb Ag Cu	186932	7925440	Christmas Creek Reef 5
5637	Zn Pb	192756	7924868	Christmas Creek Reef 7
5647	Zn Pb	187226	7931457	Ross Hill 1
5648	Zn Pb	187147	7930861	Ross Hill 2
5650	Zn Pb	187202	7930543	Ross Hill 3
5651	Zn Pb	187369	7930195	Ross Hill 4
5652	Zn Pb	187091	7930965	Ross Hill 5
5653	Zn Pb	186742	7931631	Ross Hill 6
5657	Cu Pb Zn	188618	7938317	North Laidlaw
5659	Pb Zn	194349	7931342	Teichart Hills
5660	Zn Pb	187475	7934456	Laidlaw Range
6472	Zn Pb	184323	7926076	Kapok
6473	Pb	184107	7927155	Pinnacles Lead
6483	Cu	492026	8095245	Mount Elder


	Stratabound sedimentary – clastic-hosted			
3	Zn Pb Ag	394311	8085950	Keller Creek East
4	Zn	385187	8068916	Spring Four
2508	Cu	440134	8093958	Paradise Creek (Cu A)
4543	Cu	349426	8227607	Menuairs Dome 3
5343	Cu (Zn)	368549	8022500	Little Pantan (Cu)
5371	Cu	383454	8047367	MI-10
5401	Zn	227245	7942765	Mount Ramsay 20
5519	Zn	422503	8092749	Osmond Valley (Zn 1)
5520	Zn	422172	8092484	Osmond Valley (Zn 2)
5521	Zn	424196	8093820	Osmond Valley (Zn 3)
5524	Zn	426115	8094773	Osmond Valley (Zn 4)
5527	Cu	428933	8097207	Osmond Valley (Cu 4)
5528	Zn	428363	8096228	Osmond Valley (Zn 5)
5529	Zn	428985	8096678	Osmond Valley (Zn 6)
5531	Ag Pb Zn	429935	8098717	Osmond Valley (Ag-Pb-Zn)
5532	Zn	432028	8098503	Osmond Valley (Zn 8)
5533	Zn	434237	8098252	Osmond Valley (Zn 9)
5534	Zn	435481	8098609	Osmond Valley (Zn 10)
5535	Cu Ag	436063	8098768	Osmond Valley (Cu-Ag)

	Stratabound sedimentary – undivided			
1059	Cu Zn Pb Ag (Sn)	401373	8061760	Wills Creek A
2506	Cu	456083	8152058	Discovery Yard
5307	Cu PGE (Ni)	411037	8084330	Sally Malay Bore 1
5308	Zn	391558	8085646	Mabel Downs 1
5309	Cu	392525	8080751	Mabel Downs 2
5537	Zn Pb	426737	8093661	Osmond Valley (Zn-Pb)
5538	Cu	417913	8095474	Unnamed (Cu)

	Vein and hydrothermal – undivided			
11	Cu (Ag Zn)	378791	8043735	Mini M (Cu) 2
794	Cu	250332	7930308	Eastman (Cu) B
797	Cu	248933	7930038	Eastman (Cu) D
1070	Cu Zn	402183	8060608	Wills Creek B3
1104	Pb Fl	388103	8008194	Ding Dong Downs (Pb-F)
1210	Cu	314003	7914686	McLintock Range
1211	Cu	400108	7994599	Mount Kinahan A
1212	Pb	249430	7931749	Bullock Bore 7
1213	Cu	249983	7930938	Bullock Bore 2
1216	Cu Zn Pb Fl	314532	7944558	Lamboo No. 1
1217	Cu Ag (Au)	279974	7936549	Margaret River 1
1218	Cu Ag	280733	7937658	Margaret River 2
1219	Cu Ag Au	280847	7937317	Margaret River 3
1221	Pb Ag Brl Fl	464521	8198924	Boxers
1222	Cu	411476	8074409	Calico Spring 1
1240	Cu	406891	8077361	Corkwood West
1241	Cu	445912	8163855	Excelsior Range
1242	Cu	398858	8090539	Frogghollow Spring
1243	Cu	417138	8097068	Mabel Downs East

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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BASE METAL

	Vein and hydrothermal – undivided			
1244	Cu Pb (Zn Ni)	354483	7987008	Moola Bulla
1245	Cu Zn Ag	374932	8011596	Palm Spring 1
1247	Cu	373833	8039158	Roses Bore
1249	Pb	358333	8019404	Springvale (Pb)
1250	Pb Ag	354933	8014658	Three Sisters
1251	Cu	394486	8079452	Tickalara Bore
1256	Pb Ag Fl	417133	8128958	Turkey Creek (Pb-Ag-F)
1257	Pb Ag Cu Zn (Au)	286537	7964929	Black and Glidden
1259	Pb Ag Zn (Au Cu)	280925	7974144	Dead Horse Creek
1262	Cu Pb Ag	276283	7967958	Dead Horse Bore Yard
1264	Pb Ag	256418	7956748	Hells Gate Southwest
1265	Pb Ag Zn	257065	7955676	Hells Gate Southeast
1268	Cu Pb	257226	7960747	Hells Gate North
1272	Cu Pb Zn Ag Au	248362	7956018	Me No Savvy A
1274	Cu Au Ag Pb	248816	7955305	Me No Savvy B
1277	Cu Ag	249484	7955088	Me No Savvy D
1327	Cu	237932	7953058	Taylor River
1329	Zn Cu Pb Ag	240632	7942958	Lilyhole Bore
1332	Cu	419633	8168158	Camel Creek Yard
1333	Cu	351083	8075858	Mount Remarkable 1
1334	Cu Pb	359033	8077908	Mount Remarkable 2
1335	Cu Pb	359633	8081308	Mount Remarkable 3
1337	Cu Pb	348682	8086458	Mount Remarkable 4
1340	Cu Pb Ag (Mo Au Sn)	363983	8114408	Base Metal Prospect
1341	Cu	405831	8182155	O'Donnell Range
1344	Cu	229432	7951958	Palm Spring 2
1378	Zn Cu Pb Ag	367346	8060623	Snake
1380	Cu (Ag Pb)	288689	7981410	Mount Amherst 2
1381	Pb Cu Ag	299455	7992388	Blackfellow Creek
1497	Cu Pb	391223	8210658	Speewah 1
1500	Pb Cu Ag Au Sb	391436	8206501	Martins
1501	Cu Pb Ag	389932	8204278	Calamondah
1503	Cu	387973	8194608	Speewah IV
1504	Cu Pb	386193	8191598	Speewah V
1505	Cu Pb Ag	385132	8185478	Speewah VI
1506	Cu	427733	8094458	Osmond Creek
1507	Cu	424037	8094604	Wallys
1508	Cu	429233	8098058	Mount Parker
1509	Pb	420433	8096058	Palms Yard
2209	Cu	423537	8094604	Minatome
2330	Cu Au Ag	306327	7911516	Horrays Copper Show
2331	Cu Au Ag	311982	7919158	Hanging Drum Copper Show
2346	Pb	409466	8180382	Jailhouse Creek 2
2520	Pb Ag	467624	8206438	Carr Boyd 1
2521	Pb Cu Ag	436134	8099558	Osmond Creek (Pb)
2530	Cu	326694	7974590	Ten Minute (Cu 2)
2720	Cu	414002	8070631	Frank River
2721	Cu	312282	7931358	Hangman Creek
2722	Cu	389844	8011500	Mount Coghlan
2723	Cu	326767	7959392	New Lamboo
2725	Cu	228711	7952281	Pond Spring
2758	Pb Cu	364133	8114458	Middletons
2870	Cu (Au)	344633	7999558	Five Mile A
2924	Cu	406242	8061708	Ord (Cu)
2966	Cu	350468	8210427	Campbellmerry
3155	Pb Zn Ag	390478	8028568	Alice Downs 2
3384	Cu	415033	8077858	Frank River 2
3385	Cu	420683	8082008	Frank River 3
4299	Cu	245106	7928921	Bullock Bore 3
4301	Cu Ag	244376	7929191	Bullock Bore 4
4641	Cu	353682	8073658	Mount Remarkable 9
4643	Cu	361032	8091158	Mount Remarkable 10
4644	Cu	360232	8094358	Mount Remarkable 11
4645	Cu	353233	8064958	Mount Remarkable 12
4799	Cu	281220	7975045	Dead Horse 2
4800	Cu	279034	7972689	Dead Horse 3
4801	Cu	277571	7971648	Dead Horse 4
4803	Cu	275525	7968561	Dead Horse 6
4822	Cu	377217	7997620	Sophie Downs (Cu) 1
4830	Cu	337845	7958040	Mount Angelo South
4832	Pb	466896	8209666	Carr Boyd 2
4834	Pb Ag	467161	8206279	Carr Boyd 3
4838	Pb Ag	467545	8207272	Carr Boyd 4
4842	Pb Ag (Zn)	467703	8207153	Carr Boyd 5
4846	Pb Zn Ag	467305	8207576	Carr Boyd 6
4847	Pb Zn Ag	467293	8208767	Carr Boyd 7
4864	Cu Zn Ag	377456	7999769	Gossan Y

NUMBER COMMODITY EASTING NORTHING NAME

BASE METAL



Vein and hydrothermal – undivided

4915	Zn	383947	7998948	Sophie Downs (Zn) 2
4916	Cu	384807	8000468	Sophie Downs (Cu) 2
4917	Cu	384317	7997981	Sophie Downs (Cu) 3
4919	Cu	384066	7994396	Sophie Downs (Cu) 4
4920	Cu	384053	7994238	Sophie Downs (Cu) 5
4921	Cu	384039	7993695	Sophie Downs (Cu) 6
4922	Cu	383973	7993219	Sophie Downs (Cu) 7
4930	Cu Pb	438808	8143634	Mount Pitt A
4931	Cu Pb	438404	8143634	Mount Pitt B
4932	Cu Pb	439792	8142662	Mount Pitt C
4933	Cu	439691	8142826	Mount Pitt D
4936	Cu	439678	8143445	Mount Pitt F
4937	Cu Pb	439537	8144872	Mount Pitt E
4938	Cu Pb	439352	8145163	Mount Pitt H
4951	Pb	410888	8129704	Bow River (Pb)
4958	Cu	426406	8148000	Salt Lick Creek (Cu)
4961	Pb Cu	249391	7931603	Bullock Bore 9
4965	Cu	248306	7930386	Louisa Downs 1
4966	Zn Cu Pb Co	251217	7931021	Louisa Downs 2
5053	Cu Ag	395307	8000483	Castle Creek 17
5054	Zn	319835	7917769	McClintock Range 1
5056	Cu Zn (Ni)	320270	7915639	McClintock Range 2
5058	Cu (Pb Zn)	318948	7915639	McClintock Range 3
5063	Cu (Zn)	318446	7915136	McClintock Range 4
5067	Cu (Pb Zn)	317493	7913694	McClintock Range 5
5073	Cu Ag	368713	7963339	Ruby Queen - Brockman 1
5075	Pb Zn	329904	7940638	Garden Creek 1
5076	Zn	330433	7941287	Garden Creek 2
5078	Cu Ag	400253	7994678	Mount Kinahan B
5083	Cu Ag	400637	7994824	Mount Kinahan C
5084	Cu	406060	7996958	Mount Kinahan D
5094	Zn	378037	8004571	Gossan X
5097	Cu	245980	7955458	Me No Savvy 1
5098	Cu	246178	7955431	Me No Savvy 2
5099	Cu Au Ag	245763	7954976	Me No Savvy 3
5100	Cu	391086	7999097	Saunders Creek Anticline 1
5101	Cu	391536	8000552	Saunders Creek Anticline 2
5102	Cu	391258	8000076	Saunders Creek Anticline 3
5190	Cu Ni	352387	8034748	Springvale 11
5205	Cu (Au)	395782	8077065	Tickalara Bore 7
5229	Pb	355037	7982775	Halls Creek 1
5230	Pb	354733	7981570	Halls Creek 2
5233	Cu Zn Ag	313604	7920922	Mount Dockrell 6
5234	Cu	312532	7918528	Mount Dockrell 7
5235	Ag Cu	305058	7916146	Black Bishop Hill A
5237	Zn Ag	324994	7920922	Hydrothermal Lens A
5238	Cu	324306	7920578	Hydrothermal Lens B
5240	Cu	400733	8065758	Dougals-Ord River
5248	Cu	392998	8198692	Speewah A
5249	Cu	392682	8199208	Speewah B
5250	Cu Brl Fl	482350	8048138	Brook Creek
5251	Cu	468572	8198631	Mons Mont
5252	Cu	361799	8041527	Springvale 13
5257	Cu Pb Zn Ag	321672	7960886	Emull East
5267	Cu	366727	8049651	Billymac Yard C
5271	Pb Zn Cu Ag Au	293878	7967653	Mount Amherst 1
5278	Cu Ag	246655	7955544	Me No Savvy 4
5279	Cu (Au)	245953	7955795	Me No Savvy 5
5280	Cu	247654	7954993	Me No Savvy 6
5281	Cu Ag	247832	7955147	Me No Savvy 7
5282	Cu Ag	251429	7955676	Me No Savvy 8
5294	Cu	289510	7981304	Mount Amherst 3
5295	Pb Cu	288240	7981093	Mount Amherst 4
5296	Cu	293719	7967680	Mount Amherst 5
5320	Cu	418021	8087630	John Galt (Cu) 1
5321	Cu	417916	8087432	John Galt (Cu) 2
5322	Cu	417545	8087154	John Galt (Cu) 3
5323	Cu	417690	8087048	John Galt (Cu) 4
5324	Cu	418326	8087815	John Galt (Cu) 5
5325	Cu	418431	8088186	John Galt (Cu) 6
5383	Cu Zn Pb	227997	7941993	Mount Ramsay 1
5388	Pb Zn	240098	7939442	Mount Ramsay 15B
5390	Cu	280080	7937072	Margaret River 4
5391	Cu	285687	7941538	Margaret River 5
5392	Cu	229867	7950237	Mount Ramsay 7
5393	Cu	253574	7954602	Mount Ramsay 10
5395	Cu	246959	7946664	Mount Ramsay 9
5396	Pb Zn	256432	7954337	Mount Ramsay 11

NUMBER COMMODITY EASTING NORTHING NAME

BASE METAL



Vein and hydrothermal – undivided

5397	Cu	254130	7953147	Mount Ramsay 13
5402	Pb Zn	296487	7965309	Mount Amherst 6
5541	Cu	411518	8074198	Calico Spring 2
5542	Cu	411275	8074388	Calico Spring 3
5543	Cu	411254	8074102	Calico Spring 4
5544	Cu	410894	8074039	Calico Spring 5
5545	Cu	411910	8074748	Calico Spring 6
5546	Cu	410862	8073552	Calico Spring 7
5548	Cu	410598	8073213	Calico Spring 8
5550	Cu	410111	8072949	Calico Spring 9
5551	Cu	410418	8072949	Calico Spring 10
5552	Cu	410460	8072854	Calico Spring 11
5553	Cu	410926	8072324	Calico Spring 12
5554	Cu	411392	8071679	Calico Spring 13
5556	Cu	410873	8070811	Calico Spring 14
5557	Cu	410460	8070991	Calico Spring 15
5558	Cu	409540	8071012	Calico Spring 16
5559	Cu	411582	8075203	Calico Spring 17
5560	Cu	411571	8075425	Calico Spring 18
5561	Cu	411878	8076071	Calico Spring 19
5599	Cu Ag	420335	8076157	Frank River 4
5600	Cu Ni	410535	8067957	Mount Ranford
5693	Pb	249755	7931673	Bullock Bore 8
5702	Cu (Pb)	243870	7928133	Eastmans No. 2
5720	Cu	437417	7977159	Errol Bore
5725	Cu	346273	7972569	Koongie Park 2
5727	Cu	375076	7998824	Duffer Range 2
5742	Cu Ag	323472	7910653	Surprise
5758	Cu	369046	7968462	Butcher Creek
5770	Cu	280847	7936311	Margaret River 6
5957	Cu (Au)	248608	7930016	Eastman Bore (Cu 1)
5958	Cu	248441	7929961	Eastman Bore (Cu 2)
6093	Cu	303362	7911302	Mount Carmel East 1
6212	Cu (Au)	249219	7930836	Eastman Bore (Cu 3)
6333	Pb Cu Zn Ag	341082	7993021	Moola Bulla 5
6355	Pb Zn Cu (As Au)	362122	7991010	Elvire Gossan
6362	Cu Pb Zn Ag	252955	7959010	Hells Gate 1
6363	Pb Zn Ag	253320	7959383	Hells Gate 2
6365	Pb (Au)	364678	7996904	Tiger 1
6372	Pb Cu Zn Ag Au	298337	7968128	Eastern Shear 1
6375	Pb Zn (Au Cu Ag)	281031	7973814	Dead Horse Creek 2
6388	Pb Cu Zn	228121	7932805	Bohemia
6389	Pb Zn	215409	7948076	Mount Fairbairn 1
6399	Cu Au	206901	7970784	Cabbage Tree
6405	Cu	246297	7952071	Glidden 3
6487	Cu	445022	8162377	Mount Evelyn Range
6488	Cu	357609	8188194	Plants
6493	Cu	397726	8174591	Moonlight Valley Yard
Regolith – residual to eluvial placers				
790	Cu	245184	7928242	Eastman (Cu) A
2756	Cu	479368	8059458	Copper Flats
Regolith – residual and supergene				
1128	Cu Ag Au	340032	7968758	Sandiego (oxide)

IRON



Sedimentary – banded iron-formation (supergene enriched)

3295	Fe	421132	8096358	Osmond Valley (Fe)
5344	Fe	421018	8162003	Pompeys Pillar - Sam
5345	Fe	421049	8161145	Pompeys Pillar - Central
5346	Fe	421383	8160180	Pompeys Pillar - Tony
5348	Fe	428019	8155789	Matsu
5349	Fe	421716	8160740	Pompeys Pillar - Peter
5419	Fe	490540	7871090	Black Hill

ENERGY MINERAL



Stratabound volcanic and sedimentary – undivided

6262	U (Cu)	497953	7872277	Mount Junction
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Stratabound – sandstone-hosted U

2726	U V	436239	8184154	Horse 1
5340	U	437337	8183545	Horse 2
5342	U Cu	437575	8182248	Horse 10





Stratabound sedimentary – clastic-hosted


4638	U	457133	8204358	Revolver Creek
5341	U	437456	8182804	Horse 7; DH1
5403	U REE	497223	7813617	Killi Killi 2
5797	U (Cu Zn)	386050	8003006	Rustlers

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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ENERGY MINERAL


	Vein and hydrothermal – undivided			
2348	U	414512	8181006	Frog 1
2370	U FI	412885	8181291	Frog 4
2371	U FI	414554	8180855	Frog 6
2377	U	417233	8190458	Dunham River A
2378	U	416549	8189548	Dunham River B
2380	U FI P	367088	8117406	Antares
2528	U (Pb V Cu)	326520	7974812	Ten Minute (U)
4656	U	419533	8186158	Dunham River C
4657	U	418833	8174858	Dunham River D
4804	U	416976	8190610	Donkey Creek
5329	U	343342	8051763	Anomaly A3
5330	U	414345	8180815	Frog 2
5331	U	414250	8181225	Frog 3
5332	U FI	412877	8180942	Frog 5
5333	U FI	412702	8180484	Frog 7
5335	U	410933	8183258	Last
5350	U	417373	8087346	John Galt (U)


	Vein and hydrothermal – unconformity			
5793	U Au (REE)	486462	7870419	Don
5795	U	490098	7844638	Birindudu 2
5805	U (Au Xen)	451047	7890625	Mount Mansbridge 1
5806	U	442861	7886593	Mount Mansbridge 2
5807	U	473312	7868745	Oracle
5808	U	454548	7872015	Jaimani


	Regolith – residual and supergene			
2727	U Pb	456268	8210506	Spinifex
6194	U	320175	7908656	A1


INDUSTRIAL MINERAL

	Carbonatite and alkaline igneous intrusions			
2399	FI Brt Ag	386649	8183213	West Ridge
	Orthomagmatic mafic and ultramafic – layered mafic			
5103	Tlc	322045	7956191	Lamboo (talc)
5168	Asbc	322997	7957936	Lamboo (asbestos 1)
5201	Asbc	385973	8065951	Ord Crossing (asbestos)
5486	Asbc	324254	7957639	Lamboo (asbestos 2)
5487	Asbc	323962	7957123	Lamboo (asbestos 3)
5488	Asbc	322217	7956898	Lamboo (asbestos 4)


	Skarn			
5613	Grt	384415	8055317	Garnet Hills

	Vein and hydrothermal – undivided			
1338	FI	358333	8079808	Mount Remarkable 5
1339	FI	351783	8071458	Mount Remarkable 6
2397	FI Cu (Au)	391041	8186397	A Vein
2398	FI Pb Cu	386902	8183407	West Vein
2405	FI Brt	386183	8186158	West Ridge North
2420	FI	386283	8197658	North Shear
2423	Brt	384583	8193858	Main Shear North
2428	FI	385933	8190058	Martins (fluorite)
5334	FI	412041	8177799	Frog 8
5351	FI	391161	8186497	A Vein East
5352	FI	390682	8185303	B Vein
5353	FI	391201	8186245	North B Vein
5354	FI	390950	8186298	C Vein
5355	FI	391015	8186298	C Vein East Branch
5356	FI	391518	8186495	Cross Vein
5357	FI	390420	8183135	D Vein
5362	FI	391227	8185516	G Vein
5365	FI	389388	8181600	South Vein
5366	FI	390975	8185265	East Vein
5367	FI	389758	8182024	F Vein
5379	FI	388309	8008930	Ding Dong Downs (F)
6484	Asbc	480409	8118247	J41 Trig
6494	Brt	427970	8129540	Mabel Downs (Ba)

	Stratabound volcanic and sedimentary – undivided			
2391	FI	387742	8161205	Archie Creek



	Regolith – residual and supergene			
5163	Mgs	322811	7955965	Lamboo (magnetite)

CONSTRUCTION MATERIAL

	Orthomagmatic mafic and ultramafic – layered mafic			
2765	Bsbg	360258	8033738	McIntosh B

NUMBER	COMMODITY	EASTING	NORTHING	NAME
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CONSTRUCTION MATERIAL

	Orthomagmatic mafic and ultramafic – undivided			
6474	Bsdo	361092	7988784	Halls Creek North
	Stratabound sedimentary – clastic-hosted			
6471	Sst	270947	7925740	Berrangi Creek

Appendix 2

WAMIN and EXACT databases

WAMIN database (mineral occurrences)

The WAMIN (Western Australian mineral occurrence) database of the Geological Survey of Western Australia (GSWA) contains geoscience attribute information on mineral occurrences in Western Australia. The database includes textual and numeric information on the location of the occurrences, location accuracy, mineral commodities, mineralization-style classification, order of magnitude of resource tonnage and estimated grade, ore and gangue mineralogy, details of host rocks, and both published and unpublished references. Each of the occurrences in WAMIN is identified by a unique 'deposit number'.

The WAMIN database uses a number of authority tables to constrain the essential elements of a mineral occurrence, such as the operating status, the commodity group, and the style of mineralization. In addition, there are parameters that dictate whether the presence of a mineral or an analysed element is sufficiently high to rank occurrence status; this Report only deals with mineral occurrences. These and other attributes were extracted either from open-file mineral exploration reports in WAMEX (Western Australian mineral exploration database) or from the published literature.

Those elements of the database that were used to create the symbols for mineral occurrences and tabular information displayed in Plates 1 and 1a, and Appendix 1 of this Report, are:

- occurrence number and name (deposit number and name);
- operating status (font style of deposit number);
- position and spatial accuracy (symbol position);
- commodity group (symbol colour); and
- mineralization style (symbol shape).

The elements of the database used for symbology in Plates 1 and 1a and Appendix 1 are operating status, commodity group, and mineralization style. These parameters have previously been defined for the GSWA mineralization mapping projects that have been completed for prospectivity enhancement studies of southwest Western Australia (Hassan, 1998), the north Eastern Goldfields (Ferguson, 1998), the Bangemall Basin (Cooper et al., 1998), and the west Pilbara (Ruddock, 1999).

Operating status

The database includes mineralization sites (referred to as deposits) ranging from small, but mineralogically significant, mineral occurrences up to operating mines. The classification includes all MINEDEX sites with established resources: MINEDEX is the Department of Minerals and Energy (DME) mines and mineral deposits information database (Townsend et al., 1996, 2000). All occurrences in the WAMIN database are assigned a unique, system-generated number (deposit number). The font style of this number (bold, italicized, and so on) is used as the coding to indicate operating status both on the face of the map and in Appendix 1 of this Report. The system used is:











- Mineral occurrence — any economic mineral exceeding an agreed concentration and size found in bedrock or regolith (italic serif numbers, e.g. *1212*).
- Prospect — any working or exploration activity area that has found subeconomic mineral occurrences, and from which there is no recorded production (italic serif numbers, e.g. *1138*).
- Mineral deposit — economic minerals for which there is an established resource figure (serif numbers, e.g. 1137).
- Abandoned mine — workings that are no longer operating, or are not on a care-and-maintenance basis, and for which there is recorded production, or where field evidence suggests that the workings were for more than prospecting purposes (bold-italic sans serif numbers, e.g. **2321**).
- Operating mine — workings that are operating, including on a care-and-maintenance basis, or that are in development leading to production (bold sans serif numbers, e.g. **1106**).

The names of the occurrences, and any synonyms that may have been used, are derived from the published literature and from open-file reports (in WAMEX). Names that appear in the MINEDEX database have been used where possible, although there may be differences created because MINEDEX uses site names based on overall production and resources, whereas WAMIN may show names of individual occurrences at a MINEDEX site.

Commodity group

The WAMIN database includes a broad grouping that is based on the potential end-use or typical end-use of the

Table 2.1. WAMIN authority table for commodity groups

<i>Commodity group</i>	<i>Typical commodities</i>	<i>Symbol colour</i>
Precious mineral	Diamond, semi-precious gemstones	
Precious metal	Ag, Au, PGE	
Steel-industry metal	Co, Cr, Mn, Mo, Nb, Ni, V, W	
Speciality metal	Li, REE, Sn, Ta, Ti, Zr	
Base metal	Cu, Pb, Zn, Sb	
Iron	Fe	
Aluminium	Al (bauxite)	
Energy mineral	Coal, U	
Industrial mineral	Asbestos, barite, kaolin, talc	
Construction material	Clay, dimension stone, limestone	

principal commodities comprising a mineral occurrence. The commodity group, as listed in Table 2.1, determines the particular colour for the mineral occurrence symbols in Plates 1 and 1a, and Appendix 1.

The commodity groupings are based on those published by the Mining Journal (1998) with modifications, as shown in Table 2.2, to suit the range of minerals and end-uses for the mineral output of Western Australia.

Mineralization style

There are a number of detailed schemes for classifying mineral occurrences into groups representing different

styles of mineralization, with the scheme of Cox and Singer (1986) probably being the most widely used. The application of this scheme in Western Australia would necessitate modifications to an already complex scheme, along the lines of those adopted by the Geological Survey of British Columbia (Lefebure and Ray, 1995; Lefebure and Hoy, 1996). Representing the style of mineralization on the face of a map cannot be simply and effectively achieved if the scheme adopted is too complex.

The Geological Survey of Western Australia has adopted the principles of ore deposit classification from Evans (1987) with some modifications based on Edwards and Atkinson (1986). This scheme works on the premise that 'If a classification is to be of any value it must be

Table 2.2. Modifications made to the Mining Journal Ltd (1998) commodity classification

<i>Commodity group (Mining Journal Ltd, 1998)</i>	<i>Commodities</i>	<i>Changes made for WAMIN commodity group</i>
Precious metals and minerals	Au, Ag, PGE, diamonds, other gemstones	Diamond and other gemstones in precious minerals group; Au, Ag, and PGE in precious metals group
Steel-industry metals	Iron ore, steel, ferro-alloys, Ni, Co, Mn, Cr, Mo, W, Nb, V	Fe in iron group
Speciality metals	Ti, Mg, Be, REE, Zr, Hf, Li, Ta, Rh, Bi, In, Cd, Sb, Hg	Sn added from major metals; Sb into the base metals group
Major metals	Cu, Al, Zn, Pb, Sn	Cu, Pb, and Zn into the base metals group; Al (bauxite) into aluminium group; Sn in speciality metals
Energy minerals	Coal, U	No change
Industrial minerals	Asbestos, sillimanite minerals, phosphate rock, salt, gypsum, soda ash, potash, boron, sulfur, graphite, barite, fluorspar, vermiculite, perlite, magnesite/magnesia, industrial diamonds, kaolin	No change

Table 2.3. WAMIN authority table for mineralization styles and groups

<i>Mineralization style</i>	<i>Typical commodities</i>	<i>Group symbol^(a)</i>
Carbonatite and alkaline igneous intrusions Kimberlite and lamproite intrusions	Nb, Zr, REE, P Diamond	☆
Disseminated and stockwork in plutonic intrusions Greisen Pegmatitic Skarn	Cu, Mo, Au Sn Sn, Ta, Nb, Li W, Mo, Cu, Pb, Zn, Sn	⬡
Orthomagmatic mafic and ultramafic — komatiitic or dunitic Orthomagmatic mafic and ultramafic — layered-mafic intrusions Orthomagmatic mafic and ultramafic — undivided	Ni, Cu, Co, PGE Ni, Cu, Co, V, Ti, PGE, Cr Ni, Cu, Co, V, Ti, PGE, Cr	⊕
Vein and hydrothermal — undivided Vein and hydrothermal — unconformity	Au, Ag, Cu, Pb, Zn, Ni, U, Sn, F U	◇
Stratabound volcanic and sedimentary — volcanic-hosted sulfide Stratabound volcanic and sedimentary — sedimentary-hosted sulfide Stratabound volcanic and sedimentary — volcanic oxide Stratabound volcanic and sedimentary — undivided	Cu, Zn, Pb, Ag, Au Pb, Zn, Cu, Ag Fe, P, Cu Pb, Zn, Cu, Ag, Au, Fe	△
Stratabound sedimentary — carbonate-hosted Stratabound sedimentary — clastic-hosted Stratabound sedimentary — undivided Sedimentary — banded iron-formation (supergene enriched) Sedimentary — banded iron-formation (taconite) Sedimentary — undivided	Pb, Zn, Ag, Cd Pb, Zn, Cu, Au, Ag, Ba, Cd, U Pb, Ba, Cu, Au Fe Fe Mn	□
Sedimentary — basin	Coal, bitumen	○
Regolith — alluvial to beach placers Regolith — calcrete Regolith — residual and supergene Regolith — residual to eluvial placers	Au, Fe pisolites, Ti, Zr, REE, diamond, Sn U, V Al, Au, Ni, Co, Mn, V, Fe scree Au, Sn, Ti, Zr, REE, diamond	▭
Undivided	Various	▽

NOTE: (a) The white symbol colour used in this table does not indicate the commodity group in Table 2.1

capable of including all known ore deposits so that it will provide a framework and a terminology for discussion and so be of use to the mining geologist, the prospector and the exploration geologist'. The system below is based on an environmental-rock association classification, with elements of genesis and morphology where they serve to make the system simpler and easier to apply and understand (Table 2.3).

To fully symbolize all the mineralization style groups would result in a system that is too complex. As the full details of the classification are preserved in the underlying WAMIN database, the chosen symbology has been reduced to nine shapes (Table 2.3).

Mineral occurrence determination limits

Any surface expression of mineralization (gossan or identified economic mineral) is an occurrence. Subsurface or placer mineralization is included as an occurrence where it meets the criteria given in Table 2.4.

Professional judgement is used if shorter intercepts or surface occurrences at higher grade (or vice versa) are

involved. Any diamonds or gemstones would be mineral occurrences, including diamondiferous kimberlite or lamproite.

EXACT database (exploration activities)

The EXACT* database is a GIS-based spatial index, for exploration activities in WAMEX, which has been developed by the GSWA to improve access to information in open-file mineral exploration reports (Ferguson, 1995). A major limitation to data retrieval in WAMEX, in its current form, is the difficulty in selecting reports that cover a specific area and, further, in precisely locating various individual exploration activities described within a selected report.

In the current WAMEX database, when spatial parameters are used to make data searches, the results of

* The EXACT database is a GIS-based spatial index of EXploration ACTivities. This term supersedes the acronym SPINDEX (Spatial Index) used in Cooper et al. (1998), Ferguson (1998), and Hassan (1998).

Table 2.4. Suggested minimum intersections for mineral occurrences in drillholes or trenches

<i>Element</i>	<i>Intersection length (m)</i>	<i>Grade</i>
Hard rock and lateritic deposits		
Gold	>5	>1 ppm
Silver	>10	>1 ppm
Platinum	>0.5	>1 ppm
Lead	>5	>0.5%
Zinc	>5	>2%
Copper	>5	>0.5%
Nickel	>5	>0.5%
Cobalt	>5	>0.1%
Chromium	>0.2	>5% Cr ₂ O ₃
Vanadium	>5	>0.1%
Tin	>5	>0.02%
Iron	>5	>40% Fe
Manganese	>5	>25%
Uranium	>5	>1000 ppm U
Diamonds	na	any diamonds
Tantalum	>5	>200 ppm
Tungsten	>5	>1000 ppm (0.1%)
Placer deposits		
Gold	na	>300 mg/m ³ in bulk sample
Diamonds	na	any diamonds
Heavy minerals	>5	>2% ilmenite

searches are constrained to very large areas. The smallest search polygon that can be effectively used to locate reports in WAMEX is the area of a 1:100 000-scale sheet. Even though a query may be entered as a single point (either AMG or latitude/longitude coordinates), the resulting search will produce all reports for the 1:100 000-scale sheet in which that single point is located. Hence, for example, it is not possible to restrict report selection to small areas of prospective ground of particular interest to the user. As a consequence these WAMEX searches are time consuming, and they have become more time consuming as the number of open-file reports has increased with continuing releases of data.

The EXACT spatial index overcomes this problem and allows easy access to data on specific areas of previous exploration activity. It also provides a spatial representation of the intensity of past exploration, thereby highlighting prospective areas that may have been lightly or inadequately tested by various earlier exploration methods.

The spatial index consists of an attribute database, developed in Microsoft Access, which is linked to ArcView for spatial representation. In the CD-ROM, the dataset includes tabulated textual and numeric information that has been retrieved from open-file mineral exploration reports and attached to individual exploration activities. The areas of exploration activity are digitized (as polygons, lines, or points) using the computer-assisted drafting (CAD) system Microstation, converted into Arc/Info, and then transferred into ArcView to enable an interactive display of EXACT. The positional data are digitized from hard-copy maps and plans in mineral exploration reports, using various

Table 2.5 Types of exploration activity detailed in the EXACT database

<i>Activity type</i>	<i>Description</i>
Geological	
GEOL	Geological mapping
AMS	Airborne multispectral scanning
LSAT	Landsat TM data
Geophysical	
AEM	Airborne electromagnetic surveys
AGRA	Airborne gravity surveys
AMAG	Airborne magnetic surveys
ARAD	Airborne radiometric surveys
MAG	Magnetic surveys
EM	Electromagnetic surveys (includes TEM, SIROTEM)
GEOP	Other geophysical surveys (includes IP, resistivity)
GRAV	Gravity surveys
RAD	Radiometric surveys (includes downhole logging)
SEIS	Seismic surveys
Geochemical	
SOIL	Soil surveys
SSED	Stream-sediment surveys
REGO	Regolith surveys (includes laterite, pisolite, ironstone, and lag)
NGRD	Non-gridded geochemical surveys (includes chip, channel, dump, and gossan)
ACH	Airborne geochemistry
Mineralogical	
HM	Heavy mineral surveys
Drilling	
DIAM	Diamond drilling
ROT	Rotary drilling (predominantly percussion drilling)
RAB	RAB drilling (includes other shallow geochemical drilling such as auger)
RC	RC drilling
Mineral resources	
MRE	Mineral resource estimate
Hydrogeological	
HYDR	Groundwater surveys

published sources (geological maps, topographic maps, and TENGRAPH—DME's electronic tenement-graphics system) for georeference purposes. The types of exploration activity detailed are essentially those used in WAMEX, with some rationalization, and these are listed in Table 2.5. In the table, the 25 activities are grouped as follows:

- Geological activities (and remote sensing activities)
- Geophysical activities
- Geochemical activities
- Mineralogical activities
- Drilling activities
- Mineral resources
- Hydrogeological activities.

The above groups relate to those specified in the statutory guidelines for mineral exploration reports (Department of Minerals and Energy, 1995).

For each separate exploration activity the following statistics have been compiled:

- description of activity
- sample types and numbers
- elements analyzed (asterisk symbol (*) against elements for a rough guide to anomalism)
- metres of drilling and number of holes
- scales of presentation of data in reports.

The activity data are also linked in the dataset to the following related information taken from WAMEX:

- A-numbers (WAMEX accession numbers for individual reports)
- I-numbers (WAMEX item numbers for single or groups of reports on microfiche)
- company or companies that submitted reports
- period of exploration (years)
- mineral commodities sought
- summaries (annotations) of exploration projects included in individual item numbers.

In ArcView, the exploration activities are included as spatial **themes**, which are displayed as polygons, lines, or points on the interactive on-screen map known as the **view**. The **table of contents** (i.e. map legend) provided alongside the **view** allows access to the **themes**, so that any **theme** or combination of **themes** may be displayed. Details (taken from attribute tables) of any **theme** can be accessed on screen, and **queries** can be carried out either as spatial queries through a **view** or as textual queries direct from the attribute tables. Further details (with examples) of displays, queries, charts, and view layouts are provided by Ferguson (1995).

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Appendix 3

Description of digital datasets on CD-ROM

There are three principal components of this study, which are this Report, Plates 1 and 1a, and a CD-ROM containing digital datasets for use with database or GIS software. The CD-ROM includes all the data used to compile the map and Report, and also includes files of exploration and mining activity, geophysical, remote sensing, and topographic data. The CD-ROM also includes the files necessary for viewing the data in the ArcView GIS environment, and a self-loading version of the ArcExplorer software package modified to suit this particular dataset.

Mineral occurrences (WAMIN)

The mineral occurrence dataset (from WAMIN, the Western Australian mineral occurrence database) as used in this Report and on Plates 1 and 1a is described in Appendix 2. The dataset on the CD-ROM includes textual and numeric information on:

- location of the occurrences (AMG coordinates, latitude and longitude, geological province, location method, and accuracy);
- commodities and commodity group;
- mineralization classification and morphology;
- order of magnitude of resource tonnage and estimated grade;
- mineralogy of ore and gangue;
- details of host rocks;
- both published and unpublished references.

EXACT

The EXACT dataset (from EXACT, Geological Survey of Western Australia's spatial index of exploration activities) as used in this Report is described in Appendix 2. The dataset on CD-ROM contains spatial and textual information (derived from WAMEX open-file reports) defining the locations and descriptions of exploration activities in the area. EXACT, for the east Kimberley area, was compiled between 1996 and 1998, and contains information on types of mineral exploration activity such as statistics relating to:

- report numbers
- sample types and numbers
- elements assayed
- metres of drilling and number of holes
- scales of presentation of the data.

Positional data were taken from hard-copy maps of various scales, from company reports (in WAMEX), located from coordinate and/or geographical information (from topographic maps or Landsat images), and then digitized. Table 2.5 (in Appendix 2) lists the exploration activity types.

The activity data are linked to more general data concerning the individual open-file reports (commonly defined in WAMEX by accession A-numbers) and individual exploration projects (commonly defined in WAMEX by open-file Item-numbers). This information includes the company or companies involved in the project, the commodities explored for, the timing of the project, names of localities in the project, and a summary (annotation) of the project, including exploration concept, activities, and a synopsis of results.

WAMEX

All relevant open-file company mineral exploration reports for the area, indexed in the Department of Minerals and Energy (DME) WAMEX database were referred to for this study. Information extracted from these reports was used to analyse the historical trends in exploration activity and target commodities.

MINEDEX

The MINEDEX database (Townsend et al., 1996, 2000) has current information on all mines, process plants, and deposits, excluding petroleum and gas, for Western Australia. Mineral resources included in MINEDEX must conform to the Joint Ore Reserves Committee (JORC) (1999) code to be included in the database. The database contains information relevant to WAMIN under the following general headings:

- commodity group and minerals
- corporate ownership and percentage holding
- site type and stage of development
- location data (a centroid) including map, shire, mining district, and centre
- current mineral resource estimates
- mineralization type
- tectonic unit
- tenement details.

MINEDEX contains all the relevant resource information and WAMIN uses the unique MINEDEX site

number as a cross-reference for this information. WAMIN may contain pre-resource global estimates that do not conform to the JORC (1999) code, and are not included in MINEDEX.

TENGRAPH

The TENGRAPH database (DME's electronic tenement-graphics system) shows the position of mining tenements relative to other land information. TENGRAPH provides information on the type and status of the tenement and the name(s) and address(es) of the tenement holders (Department of Minerals and Energy, 1994). It should be borne in mind that the tenement situation is constantly changing and that current tenement plans should be consulted before making any landuse-based decisions or applying for tenements.

Solid geology and regolith

The solid geology and regolith incorporates an interpretation of the study area, at 1:100 000 scale, based on a recent compilation of the Geological Survey of Western Australia (GSWA) mapping, aeromagnetism (TMI images), and Landsat TM imagery. The full details of the solid geology and regolith are on the CD-ROM. The regolith on Plates 1 and 1a is a simplified version of the digital dataset on the CD-ROM, and uses two overprints to distinguish relict and depositional regimes. The CD-ROM also includes a large number of solid geology and regolith units that are smaller than 250 000 m² in area that were omitted from Plates 1 and 1a for simplicity.

Geophysics

The aeromagnetic data covering the area are presented in the form of a total magnetic intensity (TMI) colour image. The data used to create the image were flown in 1995 for the National Geoscience Mapping Accord (between Australian Geological Survey Organisation (AGSO) and GSWA), mostly at a line spacing of 400 m, and gridded to a cell size of 800 m for the colour image. More-detailed data, gridded to a cell size of 100 m, may be obtained from AGSO.

Measurements of the background radiation using an airborne crystal usually took place concurrently with the AGSO aeromagnetic surveys over the area. The colour image on the CD-ROM shows the comparative K–Th–U ratios as red–green–blue (RGB). The data are relatively disparate in nature as variations in the crystal size and flying height were not tightly constrained over the area.

A regional gravity survey by AGSO, at a nominal station spacing of 11 km, is presented in the digital dataset as an image showing the Bouguer anomaly, gridded to a cell size of 5 km.

Landsat

Landsat TM imagery has been acquired for all the 1:250 000-scale map sheets in the east Kimberley study. The raw data are available commercially through the Remote Sensing Services section of the Department of Land Administration (DOLA). Images are included in the digital package that preserve the original 25 m pixel size, but these cannot be reverse-engineered back to any bands or band ratios of the original 6-band dataset.

Both image datasets comprise a patchwork of 1:250 000-scale map tiles. The simplest of the two uses a decorrelation stretch of the first principal component of bands 1, 2, 3, 4, 5, and 7, written out as an 8-bit dataset that can be viewed as a monochrome image. The second, more complex, image can be viewed in colour, and was created using a decorrelation stretch of bands 4, 5, and 7.

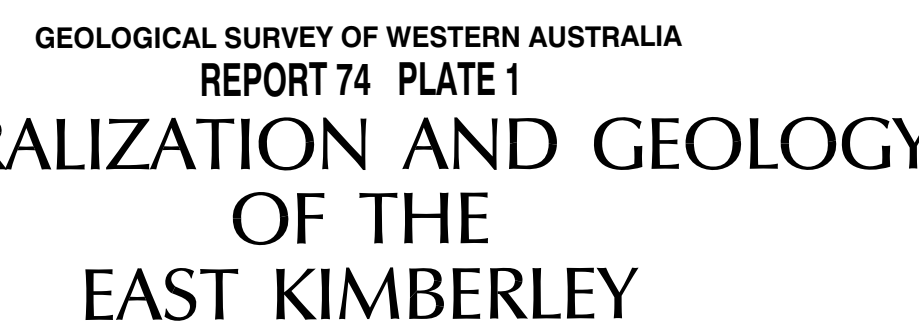
Cultural features

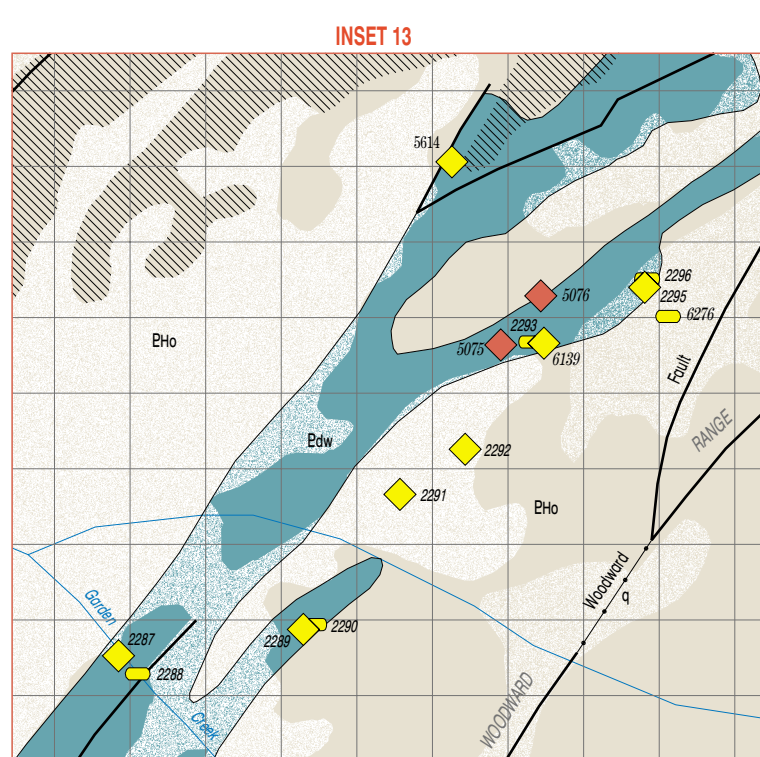
Selected roads and tracks are given as a single dataset, and range from sealed highways through shire roads to major station tracks. The digital data in this file were captured by digitizing from Landsat imagery.

Place names for the area, in a separate file, are given for major hills, stations, and communities. More-comprehensive topographical and cultural data, including drainage, can be obtained from the Australian Land Information Group (AUSLIG).

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